Efficient Data Assimilation for Spatiotemporal Chaos: a Local Ensemble Transform Kalman Filter

Brian R. Hunt

Institute for Physical Science and Technology and Department of Mathematics University of Maryland, College Park MD 20742

November 28, 2005

Abstract

Data assimilation is an iterative approach to the problem of estimating the state of a dynamical system using both current and past observations of the system together with a model for the system's time evolution. Rather than solving the problem from scratch each time new observations become available, one uses the model to "forecast" the current state, using a prior state estimate (which incorporates information from past data) as the initial condition, then uses current data to correct the prior forecast to a current state estimate. This Bayesian approach is most effective when the uncertainty in both the observations and in the state estimate, as it evolves over time, are accurately quantified. In this article, I describe a practical method for data assimilation in large, spatiotemporally chaotic systems. The method is a type of "Ensemble Kalman Filter", in which the state estimate and its approximate uncertainty are represented at any given time by an ensemble of system states. I discuss both the mathematical basis of this approach and its implementation; my primary emphasis is on ease of use and computational speed rather than improving accuracy over previously published approaches to ensemble Kalman filtering.

1 Introduction

Forecasting a physical system generally requires both a model for the time evolution of the system and an estimate of the current state of the system. In some applications, the state of the system can be measured accurately at any given time, but in other applications, such as weather forecasting, direct measurement of the system state is not feasible. Instead, the state must be inferred from available data. While a reasonable state estimate based on current data may be possible, in general one can obtain a better estimate by using both current and past data. "Data assimilation" provides such an estimate on an ongoing basis, iteratively alternating between a forecast step and a state estimation step; the latter step is often called the "analysis". The analysis step combines information from current data and from a prior short-term forecast (which is based on past data), producing a current state estimate. This estimate is used to initialize the next short-term forecast, which is subsequently used in the next analysis, and so on. The data assimilation procedure is itself a dynamical system driven by the physical system, and the practical problem is to achieve good "synchronization" [33] between the two systems.

Data assimilation is widely used to study and forecast geophysical systems [10, 21]. The analysis step is generally a statistical procedure (specifically, a Bayesian maximum likelihood estimate) involving a prior (or "background") estimate of the current state based on past data, and current data (or "observations") that are used to improve the state estimate. This procedure requires quantification of the uncertainty in both the background state and the observations. While quantifying the observation uncertainty can be a nontrivial problem, in this article I will consider that problem to be solved, and instead concentrate on the problem of quantifying the background uncertainty.

There are two main factors that create background uncertainty. One is the uncertainty in the initial conditions from the previous analysis, which produced the background state via a short-term forecast. The other is "model error", the unknown discrepancy between the model dynamics and actual system dynamics. Quantifying the uncertainty due to model error is a challenging problem, and while this problem generally can't be ignored in practice, I will discuss only crude ways of accounting for it in this article. For the time being, let us consider an idealized "perfect model" scenario, in which there is no model error.

The main purpose of this article is to describe a practical framework for data assimilation that is both relatively easy to implement and computationally efficient, even for large, spatiotemporally chaotic systems. (By "spatiotemporally chaotic" I mean a spatially extended system that exhibits temporally chaotic behavior with weak long-range spatial correlations.) The emphasis here is on methodology that scales well to high-dimensional systems and large numbers of observations, rather than on what would be optimal given unlimited computational resources. Ideally one would keep track of a probability distribution of system states, propagating the distribution using the Fokker-Planck-Kolmogorov equation during the forecast step. While this approach provides a theoretical basis for the methods used in practice [18], it would be computationally expensive even for a low-dimensional system and not at all feasible for a high-dimensional system. Instead one can use a Monte Carlo approach, using a large ensemble of system states to approximate the distribution (see [5] for an overview), or a parametric approach like the Kalman Filter [19, 20], which assumes Gaussian distributions and tracks their mean and covariance. (The latter approach was derived originally for linear problems, but serves as a reasonable approximation for nonlinear problems when the uncertainties remain sufficiently small.)

The methodology of this article is based on the Ensemble Kalman Filter [6, 8], which has elements of both approaches: it uses the Gaussian approximation, and follows the time evolution of the mean and covariance by propagating an ensemble of states. The ensemble can be reasonably small relative to other Monte Carlo methods because it is used only to parametrize the distribution, not to sample it thoroughly. The ensemble should be large enough to approximately span the space of possible system states at a given time, because the analysis essentially determines which linear combination of the ensemble members form the best estimate of the current state, given the current observations.

Many variations on the Ensemble Kalman Filter have been published in the geophysical literature, and this article draws ideas from a number of them [1, 2, 3, 12, 14, 15, 23, 29, 30, 38, 40]. These articles in turn draw ideas both from earlier work on geophysical data assimilation and from the engineering and mathematics literature on nonlinear filtering. For the most part, I will limit my citations to ensemble-based articles rather than attempt to trace all ideas to their original sources. I call the method described here a Local Ensemble Transform Kalman Filter (LETKF), because it is most closely related to the Ensemble Transform Kalman Filter [3] and the Local Ensemble Kalman Filter [29, 30].

In Section 2, I start by posing a general problem about which trajectory of a dynamical system "best fits" a time series of data; this problem is solved exactly for linear problems by the Kalman Filter and approximately for nonlinear problems by ensemble Kalman filters. Next I derive the Kalman Filter equations as a guide for what follows. Then I discuss ensemble Kalman filters in general and the issue of "localization", which is important for applications to spatiotemporally chaotic systems. Finally, I develop the basic LETKF equations, which provide a framework for data assimilation that allow a system-dependent localization strategy to be developed and tuned. I discuss also several options for "covariance inflation" to compensate for the effects of model error and the deficiencies of small sample size and linear approximation that are inherent to ensemble Kalman filters.

In Section 3, I give step-by-step instructions for efficient implementation of the approach developed in Section 2, and discuss options for further improving computational speed in certain cases. Then in Section 4, I present a generalization that allows observations gathered at different times to be assimilated simultaneously in a natural way. Finally, in Section 5 I briefly discuss preliminary results and work in progress with the LETKF approach. The notation in this article is based largely on that proposed in [17], with some elements from [30].

2 Mathematical Formulation

Consider a system governed by the ordinary differential equation

$$\frac{d\mathbf{x}}{dt} = F(t, \mathbf{x}),\tag{1}$$

where x is an *m*-dimensional vector representing the state of the system at a given time. Suppose we are given a set of (noisy) observations of the system made at various times, and we want to determine which trajectory $\{x(t)\}$ of (1) "best" fits the observations. For any given t, this trajectory gives an estimate of the system state at time t.

To formulate this problem mathematically, we need to define "best fit" in this context. Let us assume that the observations are the result of measuring quantities that depend on the system state in a known way, with Gaussian measurement errors. In other words, an observation at time t_j is a triple $(\mathbf{y}_j^o, H_j, \mathbf{R}_j)$, where \mathbf{y}_j^o is a vector of observed values, and H_j and \mathbf{R}_j describe the relationship between \mathbf{y}_j^o and $\mathbf{x}(t_j)$:

$$\mathbf{y}_j^o = H_j(\mathbf{x}(t_j)) + \varepsilon_j,$$

where ε_j is a Gaussian random variable with mean 0 and covariance matrix \mathbf{R}_j . Notice that I am assuming a perfect model here: the observations are based on a trajectory of (1), and our problem is simply to infer which trajectory produced the observations. In a real application, the observations come from a trajectory of the physical system for which (1) is only a model. So a more realistic (but more complicated) problem would be to determine a pseudo-trajectory of (1), or a trajectory of an associated stochastic differential equation, that best fits the observations. Formulating this problem mathematically then requires some assumption about the size and nature of the model error. I will use the perfect model problem as motivation and defer the consideration of model error until later.

