# **The Ensemble of Data Assimilations**

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## Outline

- KF, EKF, Reduced-rank KF, EnKF
- Hybrid methods
- The Ensemble of Data Assimilations (EDA) method
- Applications of the EDA



For a linear system the data assimilation update is:

$$\mathbf{x}_{a}^{k} = \mathbf{x}_{b}^{k} + \mathbf{K}_{k} \left( \mathbf{y}^{k} - \mathbf{H}_{k} \mathbf{x}_{b}^{k} \right)$$
$$\mathbf{x}_{b}^{k+1} = \mathbf{M}_{k} \mathbf{x}_{a}^{k}$$

Under the assumptions of statistically independent background ( $P^b$ ), observation (R) and model errors (Q), the evolution of the system error covariances is given by:

$$\mathbf{P}_{k}^{a} = \left(\mathbf{I} - \mathbf{K}_{k}\mathbf{H}_{k}\right)\mathbf{P}_{k}^{b}\left(\mathbf{I} - \mathbf{K}_{k}\mathbf{H}_{k}\right)^{T} + \mathbf{K}_{k}\mathbf{R}_{k}\mathbf{K}_{k}^{T}$$
$$\mathbf{P}_{k+1}^{b} = \mathbf{M}_{k}\mathbf{P}_{k}^{a}\mathbf{M}_{k}^{T} + \mathbf{Q}_{k}$$

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Consider now the evolution of the same system where we perturb the observations and the forecast model with random noise drawn from the respective error covariances:

$$\widetilde{\mathbf{x}}_{a}^{k} = \widetilde{\mathbf{x}}_{b}^{k} + \mathbf{K}_{k} \left( \mathbf{y}^{k} + \mathbf{\eta}_{k} - \mathbf{H}_{k} \widetilde{\mathbf{x}}_{b}^{k} \right)$$
$$\widetilde{\mathbf{x}}_{b}^{k+1} = \mathbf{M}_{k} \widetilde{\mathbf{x}}_{a}^{k} + \boldsymbol{\zeta}_{k}$$

where  $\eta_k \sim \mathcal{N}(0, \mathbf{R})$ ,  $\zeta_k \sim \mathcal{N}(0, \mathbf{Q})$ .

If we define the differences between the perturbed and unperturbed state  $\varepsilon_a \equiv \widetilde{\mathbf{x}}_a - \mathbf{x}_a$  and  $\varepsilon_b \equiv \widetilde{\mathbf{x}}_b - \mathbf{x}_b$ , their evolution is obtained by subtracting the unperturbed state evolution equations from the perturbed ones:

$$\boldsymbol{\varepsilon}_{a}^{k} = \boldsymbol{\varepsilon}_{b}^{k} + \mathbf{K}_{k} \left( \boldsymbol{\eta}_{k} - \mathbf{H}_{k} \boldsymbol{\varepsilon}_{b}^{k} \right)$$
$$\boldsymbol{\varepsilon}_{b}^{k+1} = \mathbf{M}_{k} \boldsymbol{\varepsilon}_{a}^{k} + \boldsymbol{\zeta}_{k}$$



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• i.e., the perturbations evolve with the same update equations of the state (Kalman gain and model operator).

## What about the **errors**?

If we take the statistical expectation of the outer product of the perturbations:

$$\left\langle \boldsymbol{\varepsilon}_{k}^{a} \left( \boldsymbol{\varepsilon}_{k}^{a} \right)^{T} \right\rangle = \left( \mathbf{I} - \mathbf{K}_{k} \mathbf{H}_{k} \right) \left\langle \boldsymbol{\varepsilon}_{k}^{b} \left( \boldsymbol{\varepsilon}_{k}^{b} \right)^{T} \right\rangle \left( \mathbf{I} - \mathbf{K}_{k} \mathbf{H}_{k} \right)^{T} + \mathbf{K}_{k} \mathbf{R}_{k} \mathbf{K}_{k}^{T}$$
$$\left\langle \boldsymbol{\varepsilon}_{k+1}^{b} \left( \boldsymbol{\varepsilon}_{k+1}^{b} \right)^{T} \right\rangle = \mathbf{M}_{k} \left\langle \boldsymbol{\varepsilon}_{k}^{a} \left( \boldsymbol{\varepsilon}_{k}^{a} \right)^{T} \right\rangle \mathbf{M}_{k}^{T} + \mathbf{Q}_{k}$$

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$$\left\langle \boldsymbol{\varepsilon}_{k}^{a} \left( \boldsymbol{\varepsilon}_{k}^{a} \right)^{T} \right\rangle = \left( \mathbf{I} - \mathbf{K}_{k} \mathbf{H}_{k} \right) \left\langle \boldsymbol{\varepsilon}_{k}^{b} \left( \boldsymbol{\varepsilon}_{k}^{b} \right)^{T} \right\rangle \left( \mathbf{I} - \mathbf{K}_{k} \mathbf{H}_{k} \right)^{T} + \mathbf{K}_{k} \mathbf{R}_{k} \mathbf{K}_{k}^{T} \right\rangle$$
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• These are the same equations for the evolution of the system error covariances:

$$\mathbf{P}_{k}^{a} = \left(\mathbf{I} - \mathbf{K}_{k}\mathbf{H}_{k}\right)\mathbf{P}_{k}^{b}\left(\mathbf{I} - \mathbf{K}_{k}\mathbf{H}_{k}\right)^{T} + \mathbf{K}_{k}\mathbf{R}_{k}\mathbf{K}_{k}^{T}$$
$$\mathbf{P}_{k+1}^{b} = \mathbf{M}_{k}\mathbf{P}_{k}^{a}\mathbf{M}_{k}^{T} + \mathbf{Q}_{k}$$

provided that:

- The applied perturbations **η**<sub>k</sub>, **ζ**<sub>k</sub> have the right covariances (**R**, **Q**);
- 2. At some stage in time  $\left\langle \boldsymbol{\varepsilon}_{k}^{b} \left( \boldsymbol{\varepsilon}_{k}^{b} \right)^{T} \right\rangle = \mathbf{P}_{k}^{b}$



What does all this mean in practice?

- We can use an ensemble of perturbed assimilation cycles to simulate the errors of our reference assimilation cycle;
- The ensemble of perturbed DAs should be as similar as possible to the reference DA (i.e., same or similar **K** matrix)
- The applied perturbations η<sub>k</sub>, ζ<sub>k</sub> must have the required error covariances (**R**, **Q**);

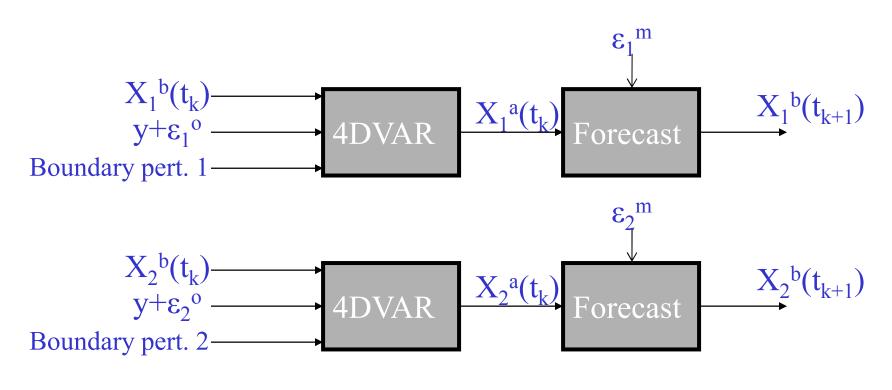
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• There is no need to explicitly perturb the background **x**<sub>b</sub>



- 25 ensemble members using 4D-Var assimilations at reduced resolution
- T399 outer loop, T95/T159 inner loops. (Reference DA: T1279 outer loop, T159/T255/T255 inner loops)
- Observations randomly perturbed according to their estimated errors
- SST perturbed with climatological error structures
- Model error represented by stochastic methods (SPPT, Leutbecher, 2009)





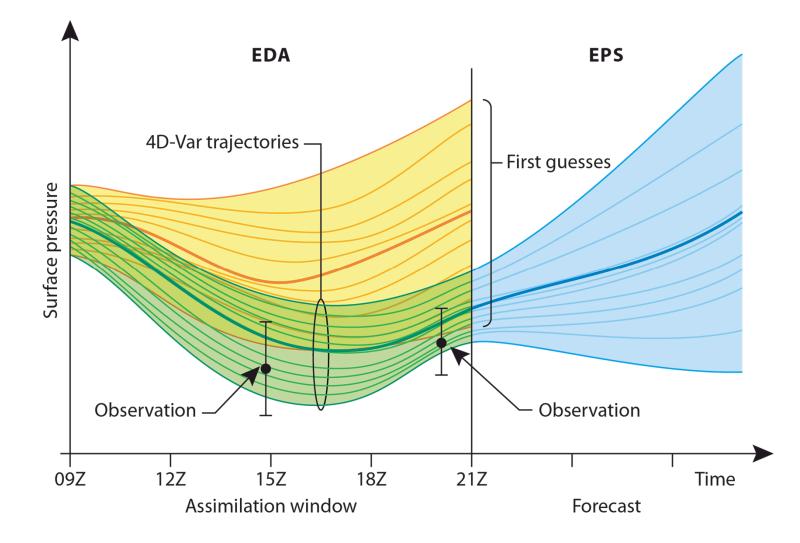
**ECMWF** 

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- The EDA simulates the error evolution of the 4DVar analysis cycle.
  As such it has two main applications:
  - 1. Provide a flow-dependent estimate of analysis errors to initialize the ensemble prediction system (EPS)
  - Provide a flow-dependent estimate of background errors for use in 4D-Var assimilation







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## • Note that the EDA impacts the EPS in two distinct ways:

- 1. Directly, by providing a flow-dependent estimate of the initial perturbations of the ensemble prediction system (EPS)
- Indirectly, through the provision of a flow-dependent estimate of background errors for use in the 4D-Var deterministic (HRES) assimilation cycle: this improves the deterministic analysis around which the EPS initial perturbations are re-centered



- 1. Directly, by providing a flow-dependent estimate of the initial perturbations of the ensemble prediction system (EPS)
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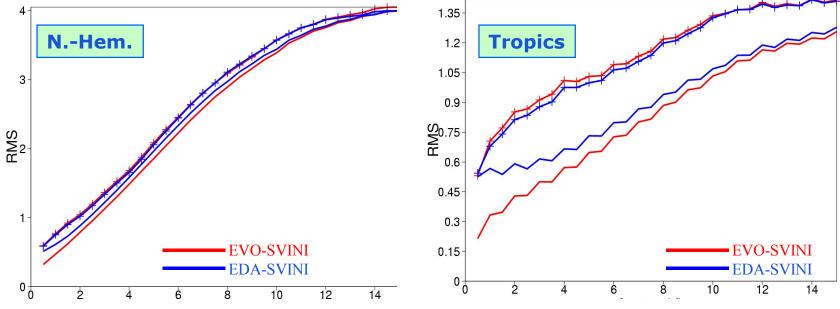
## The second effect is arguably the more important:

"... the main reason for the better performance of the EC-EPS in terms of the RMS, PAC, ROC and Brier Skill Score measures is the superiority of the ECMWF data assimilation (and perhaps numerical forecast modelling) system, and not necessarily the strategy used to simulate initial value and model related uncertainties in the EC-EPS." Buizza et.al., 2005



Improving Ensemble Prediction System by including EDA perturbations for initial uncertainty (implemented June 2010)

The Ensemble Prediction System (EPS) benefits from using EDA based perturbations. Replacing evolved singular vector perturbations by EDA based perturbations improve EPS spread, especially in the tropics. The Ensemble Mean has slightly lower error when EDA is used.



Ensemble spread and Ensemble mean RMSE for 850hPa T

- The EDA simulates the error evolution of the 4DVar analysis cycle. As such it has two main applications:
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# Hybrids: EDA

In the ECMWF 4D-Var, the **B** matrix is defined implicitly in terms of a transformation from the background departure  $(x-x_b)$  to a control variable  $\chi$ :

$$(\mathbf{x}-\mathbf{x}_{\mathbf{b}}) = \mathbf{L}\boldsymbol{\chi}$$

So that the implied  $B=LL^T$ .

In the current wavelet formulation (Fisher, 2003), the variable transform can be written as:

$$(\mathbf{x} - \mathbf{x}_b) = \mathbf{T}^{-1} \mathbf{\Sigma}_b^{1/2} \sum_j \psi_j \otimes [\mathbf{C}_j^{1/2}(\lambda, \phi) \chi_j]$$

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T is the balance operator, i.e. the operator that links the control variables to the model variables

 $\Sigma_{\rm b}$  is the gridpoint variance of background errors  $C_{\rm j}(\lambda, \varphi)$  is the vertical covariance matrix for wavelet index *j*  $\psi_i$  are the set of radial basis function that define the wavelet transform.

**ECMWF** 

## Hybrids: EDA

$$(\mathbf{x} - \mathbf{x}_b) = \mathbf{T}^{-1} \boldsymbol{\Sigma}_b^{1/2} \sum_{i} \boldsymbol{\psi}_i \otimes [\mathbf{C}_j^{1/2}(\lambda, \phi) \boldsymbol{\chi}_j]$$

 $C_{j}(\lambda, \varphi)$  are full vertical covariance matrices, function of  $(\lambda, \varphi)$ . They determine both the horizontal and vertical background error *correlation structures*;

In standard 4D-Var T and  $C_j$  are computed off-line using a climatology of EDA perturbations.

 $\Sigma_{\rm b}$  is computed by random sampling of the static **B** matrix (randomization procedure, Fisher and Courtier, 1995)

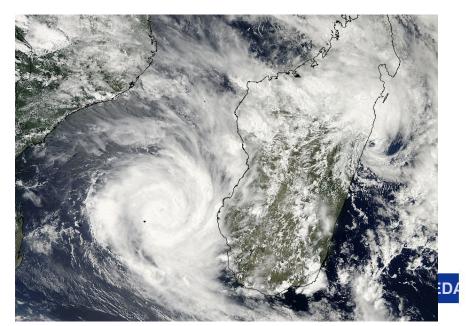
## How do we make this error covariance model flow-dependent?