Given our assumptions about the observations, we can formulate a maximum likelihood estimate for the trajectory of (1) that best fits the observations at times $t_1 < t_2 < \cdots < t_n$. The likelihood of a trajectory $\mathbf{x}(t)$ is proportional to

$$\prod_{j=1}^{n} \exp\left(-[\mathbf{y}_{j}^{o} - H_{j}(\mathbf{x}(t_{j}))]^{T} \mathbf{R}_{j}^{-1}[\mathbf{y}_{j}^{o} - H_{j}(\mathbf{x}(t_{j}))]\right)$$

The most likely trajectory is the one that maximizes this expression, or equivalently minimizes the "cost function"

$$J^{o}(\{\mathbf{x}(t)\}) = \sum_{j=1}^{n} [\mathbf{y}_{j}^{o} - H_{j}(\mathbf{x}(t_{j}))]^{T} \mathbf{R}_{j}^{-1} [\mathbf{y}_{j}^{o} - H_{j}(\mathbf{x}(t_{j}))].$$
(2)

Thus, the "most likely" trajectory is also the one that best fits the observations in a least square sense.

Notice that (2) expresses the cost J^o as a function of the trajectory $\{\mathbf{x}(t)\}$. To minimize the cost, it is more convenient to write it as a function of the system state at a particular time t. Let $M_{t,t'}$ be the map that propagates a solution of (1) from time t to time t'. Then

$$J_t^o(\mathbf{x}) = \sum_{j=1}^n [\mathbf{y}_j^o - H_j(M_{t,t_j}(\mathbf{x}))]^T \mathbf{R}_j^{-1} [\mathbf{y}_j^o - H_j(M_{t,t_j}(\mathbf{x}))]$$
(3)

expresses the cost in terms of the system state x at time t. Thus to estimate the state at time t, we attempt to minimize J_t^o .

For a nonlinear model, there is no guarantee that a unique minimum exists. And even if it does, evaluating J_t^o is apt to be computationally expensive, and minimizing it may be impractical. But if both the model and the observation operators H_j are linear, the minimization is quite tractable, because J_t^o is then quadratic. Furthermore, one can compute the minimum by an iterative method, namely the Kalman Filter [19, 20], which I will now describe. This method forms the basis for the approach we will use in the nonlinear scenario.

2.1 Linear Scenario: the Kalman Filter

In the linear scenario, we can write $M_{t,t'}(\mathbf{x}) = \mathbf{M}_{t,t'}\mathbf{x}$ and $H_j(\mathbf{x}) = \mathbf{H}_j\mathbf{x}$ where $\mathbf{M}_{t,t'}$ and \mathbf{H}_j are matrices. Using the terminology from the introduction, we now describe how to perform a forecast step from time t_{n-1} to time t_n followed by an analysis step at time t_n , in such a way that if we start with the most likely system state, in the sense described above, given the observations up to time t_{n-1} , we end up with the most likely state given the observations up to time t_n . The forecast step propagates the solution from time t_{n-1} to time t_n , and the analysis step combines the information provided by the observations at that time with the propagated information from the prior observations. This iterative approach requires that we keep track not only the most likely state, but also its uncertainty, in the sense described below. (Of course, the fact that the Kalman Filter computes the uncertainty in its state estimate may be viewed as a virtue.)

Suppose the analysis at time t_{n-1} has produced a state estimate $\bar{\mathbf{x}}_{n-1}^a$ and an associated covariance matrix \mathbf{P}_{n-1}^a . In probabilistic terms, $\bar{\mathbf{x}}_{n-1}^a$ and \mathbf{P}_{n-1}^a represent the mean and covariance of a Gaussian probability distribution that represents the relative likelihood of the possible system states given the observations from time t_1 to t_{n-1} . Algebraically, what we assume is that for some constant c,

$$\sum_{j=1}^{n-1} [\mathbf{y}_{j}^{o} - \mathbf{H}_{j} \mathbf{M}_{t_{n-1}, t_{j}} \mathbf{x}]^{T} \mathbf{R}_{j}^{-1} [\mathbf{y}_{j}^{o} - \mathbf{H}_{j} \mathbf{M}_{t_{n-1}, t_{j}} \mathbf{x}] = [\mathbf{x} - \bar{\mathbf{x}}_{n-1}^{a}]^{T} (\mathbf{P}_{n-1}^{a})^{-1} [\mathbf{x} - \bar{\mathbf{x}}_{n-1}^{a}] + c.$$
(4)

In other words, the analysis at time t_{n-1} has "completed the square" to express the part of the quadratic cost function $J_{t_{n-1}}^o$ that depends on the observations up to that time as a single quadratic

form plus a constant. The Kalman Filter determines $\bar{\mathbf{x}}_n^a$ and \mathbf{P}_n^a such that an analogous equation holds at time t_n .

First we propagate the analysis state estimate $\bar{\mathbf{x}}_{n-1}^a$ and its covariance matrix \mathbf{P}_{n-1}^a using the forecast model to produce a background state estimate $\bar{\mathbf{x}}_n^b$ and covariance \mathbf{P}_n^b for the next analysis:

$$\bar{\mathbf{x}}_n^b = \mathbf{M}_{t_{n-1}, t_n} \bar{\mathbf{x}}_{n-1}^a, \tag{5}$$

$$\mathbf{P}_{n}^{b} = \mathbf{M}_{t_{n-1},t_{n}} \mathbf{P}_{n-1}^{a} \mathbf{M}_{t_{n-1},t_{n}}^{T}.$$
(6)

Under a linear model, a Gaussian distribution of states at one time propagates to a Gaussian distribution at any other time, and the equations above describe how the model propagates the mean and covariance of such a distribution. (Often a constant matrix is added to the right side of (6) to represent additional uncertainty due to model error.)

Next, we want to rewrite the cost function $J_{t_n}^o$ given by (3) in terms of the background state estimate and the observations at time t_n . (This step is often formulated as applying Bayes' Rule to the corresponding probability density functions.) In (4), x represents a hypothetical system state at time t_{n-1} . In our expression for $J_{t_n}^o$, we want x to represent instead a hypothetical system state at time t_n , so we first replace x by $\mathbf{M}_{t_n,t_{n-1}}\mathbf{x} = \mathbf{M}_{t_{n-1},t_n}^{-1}\mathbf{x}$ in (4). Then using (5) and (6) yields

$$\sum_{j=1}^{n-1} [\mathbf{y}_j^o - \mathbf{H}_j \mathbf{M}_{t_n, t_j} \mathbf{x}]^T \mathbf{R}_j^{-1} [\mathbf{y}_j^o - \mathbf{H}_j \mathbf{M}_{t_n, t_j} \mathbf{x}] = [\mathbf{x} - \bar{\mathbf{x}}_n^b]^T (\mathbf{P}_n^b)^{-1} [\mathbf{x} - \mathbf{x}_n^b] + c.$$

It follows that

$$J_{t_n}^o(\mathbf{x}) = [\mathbf{x} - \bar{\mathbf{x}}_n^b]^T (\mathbf{P}_n^b)^{-1} [\mathbf{x} - \bar{\mathbf{x}}_n^b] + [\mathbf{y}_n^o - \mathbf{H}_n \mathbf{x}]^T \mathbf{R}_n^{-1} [\mathbf{y}_n^o - \mathbf{H}_n \mathbf{x}] + c.$$
(7)

To complete the data assimilation cycle, we determine the state estimate $\bar{\mathbf{x}}_n^a$ and its covariance \mathbf{P}_n^a so that

$$J_{t_n}^o(\mathbf{x}) = [\mathbf{x} - \bar{\mathbf{x}}_n^a]^T (\mathbf{P}_n^a)^{-1} [\mathbf{x} - \bar{\mathbf{x}}_n^a] + c'$$

for some constant c'. Equating the terms of degree 2 in x, we get

$$\mathbf{P}_{n}^{a} = \left[(\mathbf{P}_{n}^{b})^{-1} + \mathbf{H}_{n}^{T} \mathbf{R}_{n}^{-1} \mathbf{H}_{n} \right]^{-1}.$$
(8)

Equating the terms of degree 1, we get

$$\bar{\mathbf{x}}_{n}^{a} = \mathbf{P}_{n}^{a} \left[(\mathbf{P}_{n}^{b})^{-1} \bar{\mathbf{x}}_{n}^{b} + \mathbf{H}_{n}^{T} \mathbf{R}_{n}^{-1} \mathbf{y}_{n}^{o} \right].$$
(9)

This equation in some sense (consider, for example, the case where H_n is the identity matrix) expresses the analysis state estimate as a weighted average of the background state estimate and the observations, weighted according to the inverse covariance of each.

Equations (8) and (9) can be written in many different but equivalent forms, and it will be useful later to rewrite both of them now. Using (8) to eliminate $(\mathbf{P}_n^b)^{-1}$ from (9) yields

$$\bar{\mathbf{x}}_n^a = \bar{\mathbf{x}}_n^b + \mathbf{P}_n^a \mathbf{H}_n^T \mathbf{R}_n^{-1} (\mathbf{y}_n^o - \mathbf{H}_n \bar{\mathbf{x}}_n^b).$$
(10)

The matrix $\mathbf{P}_n^a \mathbf{H}_n^T \mathbf{R}_n^{-1}$ is called the "Kalman gain". It multiplies the difference between the observations at time t_n and the values predicted by the background state estimate to yield the increment between the background and analysis state estimates. Next, multiplying (8) on the right by $(\mathbf{P}_n^b)^{-1}\mathbf{P}_n^b$ and combining the inverses yields

$$\mathbf{P}_n^a = (\mathbf{I} + \mathbf{P}_n^b \mathbf{H}_n^T \mathbf{R}_n^{-1} \mathbf{H}_n)^{-1} \mathbf{P}_n^b.$$
(11)

This expression is better than the previous one from a practical point of view, since it does not require inverting \mathbf{P}_n^b .

Initialization. The derivation above of the Kalman Filter avoided the issue of how to initialize the iteration. In order to solve the best fit problem we originally posed, we should make no assumptions about the system state prior to the analysis at time t_1 . Formally we can regard the background covariance \mathbf{P}_1^b to be infinite, and for n = 1 use (8) and (9) with $(\mathbf{P}_1^b)^{-1} = \mathbf{0}$. This works if there are enough observations at time t_1 to determine (aside from the measurement errors) the system state; that is, if \mathbf{H}_1 has rank equal to the number of model variables m. The analysis then determines $\bar{\mathbf{x}}_1^a$ in the appropriate least-square sense. However, if there are not enough observations, then the matrix to be inverted in (8) does not have full rank. To avoid this difficulty, one can assume a prior background distribution at time t_1 , with \mathbf{P}_1^b reasonably large but finite. This adds a small quadratic term to the cost function being minimized, but with sufficient observations over time, the effect of this term on the analysis at time t_n decreases in significance as t increases.

2.2 Nonlinear Scenario: Ensemble Kalman Filtering

Many approaches to data assimilation for nonlinear problems are based on the Kalman Filter, or at least on minimizing a cost function similar to (7). At a minimum, a nonlinear model forces a change in the forecast equations (5) and (6), while nonlinear observation operators H_n force a change in the analysis equations (10) and (11). The Extended Kalman Filter (see, for example, [18]) computes $\bar{\mathbf{x}}_n^b = M_{t_{n-1},t_n}(\bar{\mathbf{x}}_{n-1}^a)$ using the nonlinear model, but computes \mathbf{P}_n^b using the linearization \mathbf{M}_{t_{n-1},t_n} of M_{t_{n-1},t_n} around $\bar{\mathbf{x}}_{n-1}^a$. The analysis then uses the linearization \mathbf{H}_n of H_n around $\bar{\mathbf{x}}_n^b$. This approach is problematic for complex, high-dimensional models such as a global weather model for (at least) two reasons. First, it is not easy to linearize such a model. Second, the number of model variables m is several million, and as a result the $m \times m$ matrix inverse required by the analysis cannot be performed in a reasonable amount of time. Approaches used in operational weather forecasting generally eliminate for pragmatic reasons the time iteration of the Kalman Filter. The U.S. National Weather Service performs data assimilation every 6 hours using the "3D-VAR" method [25, 31], in which the background covariance \mathbf{P}_n^b in (7) is replaced by a constant matrix **B** representing typical uncertainty in a 6 hour forecast. This simplification allows the analysis to be formulated in a manner that does not require a large matrix to be inverted each time. The 3D-VAR cost function also includes a nonlinear observation operator H_n , and is minimized numerically to produce the analysis state estimate \mathbf{x}_n^a .

The "4D-VAR" method [24, 35] used by the European Centre for Medium-Range Weather Forecasts uses a cost function that includes a constant-covariance background term as in 3D-VAR together with a sum like (2) accounting for the observations collected over a 12 hour time span. Again the cost function is minimized numerically; this procedure is computationally intensive, because both the the nonlinear model and its linearization must be integrated over the 12 hour interval to compute the gradient of the 4D-VAR cost function, and this procedure is repeated until a satisfactory approximation to the minimum is found.

The key idea of ensemble Kalman filtering [6] is to choose at time t_{n-1} an ensemble of initial conditions whose spread around $\bar{\mathbf{x}}_{n-1}^a$ characterizes the analysis covariance \mathbf{P}_{n-1}^a , propagate each ensemble member using the nonlinear model, and compute \mathbf{P}_n^b based on the resulting ensemble at time t_n . Thus like the Extended Kalman Filter, the (approximate) uncertainty in the state estimate is propagated from one analysis to the next, unlike 3D-VAR (which does not propagate the uncertainty at all) or 4D-VAR (which propagates it only for a limited time). Furthermore, it does this without requiring a linearized model. While these are advantages over the other methods, there are some potential disadvantages as well.