We look for flow-dependent EDA estimates of  $\Sigma_{b}$  and  $C_{i}(\lambda, \varphi)$ 

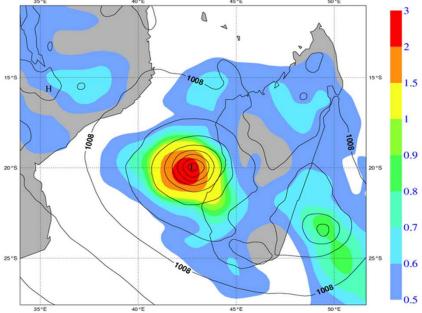


- We want to use EDA perturbations to simulate 4DVar flowdependent background error covariance evolution
- We start with the EDA flow-dependent estimates of background error variances (diagonal of the B matrix,  $\Sigma_b$ )

### Hurricane Fanele, 20 January 2009

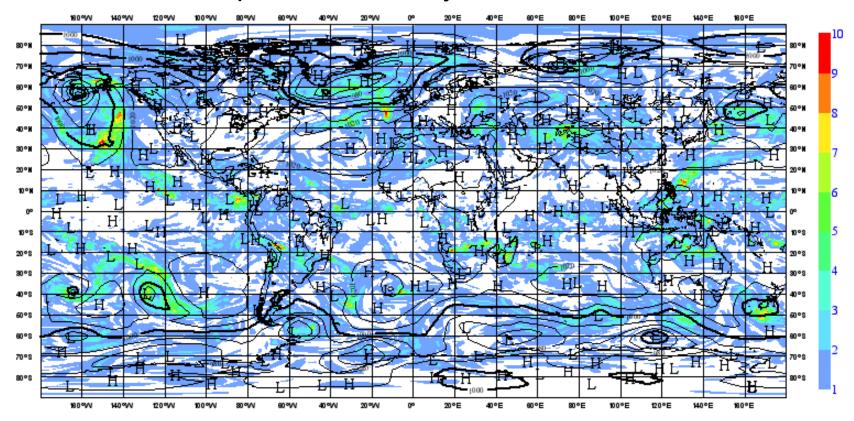


EDA based background error variance for surface pressure



## What do raw ensemble variances look like?

## Spread of Vorticity t+9h 500hPa



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- Noise level is due to sampling errors: 25 member ensemble
- EDA is a stochastic system: error variance of variance estimator ~ 1/N<sub>ens</sub>
- We need a system to effectively filter out noise from first guess ensemble forecast variances: Reduce the random component of the estimation error



Mallat et al.: 1998, Annals of Statistics, 26,1-47

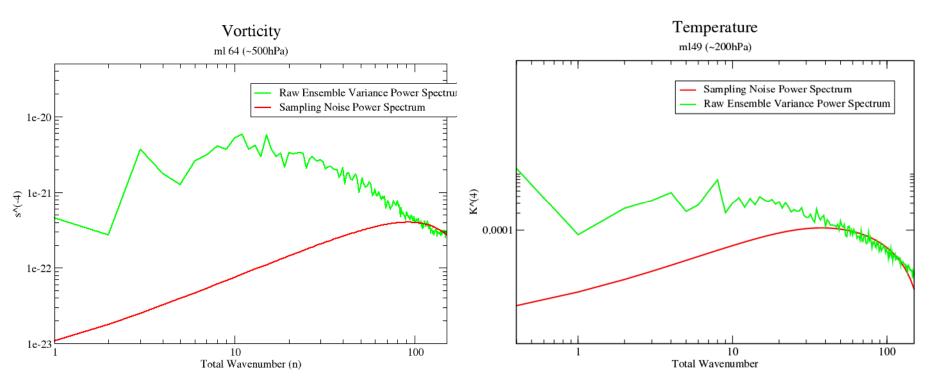
Define G<sup>e</sup>(i) as the random component of the sampling error in the estimated ensemble variance at gridpoint *i*:

$$G^{e}(i) \equiv \widetilde{B}_{ii} - E\left[\widetilde{B}_{ii}\right]$$

Then the covariance of the sampling noise can be shown to be a simple function of the expectation of the ensemble-based covariance matrix:

$$E\left[G^{e}(i)G^{e}(j)\right] = \frac{2}{N-1} \left(E\left[\widetilde{B}_{ij}\right]\right)^{2} \qquad (1)$$