Perhaps the most important difference between ensemble Kalman filtering and the other methods described above is that it only quantifies uncertainty in the space spanned by the ensemble. Assuming that computational resources restrict the number of ensemble members k to be much smaller than the number of model variables m, this can be a severe limitation. On the other hand, if this limitation can be overcome (see the section on "Localization" below), then the analysis can be performed in a much lower-dimensional space (k versus m). Thus, ensemble Kalman filtering has the potential to be more computationally efficient than the other methods. Indeed, the main point of this article is to describe how to do ensemble Kalman filtering efficiently without sacrificing accuracy.

Notation. We start with an ensemble $\{\mathbf{x}_{n-1}^{a(i)} : i = 1, 2, ..., k\}$ of *m*-dimensional model state vectors at time t_{n-1} . One approach would be to let one of the ensemble members represent the best estimate of the system state, but here we assume the ensemble to be chosen so that its average represents the analysis state estimate. We evolve each ensemble member according to the nonlinear

model to obtain a background ensemble $\{\mathbf{x}_n^{b(i)} : i = 1, 2, ..., k\}$ at time t_n :

$$\mathbf{x}_{n}^{b(i)} = M_{t_{n-1},t_n}(\mathbf{x}_{n-1}^{a(i)}).$$

For the rest of this article, I will discuss what to do at the analysis time t_n , and so I now drop the subscript n. Thus, for example, H and \mathbf{R} will represent respectively the observation operator and the observation error covariance matrix at the analysis time. Let ℓ be the number of scalar observations used in the analysis.

For the background state estimate and its covariance we use the sample mean and covariance of the background ensemble:

$$\bar{\mathbf{x}}^{b} = k^{-1} \sum_{i=1}^{k} \mathbf{x}^{b(i)},$$
$$\mathbf{P}^{b} = (k-1)^{-1} \sum_{i=1}^{k} (\mathbf{x}^{b(i)} - \bar{\mathbf{x}}^{b}) (\mathbf{x}^{b(i)} - \bar{\mathbf{x}}^{b})^{T} = (k-1)^{-1} \mathbf{X}^{b} (\mathbf{X}^{b})^{T},$$
(12)

where \mathbf{X}^{b} is the $m \times k$ matrix whose *i*th column is $\mathbf{x}^{b(i)} - \bar{\mathbf{x}}^{b}$. The analysis must determine not only an state estimate $\bar{\mathbf{x}}^{a}$ and covariance \mathbf{P}^{a} , but also an ensemble { $\mathbf{x}^{a(i)} : i = 1, 2, ..., k$ } with the appropriate sample mean and covariance:

$$\bar{\mathbf{x}}^a = k^{-1} \sum_{i=1}^k \mathbf{x}^{a(i)}$$

$$\mathbf{P}^{a} = (k-1)^{-1} \sum_{i=1}^{k} (\mathbf{x}^{a(i)} - \bar{\mathbf{x}}^{a}) (\mathbf{x}^{a(i)} - \bar{\mathbf{x}}^{a})^{T} = (k-1)^{-1} \mathbf{X}^{a} (\mathbf{X}^{a})^{T},$$
(13)

where \mathbf{X}^a is the $m \times k$ matrix whose *i*th column is $\mathbf{x}^{a(i)} - \bar{\mathbf{x}}^a$.

In Section 2.3, I will describe how to determine $\bar{\mathbf{x}}^a$ and \mathbf{P}^a for a (possibly) nonlinear observation operator H in a way that agrees with the Kalman Filter equations (10) and (11) in the case that H is linear.

Choice of analysis ensemble. Once $\bar{\mathbf{x}}^a$ are \mathbf{P}^a specified, there are still many possible choices of an analysis ensemble (or equivalently, a matrix \mathbf{X}^a that satisfies (13) and the sum of whose columns is zero). A variety of ensemble Kalman filters have been proposed, and one of the main differences among them is how the analysis ensemble is chosen. The simplest approach is to apply the Kalman filter update (10) separately to each background ensemble member (rather than their mean) to get the corresponding analysis ensemble member. However, this results in an analysis ensemble whose sample covariance is smaller than \mathbf{P}^a , unless the observations are artificially perturbed so that each ensemble member is updated using different random realization of the perturbed but deterministic algorithms to generate an analysis ensemble with the desired sample mean and covariance. As such,

their analyses coincide exactly with the standard Kalman Filter in the linear scenario of the previous section. I will use this deterministic approach below.

Localization. Another important issue in ensemble Kalman filtering of spatiotemporally chaotic systems is spatial localization. If the ensemble has k members, then the background covariance matrix \mathbf{P}^{b} given by (12) describes nonzero uncertainty only in the k-dimensional subspace spanned by the ensemble, and a global analysis will allow adjustments to the system state only in this subspace. If the system is high-dimensionally unstable, then forecast errors will grow in directions not accounted for by the ensemble, and these errors will not be corrected by the analysis. On the other hand, in a sufficiently small local region, the system may behave like a low-dimensionally unstable system driven by the dynamics in neighboring region; such behavior was observed for a global weather model in [32]. Performing the analysis locally requires only that the ensemble in some sense span the local unstable space; by allowing the local analyses to choose different linear combinations of the ensemble members in different regions, the global analysis is not confined to the k-dimensional ensemble space and instead explores a much higher dimensional space [9, 29, 30]. Another view on the necessity of localization for spatiotemporally chaotic systems is that the limited sample size provided by an ensemble will produce spurious correlations between distant locations in the background covariance matrix \mathbf{P}^{b} [14, 12]. Unless they are suppressed, these spurious correlations will cause observations from one location to affect, in an essentially random manner, the analysis an arbitrarily large distance away. If the system has a characteristic "correlation distance", then the analysis should ignore ensemble correlations over much larger distances. In addition to providing better results in many cases, localization allows the analysis to be done more efficiently as a parallel computation [23, 29, 30].

Localization is generally done either explicitly, considering only the observations from a region surrounding the location of the analysis [22, 14, 23, 1, 29, 30], or implicitly, by multiplying the entries in \mathbf{P}^b by a distance-dependent function that decays to zero beyond a certain distance, so that observations do not affect the model state beyond that distance [15, 12, 38]. I will follow the explicit approach here, doing a separate analysis for each spatial grid point of the model. (My use of "grid point" assumes the model to be a discretization of a partial differential equation, or otherwise be defined on a lattice, but the method is also applicable to systems with other geometries.) The choice of which observations to use for each grid point is up to the user of the method, and a good choice will depend both on the particular system being modeled and the size of the ensemble (more ensemble members will generally allow more distant observations to be used gainfully). It is important however to have significant overlap between the observations used for one grid point and the observations used for a neighboring grid point; otherwise the analysis ensemble may change suddenly from one grid point to the next. For an atmospheric model, a reasonable approach would

be to use observations within a cylinder of a given radius and height and determine empirically which values of these two parameters work best. At its simplest, the method I describe gives all of the chosen observations the same weight, but I will also describe how to make the weights given to the observations decay more smoothly to zero as the distance from the analysis location.

2.3 LETKF: A local ensemble transform Kalman filter

I will now describe an efficient means of performing the analysis that transforms a background ensemble $\{\mathbf{x}^{b(i)} : i = 1, 2, ..., k\}$ into an appropriate analysis ensemble $\{\mathbf{x}^{a(i)} : i = 1, 2, ..., k\}$, using the notation defined above. I assume that the number of ensemble members k is smaller than both the number of model variables m and the number of observations ℓ , even when localization has reduced the effective values of m and ℓ considerably compared to a global analysis. (In this section I will assume the choice of observations to use for the local analysis to have been performed already, and consider \mathbf{y}^o , H and **R** to be truncated to these observations; as such, correlations between errors in the chosen observations and errors in other observations are ignored.) Most of the analysis will take place in a k-dimensional space, with as few operations as possible in the model and observation spaces.

Formally, we want the analysis mean $\bar{\mathbf{x}}^a$ to minimize the Kalman filter cost function (7), modified to allow for a nonlinear observation operator *H*:

$$J(\mathbf{x}) = (\mathbf{x} - \bar{\mathbf{x}}^b)^T (\mathbf{P}^b)^{-1} (\mathbf{x} - \bar{\mathbf{x}}^b) + [\mathbf{y}^o - H(\mathbf{x})]^T \mathbf{R}^{-1} [\mathbf{y}^o - H(\mathbf{x})].$$
(14)

However, the $m \times m$ background covariance matrix $\mathbf{P}^b = (k-1)^{-1} \mathbf{X}^b (\mathbf{X}^b)^T$ can have rank at most k-1, and is therefore not invertible. Nonetheless, its inverse is well-defined on the space S spanned by the background ensemble perturbations, that is, the columns of \mathbf{X}^b . Thus J is also well-defined for $\mathbf{x} - \bar{\mathbf{x}}^b$ in S, and the minimization can be carried out in this subspace.¹ As we have said, this reduced dimensionality is an advantage from the point of view of efficiency, though the restriction of the analysis mean to S is sure to be detrimental if k is too small.

In order to perform the analysis on S, we must choose an appropriate coordinate system. A natural approach is to use the singular vectors of \mathbf{X}^{b} (the eigenvectors of \mathbf{P}^{b}) to form a basis for S [1, 29, 30]. Here we avoid this step by using instead the columns of \mathbf{X}^{b} to span S, as in [3]. One

¹Considerably more general cost functions could be used in the relatively low-dimensional ensemble space S. In particular, one could consider non-Gaussian background distributions as follows. Given a distribution that can be parametrized solely by a mean and covariance matrix, substitute the negative logarithm of its probability distribution function for the first term on the right side of (14). Though the formulas below that determine the analysis mean and covariance would not be valid, one could numerically determine the appropriate mean and covariance. In principle, distributions that are parametrized by higher order moments could be considered, but this would require significantly larger ensembles.

conceptual difficulty in this approach is that the sum of these columns is zero, so they are necessarily not linearly independent. We could assume the first k - 1 columns to be independent and use them as a basis, but this assumption is unnecessary and clutters the resulting equations. Instead, we regard \mathbf{X}^b as a linear transformation from a k-dimensional space \tilde{S} onto S, and perform the analysis in \tilde{S} . Let w denote a vector in \tilde{S} ; then \mathbf{X}^b w belongs to the space S spanned by the background ensemble perturbations, and $\mathbf{x} = \bar{\mathbf{x}}^b + \mathbf{X}^b$ w is the corresponding model state.

Notice that if w is a Gaussian random vector with mean 0 and covariance $(k - 1)^{-1}\mathbf{I}$, then $\mathbf{x} = \bar{\mathbf{x}}^b + \mathbf{X}^b \mathbf{w}$ is Gaussian with mean $\bar{\mathbf{x}}^b$ and covariance $\mathbf{P}^b = (k - 1)^{-1} \mathbf{X}^b (\mathbf{X}^b)^T$. This motivates the cost function

$$\tilde{J}(\mathbf{w}) = (k-1)\mathbf{w}^T\mathbf{w} + [\mathbf{y}^o - H(\bar{\mathbf{x}}^b + \mathbf{X}^b\mathbf{w})]^T\mathbf{R}^{-1}[\mathbf{y}^o - H(\bar{\mathbf{x}}^b + \mathbf{X}^b\mathbf{w})]$$
(15)

on \tilde{S} . In particular, I claim that if $\bar{\mathbf{w}}^a$ minimizes \tilde{J} , then $\bar{\mathbf{x}}^a = \bar{\mathbf{x}}^b + \mathbf{X}^b \bar{\mathbf{w}}^a$ minimizes the cost function J. Substituting the change of variables formula into (14) and using (12) yields the identity

$$\tilde{J}(\mathbf{w}) = (k-1)\mathbf{w}^T (\mathbf{I} - (\mathbf{X}^b)^T [\mathbf{X}^b (\mathbf{X}^b)^T]^{-1} \mathbf{X}^b) \mathbf{w} + J(\bar{\mathbf{x}}^b + \mathbf{X}^b \mathbf{w}).$$
(16)

The matrix $\mathbf{I} - (\mathbf{X}^b)^T [\mathbf{X}^b (\mathbf{X}^b)^T]^{-1} \mathbf{X}^b$ is the orthogonal projection onto the null space N of \mathbf{X}^b . (Generally N will be one-dimensional, spanned by the vector $(1, 1, ..., 1)^T$, but it could be higherdimensional.) Thus, the first term on the right side of (16) depends only on the component of \mathbf{w} in N, while the second term depends only on its component in the space orthogonal to N (which is in one-to-one correspondence with S under \mathbf{X}^b). Thus if $\bar{\mathbf{w}}^a$ minimizes \tilde{J} , then it must be orthogonal to N, and the corresponding vector $\bar{\mathbf{x}}^a$ minimizes J.

Nonlinear Observations. The most accurate way to allow for a nonlinear observation operator H would be to numerically minimize \tilde{J} in the k-dimensional space \tilde{S} , as in [40]. If H is sufficiently nonlinear, then \tilde{J} could have multiple minima, but a numerical minimization using $\mathbf{w} = \mathbf{0}$ (corresponding to $\mathbf{x} = \bar{\mathbf{x}}^b$) as an initial guess would still be a reasonable approach. Having determined $\bar{\mathbf{w}}^a$ in this manner, one would compute the analysis covariance $\tilde{\mathbf{P}}^a$ in \tilde{S} from the second partial derivatives of \tilde{J} at $\bar{\mathbf{w}}^a$, then use \mathbf{X}^b to transform the analysis results into the model space, as below. But in order to formulate the analysis more explicitly, we now linearize H about the background ensemble mean $\bar{\mathbf{x}}^b$. Of course, if H is linear then we will find the minimum of \tilde{J} exactly. And if the spread of the background ensemble is not too large, the linearization should be a decent approximation, similar to the approximation we have already made that a linear combination of background ensemble members is also a plausible background model state.