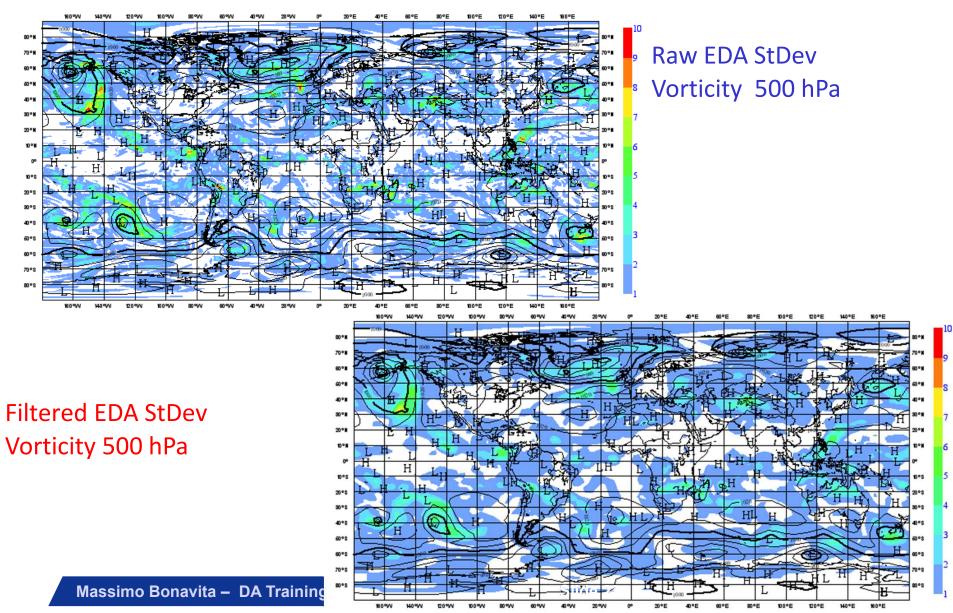


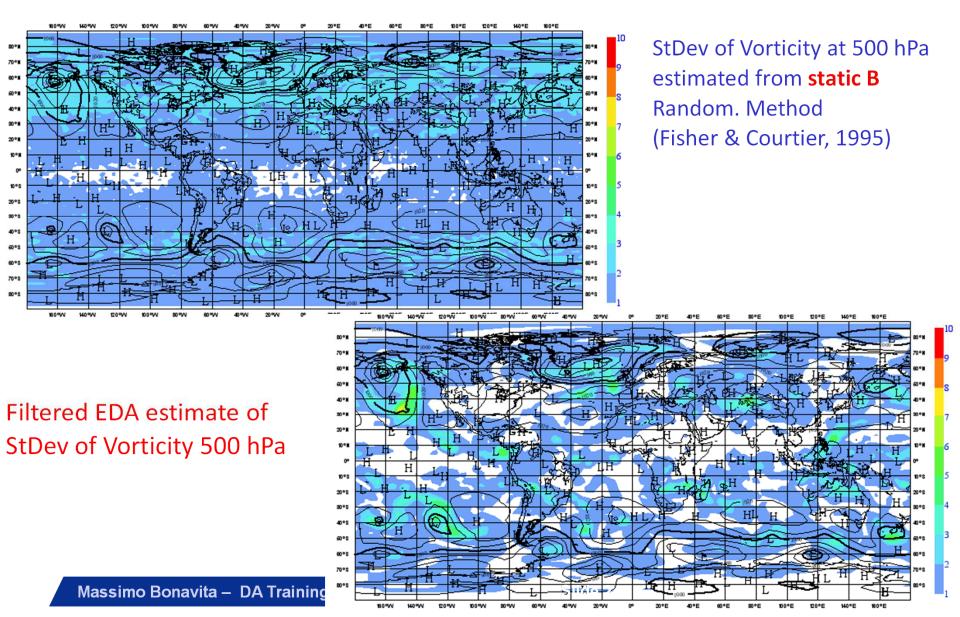


- We can use a spectral filter to disentangle noise from signal
- Truncation wavenumber is determined by maximizing signalto-noise ratio of filtered variances (Raynaud *et al.,* 2009; Bonavita *et al.,* 2011)

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**CECMWF** 





# Is there also a systematic error in our EDA sampled variances?

A statistically consistent ensemble satisfies:

(1-1/N<sub>ens</sub>)<sup>-1</sup> (1+1/N<sub>ens</sub>)<sup>-1</sup> Ens\_Mean\_Square\_Error

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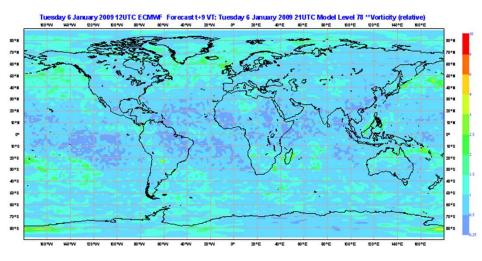


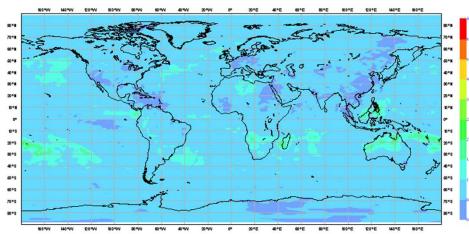
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## Vorticity ml 78 (~850hPa)

#### **Ensemble Error**

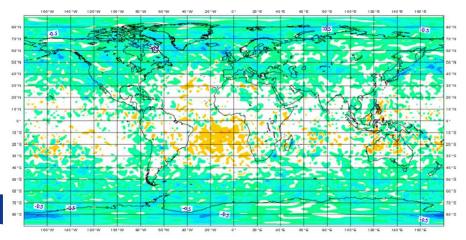
#### **Ensemble Spread**



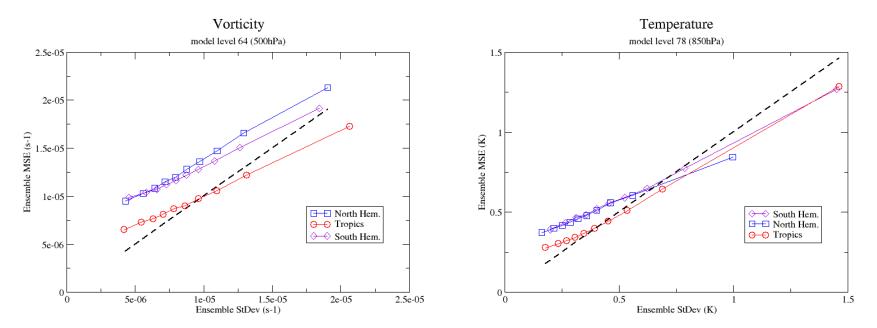


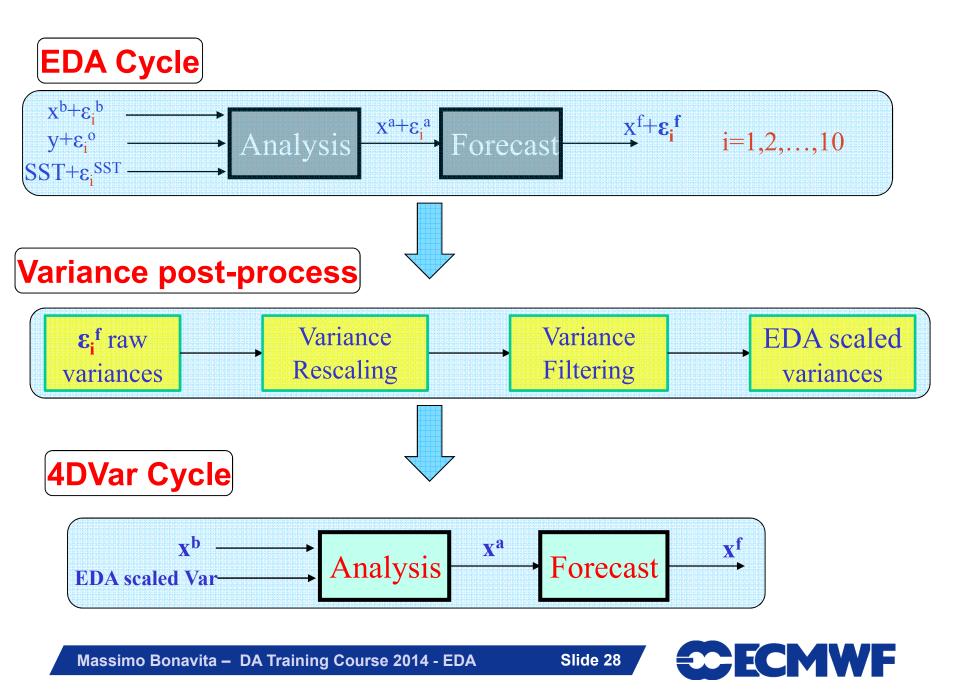
#### **Spread - Error**

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- To get statistically consistent EDA variances we need to perform an online calibration (Ensemble Variance Calibration; Kolczynsky et al., 2009, 2011; Bonavita et al., 2011)
- Calibration factors are also state-dependent, i.e. depend on the size of the expected error
- Need to perform calibration of variances reflects underlying problem in Q and R models, system non-linearities, ensemble size



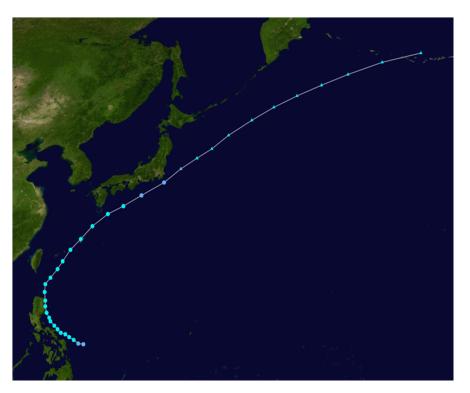


- How does the flow-dependent structure of the EDA variances affect the 4DVar analysis?
- Two mechanisms:
  - 1. Inside 4DVar EDA variances change the shape and size of analysis increments
  - 2. Before 4DVar they affect the observation quality control decisions



- 1. Inside 4DVar EDA variances change the shape and size of analysis increments
- Tropical Cyclone Aere, Philippines 8-9 May 2011.