Since we only need to evaluate H in the ensemble space (or equivalently to evaluate $H(\bar{\mathbf{x}}^b + \mathbf{X}^b \mathbf{w})$ for \mathbf{w} in \tilde{S}), the simplest way to linearize H is to apply it to each of the ensemble members $\mathbf{x}^{b(i)}$ and interpolate. To this end, we define an ensemble $\mathbf{y}^{b(i)}$ of background observation vectors by

$$\mathbf{y}^{b(i)} = H(\mathbf{x}^{b(i)}). \tag{17}$$

We define also their mean $\bar{\mathbf{y}}^b$, and the $\ell \times k$ matrix \mathbf{Y}^b whose *i*th column is $\mathbf{y}^{b(i)} - \bar{\mathbf{y}}^b$. We then make the linear approximation

$$H(\bar{\mathbf{x}}^b + \mathbf{X}^b \mathbf{w}) \approx \bar{\mathbf{y}}^b + \mathbf{Y}^b \mathbf{w}.$$
(18)

The same approximation is used in, for example, [15], and is equivalent to the joint state-observation space method in [1].

Analysis. The linear approximation we have just made yields the quadratic cost function

$$\tilde{J}^*(\mathbf{w}) = (k-1)\mathbf{w}^T\mathbf{w} + [\mathbf{y}^o - \bar{\mathbf{y}}^b - \mathbf{Y}^b\mathbf{w}]^T\mathbf{R}^{-1}[\mathbf{y}^o - \bar{\mathbf{y}}^b - \mathbf{Y}^b\mathbf{w}].$$
(19)

This cost function is in the form of the Kalman filter cost function (7), using the background mean $\bar{\mathbf{w}}^b = \mathbf{0}$ and background covariance matrix $\tilde{\mathbf{P}}^b = (k-1)^{-1}\mathbf{I}$, with \mathbf{Y}^b playing the role of the observation operator. The analogues of the analysis equations (10) and (11) are then

$$\mathbf{w}^{a} = \tilde{\mathbf{P}}^{a} (\mathbf{Y}^{b})^{T} \mathbf{R}^{-1} (\mathbf{y}^{o} - \bar{\mathbf{y}}^{b}),$$
(20)

$$\tilde{\mathbf{P}}^{a} = [(k-1)\mathbf{I} + (\mathbf{Y}^{b})^{T}\mathbf{R}^{-1}\mathbf{Y}^{b}]^{-1}.$$
(21)

In model space, the analysis mean and covariance are then

$$\bar{\mathbf{x}}^a = \bar{\mathbf{x}}^b + \mathbf{X}^b \bar{\mathbf{w}}^a,\tag{22}$$

$$\mathbf{P}^a = \mathbf{X}^b \tilde{\mathbf{P}}^a (\mathbf{X}^b)^T.$$
(23)

In order to initiate the ensemble forecast that will produce the background for the next analysis, we must choose an analysis ensemble whose sample mean and covariance are equal to $\bar{\mathbf{x}}^a$ and \mathbf{P}^a . As mentioned above, this amounts to choosing a matrix \mathbf{X}^a so that the sum of its columns is zero and (13) holds. Then one can form the analysis ensemble by adding $\bar{\mathbf{x}}^a$ to each of the columns of \mathbf{X}^a .

Symmetric Square Root. Our choice of analysis ensemble is described by $X^a = X^b W^a$, where

$$\mathbf{W}^a = [(k-1)\tilde{\mathbf{P}}^a]^{1/2} \tag{24}$$

and by the 1/2 power of a symmetric matrix we mean its symmetric square root. Then $\tilde{\mathbf{P}}^a = (k-1)^{-1}\mathbf{W}^a(\mathbf{W}^a)^T$, and (13) follows from (23). The use of the symmetric square root to determine \mathbf{W}^a from $\tilde{\mathbf{P}}^a$ (as compared to, for example, a Cholesky factorization, or the choice described in [3]), is important for two main reasons. First, as we will see below, it ensures that the sum of the columns of \mathbf{X}^a is zero, so that the analysis ensemble has the correct sample mean (this is also shown for the symmetric square root in [37]). Second, it ensures that \mathbf{W}^a depends continuously on $\tilde{\mathbf{P}}^a$; while this

may be a desirable property in general, it is crucial in a local analysis scheme, so that neighboring grid points with slightly different matrices $\tilde{\mathbf{P}}^a$ do not yield very different analysis ensembles.

Another potentially desirable property of the symmetric square root is that it minimizes the (mean-square) distance between W^a and the identity matrix [29, 30], though because of the different choice of basis, it does not minimize the same quantity, and thus does not yield the same analysis ensemble, as in that article. Numerical experiments to be published elsewhere produce similar quality results to other reasonable choices of the analysis ensemble in a square-root filter; see Section 5.

To see that the sum of the columns of \mathbf{X}^a is zero, we express this condition as $\mathbf{X}^a \mathbf{v} = \mathbf{0}$, where **v** is a column vector of k ones: $\mathbf{v} = (1, 1, ..., 1)^T$. Notice that by (21), **v** is an eigenvector of $\tilde{\mathbf{P}}^a$ with eigenvalue $(k-1)^{-1}$:

$$(\tilde{\mathbf{P}}^a)^{-1}\mathbf{v} = [(k-1)\mathbf{I} + (\mathbf{Y}^b)^T \mathbf{R}^{-1} \mathbf{Y}^b]\mathbf{v} = (k-1)\mathbf{v},$$

because the sum of the columns of \mathbf{Y}^b is zero. Then by (24), \mathbf{v} is also an eigenvector of \mathbf{W}^a with eigenvalue 1. Since the sum of the columns of \mathbf{X}^b is zero, $\mathbf{X}^a \mathbf{v} = \mathbf{X}^b \mathbf{W}^a \mathbf{v} = \mathbf{X}^b \mathbf{v} = \mathbf{0}$ as desired.

Finally, notice that we can form the analysis ensemble first in \tilde{S} by adding $\bar{\mathbf{w}}^a$ to each of the columns of \mathbf{W}^a ; let $\{\mathbf{w}^{a(i)}\}$ be the columns of the resulting matrix. These "weight" vectors specify what linear combinations of the background ensemble perturbations to add to the background mean in order to get the analysis ensemble in model space:

$$\mathbf{x}^{a(i)} = \bar{\mathbf{x}}^b + \mathbf{X}^b \mathbf{w}^{a(i)},\tag{25}$$

Local Implementation. Notice that once the background ensemble has been used to form $\bar{\mathbf{y}}^b$ and \mathbf{Y}^b , it is no longer needed in the analysis, except in (25) to translate the results from \tilde{S} to model space. This point is useful to keep in mind when implementing a local filter that computes a separate analysis for each model grid point. In principle, one should form a global background observation ensemble $\mathbf{y}_{[g]}^{b(i)}$ from the global background vectors, though in practice this can be done locally when the global observation operator $H_{[g]}$ uses local interpolation. After the background observation ensemble is formed, the analyses at different grid points are completely independent of each other and can be computed in parallel. The observations chosen for a given grid point will dictate which coordinates of $\mathbf{y}_{[g]}^{b(i)}$ are used to form the local background observation ensemble $\mathbf{y}^{b(i)}$ for that analysis in \tilde{S} will produce the weight vectors { $\mathbf{w}^{a(i)}$ } for that grid point. Thus computing the analysis ensemble { $\mathbf{x}^{a(i)}$ } for that grid point using (25) requires only using the background model states at that grid point.