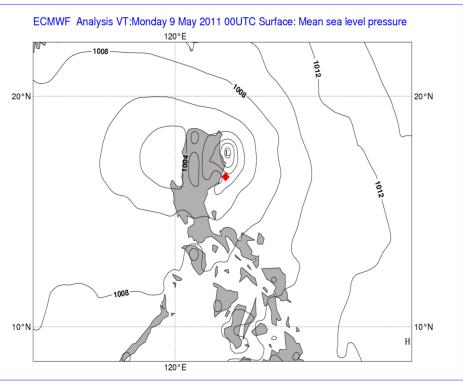


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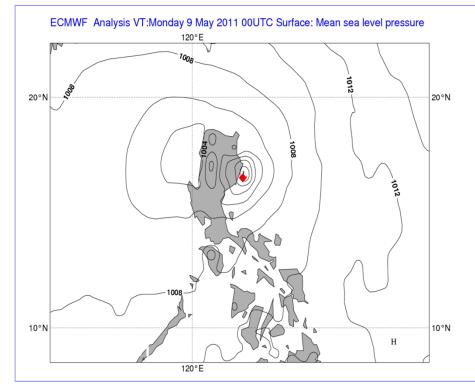


- 1. Inside 4DVar EDA variances change the shape and size of analysis increments
- Significant operational analysis error, corrected by 4DVar with EDA variances

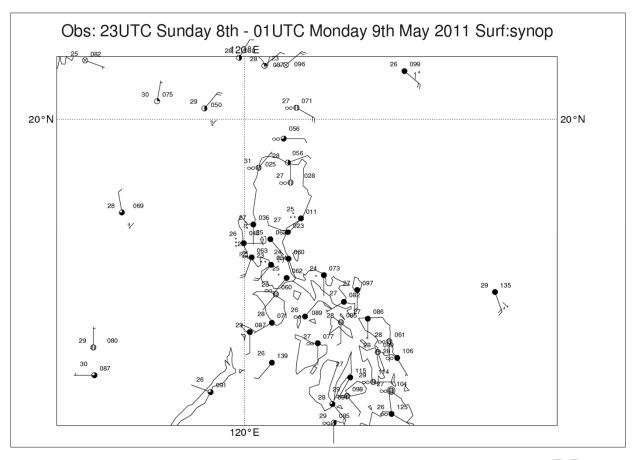
### 4DVar with Static errors



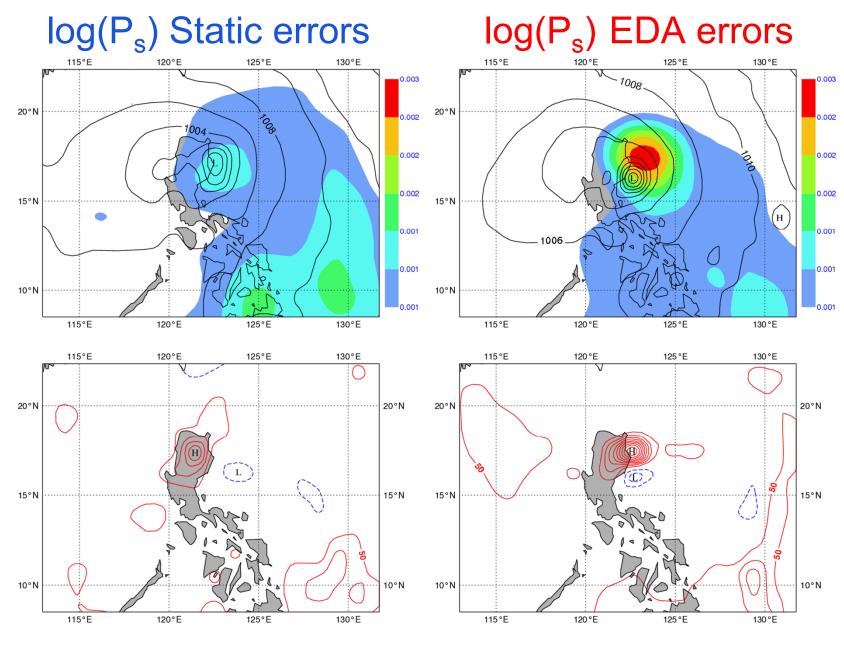




 Inside 4DVar EDA variances change the shape and size of analysis increments



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Static mslp ana incr.

EDA mslp ana incr.

- Flow-dependent EDA errors have been used operationally since May 2012 (CY37R2)
- The effect of using flow-dependent EDA estimated errors is large on average skill scores



## **Geopotential RMSE reduction**

RMS forecast errors in Z(ffg8-fezi), 11-Jan-2010 to 30-Mar-2010, from 72 to 79 samples. RMS forecast errors in Z(ffge-0051), 2-Aug-2010 to 30-Oct-2010, from 83 to 90 samples. Point confidence 99.5% to give multiple-comparison adjusted confidence 90%. Verified against own-analysis.

Point confidence 99.5% to give multiple-comparison adjusted confidence 90%. Verified against own-analysis.

 0.10

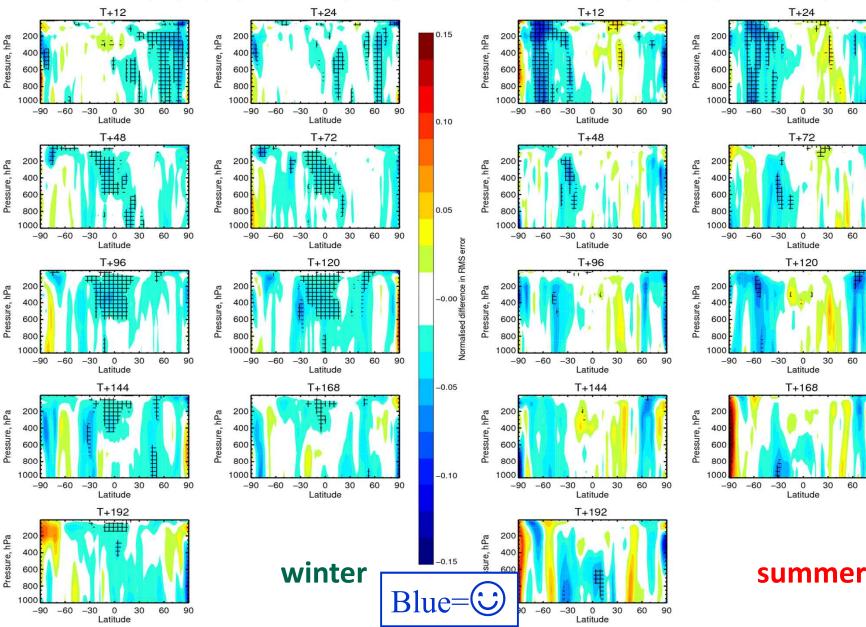
0.05

Normalised difference in RMS error

0.00

-0.05

-0.10



$$(\mathbf{x} - \mathbf{x}_b) = \mathbf{T}^{-1} \boldsymbol{\Sigma}_b^{1/2} \sum_j \boldsymbol{\psi}_j \otimes \left[ \mathbf{C}_j^{1/2} (\lambda, \phi) \boldsymbol{\chi}_j \right]$$

T is the balance operator

 $\Sigma_{\rm b}$  is the gridpoint variance of background errors

 $C_j(\lambda, \varphi)$  is the vertical covariance matrix for wavelet index *j*  $\psi_j$  are the set of radial basis function that define the wavelet transform

 $C_{j}(\lambda, \varphi)$  are fields of full vertical covariance matrices, defined for each wavelet band. They determine both the horizontal and vertical background error correlation structures.