As long as the sets of observations used at a pair of neighboring grid points overlap heavily, the linear combinations used at the two grid points will be similar, and thus the global analysis ensemble members formed by these spatially varying linear combinations will change slowly from one grid point to the next. In a local region of several grid points, they will approximately be linear combinations of the background ensemble members, and thus should represent reasonably "physical" initial conditions for the forecast model. However, if the model requires of its initial conditions high-order smoothness and/or precise conformance to an conservation law, it may be necessary to post-process the analysis ensemble members to smooth them and/or project them onto the manifold determined by the conserved quantities before using them as initial conditions (this procedure is often called "balancing" in geophysical data assimilation).

In other localization approaches [15, 12, 38], the influence of an observation at a particular point on the analysis at a particular model grid point decays smoothly to zero as the distance between the two points increases. A similar effect can be achieved here by multiplying the entries in the inverse observation error covariance matrix \mathbf{R}^{-1} by a factor that decays from one to zero as the distance of the observations from the analysis grid point increases. This "smoothed localization" corresponds to gradually increasing the uncertainty assigned to the observations until beyond a certain distance they have infinite uncertainty and therefore no influence on the analysis.

Covariance Inflation. In practice, an ensemble Kalman filter that adheres strictly to the Kalman Filter equations (10) and (11) may fail to synchronize with the "true" system trajectory that produces the observations. One reason for this is model error, but even with a perfect model the filter will tend to underestimate the uncertainty in its state estimate [38]. Regardless of the cause, underestimating the uncertainty leads to overconfidence in the background state estimate, and hence underweighting of the observations by the analysis. If the discrepancy becomes too large over time, the observations are essentially ignored by the analysis, and the dynamics of the data assimilation system become decoupled from the truth.

Generally this effect is countered by an ad hoc procedure (with at least one tunable parameter) that inflates either the background covariance or the analysis covariance during each data assimilation cycle. One "hybrid" approach adds a multiple of the background covariance matrix **B** from the 3D-VAR method to the background covariance prior to the analysis [11]. "Multiplicative inflation" [2, 12] instead multiplies the background covariance matrix (or equivalently, the differences or "perturbations" between the background ensemble members and their mean) by a constant factor larger than one. "Additive inflation" adds a small multiple of the identity matrix to either the background covariance during each cycle [29, 30]. Finally, if one chooses the analysis ensemble in such a way that each member has a corresponding member of the background ensemble; replacing each analysis perturbation from the mean by a weighted average of itself and the corresponding background perturbation [39].

Within the framework described in this article, the hybrid approach is not feasible because it

requires the analysis to consider uncertainty outside the space spanned by the background ensemble. However, once the analysis ensemble is formed, one could develop a means of inflating it in directions (derived from the 3D-VAR background covariance matrix **B** or otherwise) outside the ensemble space so that uncertainty in these directions is reflected in the background ensemble at the next analysis step. Additive inflation is feasible, but requires substantial additional computation in order to determine the adjustment necessary in the k-dimensional space \tilde{S} that corresponds to adding a multiple of the identity matrix to the model space covariance \mathbf{P}^b or \mathbf{P}^a . Relaxation is simple to implement, and is most efficiently done in \tilde{S} by replacing \mathbf{W}^a with a weighted average of it and the identity matrix.

Multiplicative inflation can be performed most easily on the analysis ensemble by multiplying \mathbf{W}^a by an appropriate factor (namely $\sqrt{\rho}$ in order to multiply the analysis covariance by ρ). To perform multiplicative inflation on the background covariance instead, one should theoretically multiply \mathbf{X}^b by such a factor, and adjust the background ensemble $\{\mathbf{x}^{b(i)}\}$ accordingly before applying the observation operator H to form the background observation ensemble $\{\mathbf{y}^{b(i)}\}$. However, a more efficient approach, which is equivalent if H is linear, and is a close approximation even for nonlinear H if the inflation factor ρ is close to one, is simply to replace $(k - 1)\mathbf{I}$ by $(k - 1)\mathbf{I}/\rho$ in (21), since $(k - 1)\mathbf{I}$ is the inverse of the background covariance matrix $\tilde{\mathbf{P}}^b$ in the k-dimensional space \tilde{S} . One can check that this has the same effect on the analysis mean $\bar{\mathbf{x}}^a$ and covariance \mathbf{P}^a as multiplying \mathbf{X}^b and \mathbf{Y}^b by $\sqrt{\rho}$.

3 Efficient Computation of the Analysis

Here is a step-by-step description of how to perform the analysis described in the previous section, designed for efficiency both in ease of implementation and in the amount of computation and memory usage. Of course there are some trade-offs between these objectives, so in each step I first describe the simplest approach and then in some cases mention alternate approaches and possible gains in computational efficiency. I also give a rough accounting of the computational complexity of each step, and at the end discuss the overall computational complexity. After that, I describe an approach that in some cases will produce a significantly faster analysis, at the expense of more memory usage and more difficult implementation, by reorganizing some of the steps. As before, I will use "grid point" in this section to mean a spatial location in the forecast model, whether or not the model is actually based on a grid geometry; I will use "array" to mean a vector or matrix. The use of "columns" and "rows" below is for exposition only; one should of course store arrays in whatever manner is most efficient for one's computing environment.

The inputs to the analysis are a background ensemble of $m_{[g]}$ -dimensional model state vectors $\{\mathbf{x}_{[g]}^{b(i)} : i = 1, 2, ..., k\}$, a function $H_{[g]}$ from the $m_{[g]}$ -dimensional model space to the $\ell_{[g]}$ -

dimensional observation space, an $\ell_{[g]}$ -dimensional vector $\mathbf{y}_{[g]}^{o}$ of observations, and an $\ell_{[g]} \times \ell_{[g]}$ observation error covariance matrix $\mathbf{R}_{[g]}$. The subscript g here signifies that these inputs reflect the global model state and all available observations, from which a local subset should be chosen for each local analysis. How to choose which observations to use is entirely up to the user of this method, but a reasonable general approach is to choose those observations made within a certain distance of the grid point at which one is doing the local analysis and determine empirically which value of the cutoff distance produces the "best" results. If one deems localization to be unnecessary in a particular application, then one can ignore the distinction between local and global, and skip Steps 3 and 9 below.

Steps 1 and 2 are essentially global operations, but may be done locally in a parallel implementation. Steps 3–8 should be performed separately for each local analysis (generally this means for each grid point, but see the parenthetical comment at the end of Step 3). Step 9 simply combines the results of the local analyses to form a global analysis ensemble $\{\mathbf{x}_{[g]}^{a(i)}\}$, which is the final output of the analysis.