In order to get flow-dependent error covariances we need flow-dependent estimates of  $C_i(\lambda, \varphi)$ .

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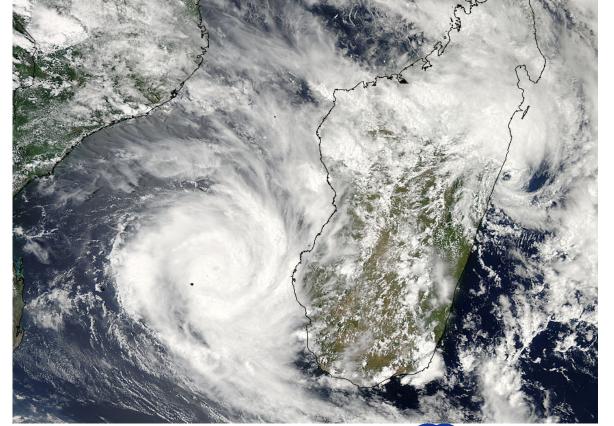
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**Diagnosing the Background Error Correlation Length-Scales** 

 $L = \sqrt{-\frac{\rho(r)}{\partial^2 \rho(r)/\partial r^2}}$ 

#### Hurricane Fanele 20 January 2009



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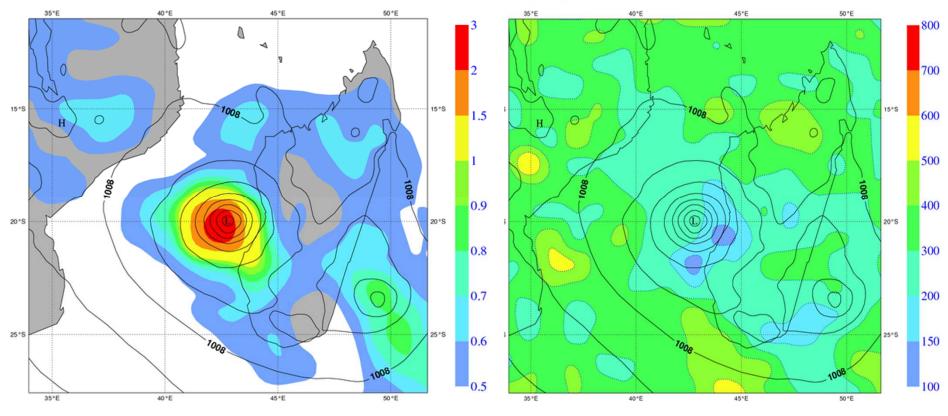


#### 20 member EDA

#### Surf. Press. Background Err. St.Dev. Surf. Press. BG Err. Correlation L. Scale

Tuesday 20 January 2009 00UTC ECMWF. Forecast t+9 VT: Tuesday 20 January 2009 09UTC Surface: Mean sea level pressure

#### Tuesday 20 January 2009 00UTC ECMWF Forecast t+9 VT: Tuesday 20 January 2009 09UTC Surface: Mean sea level pressure



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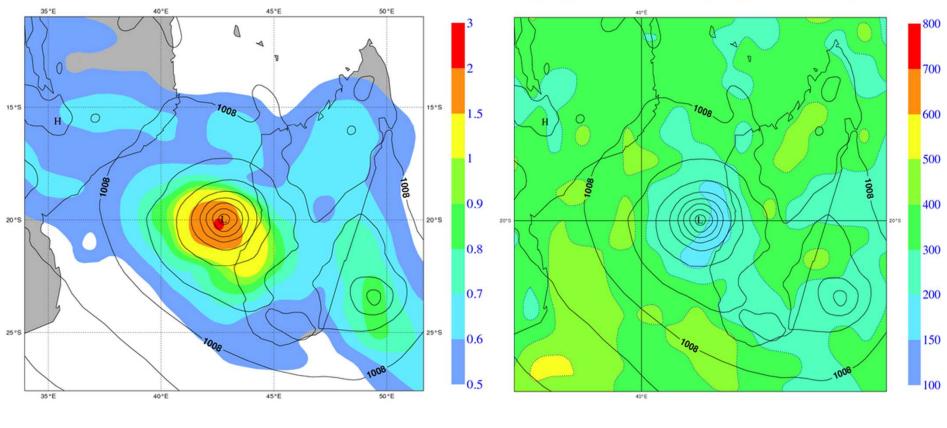


#### **50 member EDA**

Surf. Press. Background Err. St.Dev. Surf. Press. BG Err. Correlation L. Scale

Tuesday 20 January 2009 00UTC ECMWF Forecast I+9 VT: Tuesday 20 January 2009 09UTC Surface: Mean sea level pressure

Tuesday 20 January 2009 00UTC ECMWF. Forecast t+9 VT: Tuesday 20 January 2009 09UTC Surface: Mean sea level pressure



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- a) From November 2013 (CY40R1) background error covariances (wavelet JB) are computed online, i.e. they are updated at every assimilation time (00, 12 UTC)
- b) EDA perturbations from the past 12 days are used, with an exponential decay factor (i.e., reduce noise at the cost of losing some flow-dependent detail)
- c) Continuously updated JB is used in High Resolution 4D-Var



#### Why is a flow-dependent JB better? Correlation length scale of Vorticity errors, ~500 hPa 120°E 140°E 160°E

Online wavelet JB, valid 20120110 00Z

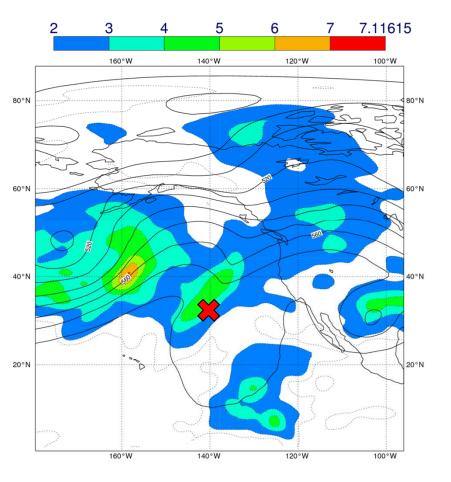
170 80°N **AF** 165 60°N Œ 160 40°N 155 20°N 150 145 0° 140 20°S H 135 40°S 130 60°S 125 520 80°S H 120 160°W 140°W 120°W 100°W 80°W 60°W 60°E 80°E 100°E 120°E 140°E 160°E 40°W 20°W 0° 20°E 40°E 80°E 120°E 140°E 160°E 160°W 140°W 120°W 100°W 80°W 60°W 40°W 20°W 0° 20°E 40°E 60°E 100°E 170 80°N 165 60°N T 160 40°N 155 20°N 150 145 0° 140 20°S F 135 H 40°S 130 0 Н Massimo Bonavita – DA Train 620 125 520 H 80°S 120

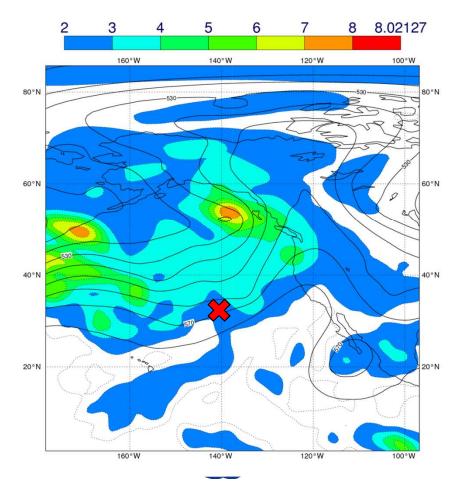
Static wavelet JB

#### Why is a flow-dependent JB better?

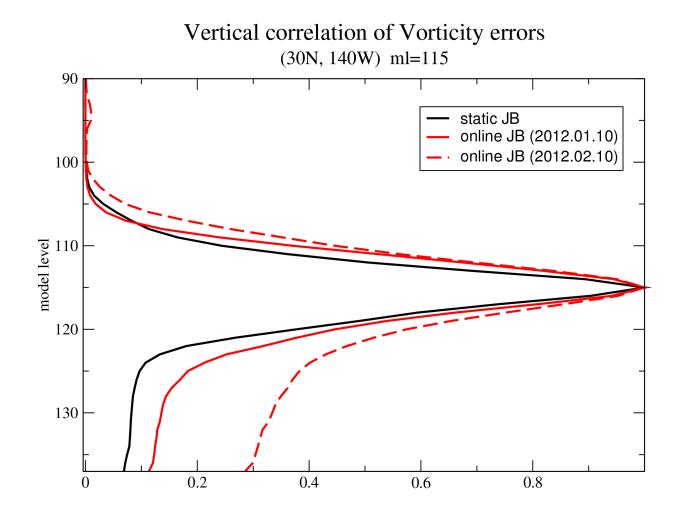
#### 2012-01-01 00Z

#### 2012-02-01 00Z





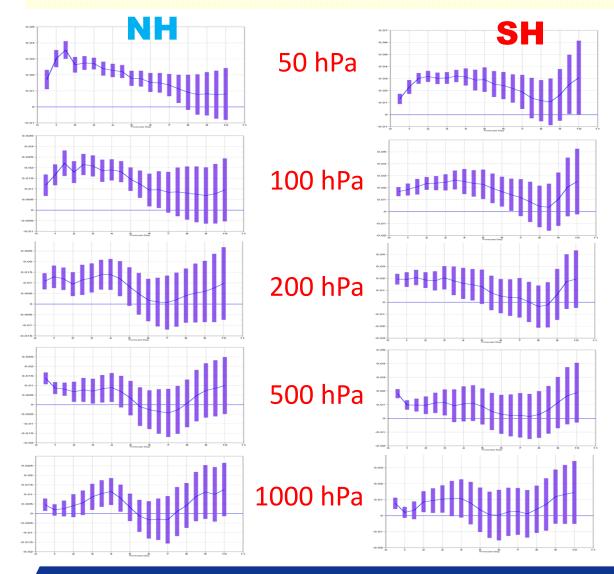
### Why is a flow-dependent JB better?





# Impact of online JB

**Reduction in Geopotential RMSE - 95% confidence** 



Period: Feb - June 2012

T511L91, 3 Outer Loops (T159/T255/T255)

Verified against operational analysis



- The EDA is the best available tool at ECMWF to estimate analysis and background errors of the ECMWF analysis
- The EDA derived analysis errors contribute to the estimation of initial perturbations of the Ensemble Prediction System (EPS)
- The EDA background perturbations provide flow-dependent estimates of background errors variances and covariances for use in the High-Resolution 4DVar analysis



- The use of EDA background perturbations in a variational analysis is currently done in two ways:
  - a) adding an ensemble, flow-dependent component to the static B used in 4D-Var (UK MetO, NCEP; alpha control variable method: Lorenc, 2003)

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 b) using EDA perturbations to get an on-line, flowdependent estimate of parameterised B (ECMWF, Meteo France; EDA approach)



- Flow-dependent background errors from EDA variances have been used in ECMWF 4D-Var since April 2011 (Cycle 37R2)
- They benefit the analysis and forecast skill by:
  - a) changing the weight given to observations near dynamically active zones;
  - b) introducing a level of flow-dependency in the analysis increments
  - c) Allowing a state and flow-dependent QC of observations
- Increase in ensemble size will benefit the system allowing less aggressive filtering of raw variances (Bonavita *et al.,* 2011)

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- Flow-dependent background error covariances (B) estimated from EDA perturbations has been implemented recently (CY40R1, November 2013)
- They also have an important effect on the analysis by providing flow-dependent, non homogeneous, correlation structures
- The introduction of flow-dependent background errors and covariances estimated from the EDA has been one of the largest source of improvement in recent years in the analysis and forecast skill of the IFS



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Massimo Bonavita – DA Training Course 2014 - EDA

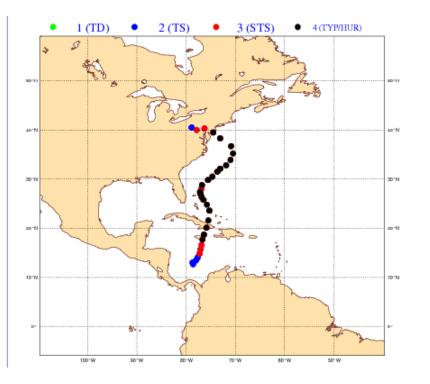
**CECMWF** 

#### **Additional Slides**

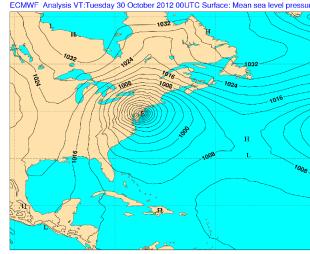




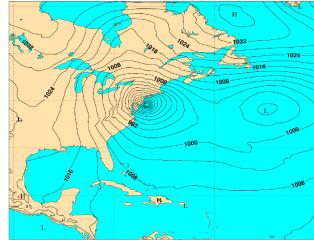
- 2. Before 4DVar they affect the observation quality control decisions
- Super Storm Sandy



Massimo Bonavita - DA Training Course 2014 - El



Monday 22 October 2012 00UTC ECMWF Forecast I+204 VT: Tuesday 30 October 2012 12UTC Surface: Mean sea level pressure



Mslp Ana 30/10/2012 00UTC

Mslp t+204h Fcst valid at 30/10/2012 00UTC

What happens if we withhold polar-orbiters observations (i.e., approx. 90% of obs. counts)?

The forecast performance is obviously degraded, and only 5 days before landfall the system recovers the correct track

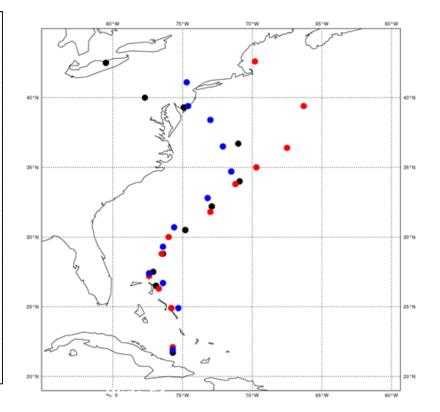
Sandy's forecast tracks 25 Oct 2013 00UTC

**Operational forecast** 

Forecast from HRES assimilation cycle without polar orbiters and errors from operational EDA

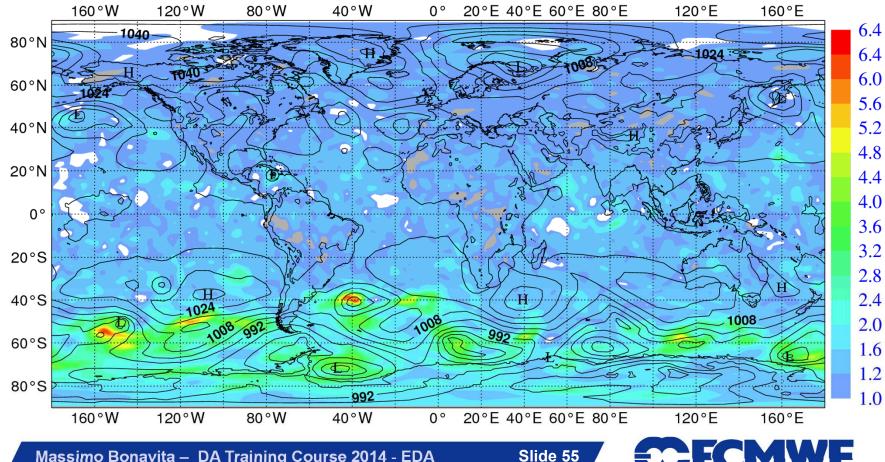
Forecast from HRES assimilation cycle and EDA both without polar orbiters data





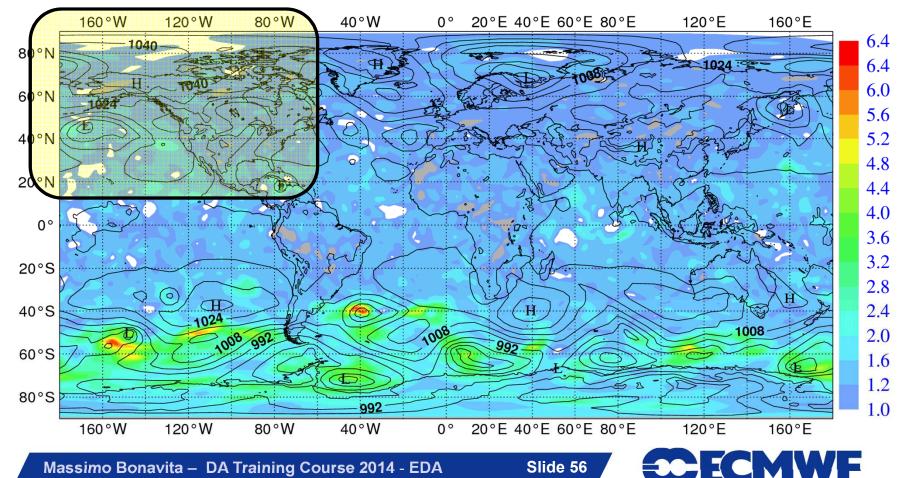
EDA without polar orbiters' data has larger spread than operational EDA

EDA spread for u-wind component at 850 hPa: No-Polar/Oper ratio



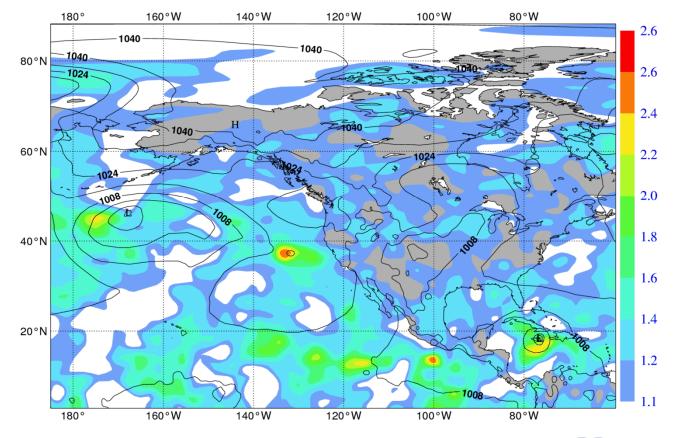
EDA without polar orbiters' data has larger spread than operational EDA

EDA spread for u-wind component at 850 hPa: No-Polar/Oper ratio



#### EDA without polar orbiters' data has larger spread than operational EDA

EDA spread for u-wind component at 850 hPa: No-Polar/Oper ratio

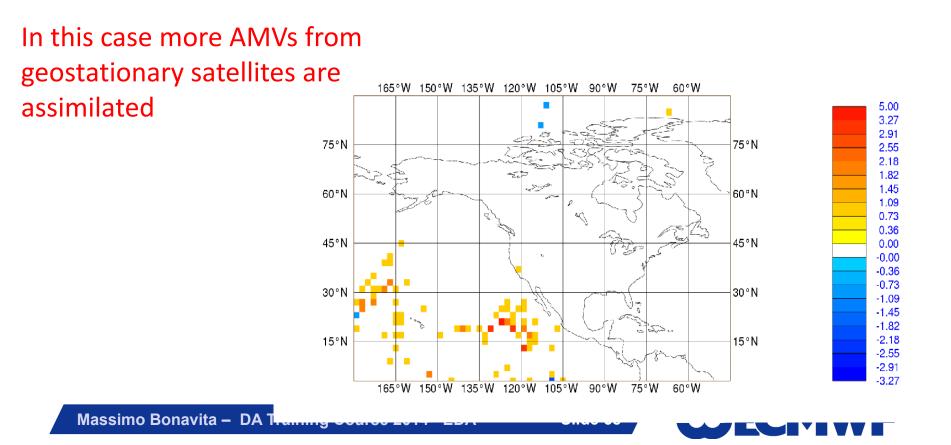


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EDA without polar orbiters' data has larger spread than operational EDA

This has two effects: a) Observations are more closely fit and b) More observations pass first guess quality control:  $(y-\mathcal{H}(x))^2 \le \alpha(\sigma_b^2 + \sigma_o^2)$ 



#### In this case more AMVs from geostationary satellites are assimilated