- 1. Apply $H_{[g]}$ to each $\mathbf{x}_{[g]}^{b(i)}$ to form the global background observation ensemble $\{\mathbf{y}_{[g]}^{b(i)}\}$, and average the latter vectors to get the $\ell_{[g]}$ -dimensional column vector $\bar{\mathbf{y}}_{[g]}^{b}$. Subtract this vector from each $\{\mathbf{y}_{[g]}^{b(i)}\}$ to form the columns of the $\ell_{[g]} \times k$ matrix $\mathbf{Y}_{[g]}^{b}$. (This subtraction can be done "in place", since the vectors $\{\mathbf{y}_{[g]}^{b(i)}\}$ are no longer needed.) This requires k applications of H, plus $2k\ell_{[g]}$ (floating-point) operations. If H is an interpolation operator that requires only a few model variables to compute each observation variable, then the total number of operations for this step is proportional to $k\ell_{[g]}$ times the average number of model variables required to compute each scalar observation.
- 2. Average the vectors $\{\mathbf{x}_{[g]}^{b(i)}\}$ to get the $m_{[g]}$ -dimensional vector $\bar{\mathbf{x}}_{[g]}^{b}$, and subtract this vector from each $\mathbf{x}_{[g]}^{b(i)}$ to form the columns of the $m_{[g]} \times k$ matrix $\mathbf{X}_{[g]}^{b}$. (Again the subtraction can be done "in place"; the vectors $\{\mathbf{x}_{[g]}^{b(i)}\}$ are no longer needed). This step requires a total of $2km_{[g]}$ operations. (If H is linear, one can equivalently perform Step 2 before Step 1, and obtain $\bar{\mathbf{y}}_{[g]}^{b}$ and $\mathbf{Y}_{[g]}^{b}$ by applying H to $\bar{\mathbf{x}}_{[g]}^{b}$ and $\mathbf{X}_{[g]}^{b}$.)
- 3. This step selects the necessary data for a given grid point (whether it is better to form the local arrays described below explicitly or select them later as needed from the global arrays depends on one's implementation). Select the rows of x
 ^b_[g] and X^b_[g] corresponding to the given grid point, forming their local counterparts: the m-dimensional vector x
 ^b and the m × k matrix X^b, which will be used in Step 8. Likewise, select the rows of y
 ^b_[g] and Y^b_[g] corresponding to the d-dimensional vector y
 ^b and the ℓ×k matrix Y^b. Select the corresponding rows of y
 ^o_[g] and rows and columns of R_[g] to form the ℓ-dimensional vector y^o and the ℓ×ℓ matrix R. (For a high-resolution

model, it may be reasonable to use the same set of observations for multiple grid points, in which case one should select here the rows of $\mathbf{X}_{[g]}^b$ and $\bar{\mathbf{x}}_{[g]}^b$ corresponding to all of these grid points.)

- 4. Compute the k × ℓ matrix C = (Y^b)^TR⁻¹. Since this is the only step in which R is used, it may be most efficient to compute C by solving the linear system RC^T = Y^b rather than inverting R. In some applications, R may be diagonal, but in others R will be block diagonal with each block representing a group of correlated observations. As long as the size of each block is relatively small, inverting R or solving the linear system above will not be computationally expensive. Furthermore, many or all of the blocks that make up R may be unchanged from one analysis time to the next, so that their inverses need not be recomputed each time. Based on these considerations, the number of operations required (at each grid point) for this step in a typical application should be proportional to kℓ, multiplied by a factor related to the typical block size of R.
- 5. Compute the $k \times k$ matrix $\tilde{\mathbf{P}}^a = \left[(k-1)\mathbf{I}/\rho + \mathbf{C}\mathbf{Y}^b \right]^{-1}$, as in (21). Here $\rho > 1$ is a multiplicative covariance inflation factor, as described at the end of the previous section. Though trying some of the other approaches described there may be fruitful, a reasonable general approach is to start with $\rho > 1$ and increase it gradually until one finds a value that is optimal according to some measure of analysis quality. Multiplying \mathbf{C} and \mathbf{Y}^b requires less than $2k^2\ell$ operations, while the number of operations needed to invert the $k \times k$ matrix is proportional to k^3 .
- 6. Compute the $k \times k$ matrix $\mathbf{W}^a = [(k-1)\tilde{\mathbf{P}}^a]^{1/2}$, as in (24). Again the number of operations required is proportional to k^3 ; it may be most efficient to compute the eigenvalues and eigenvectors of $[(k-1)\mathbf{I}/\rho + \mathbf{C}\mathbf{Y}^b]$ in the previous step and then use them to compute both $\tilde{\mathbf{P}}^a$ and \mathbf{W}^a .
- Compute the k-dimensional vector w^a = P̃^aC(y^o − ȳ^b), as in (20), and add it to each column of W^a, forming a k × k matrix whose columns are the analysis vectors {w^{a(i)}}. Computing the formula for w^a from right-to-left, the total number of operations required for this step is less than 3k(ℓ + k).
- 8. Multiply \mathbf{X}^{b} by each $\mathbf{w}^{a(i)}$ and add $\bar{\mathbf{x}}^{b}$ to get the analysis ensemble members $\{\mathbf{x}^{a(i)}\}$ at the analysis grid point, as in (25). This requires $2k^{2}m$ operations.
- 9. After performing Steps 3–8 for each grid point, the outputs of Step 8 form the global analysis ensemble $\{\mathbf{x}_{[q]}^{a(i)}\}$.

We now summarize the overall computational complexity of the algorithm described above. If p is the number local analyses performed (equal to the number of grid points in the most basic approach), then notice that $pm = m_{[g]}$, while $p\bar{\ell} = q\ell_{[g]}$, where $\bar{\ell}$ is the average number of observations used in a local analysis and q is the average number of local analyses in which a particular observation is used. If $\bar{\ell}$ is large compared to k and m, then the most computationally expensive step is either Step 5, requiring approximately $2k^2p\bar{\ell} = 2k^2q\ell_{[g]}$ operations over all the local analyses, or Step 4, whose overall number of operations is proportional to $kp\bar{\ell} = kq\ell_{[g]}$, but with a proportionality constant dependent on the correlation structure of $\mathbf{R}_{[g]}$. In any case, as long as the typical number of correlated observations in a block of $\mathbf{R}_{[g]}$ remains constant, the overall computation time grows at most linearly with the total number $\ell_{[g]}$ of observations. It also grows at most linearly with the total number $\ell_{[g]}$ of observations. It also grows at most linearly with the total number $\ell_{[g]}$ overall operations. The computation time tends to be roughly quadratic in the number k of ensemble members, though for a very large ensemble the terms of order k^3 above would become significant.

Batch Processing of Observations. Some of the steps above have a q-fold redundancy, in that computations involving a given observation are repeated over an average of q different local analyses. For a general observation error covariance matrix $\mathbf{R}_{[a]}$ this redundancy may be unavoidable, but it can be avoided as described below if the global observations can be partitioned into local groups (or "batches") numbered 1, 2, ... that meet the following conditions. First, all of the observations in a given batch must be used in the exact same subset of the local analyses. Second, observations in different batches must have uncorrelated errors, so that each batch j corresponds to a block \mathbf{R}_{j} in a block diagonal decomposition of $\mathbf{R}_{[q]}$. (These conditions can always met if $\mathbf{R}_{[q]}$ is diagonal, by making each batch consist of a single observation. However, as explained below, for efficiency one should make the batches as large as possible while still meeting the first condition.) Then at Step 3, instead of selecting (overlapping) submatrices of $\bar{\mathbf{y}}_{[g]}^{b}$, $\mathbf{Y}_{[g]}^{b}$, $\mathbf{y}_{[g]}^{o}$, and $\mathbf{R}_{[g]}$, for each grid point, let $\bar{\mathbf{y}}_{j}^{b}$, $\mathbf{Y}_{j}^{b}, \mathbf{y}_{j}^{o}$, represent the rows corresponding to the observations in batch j, and do the following for each batch. Compute and store the $k \times k$ matrix $\mathbf{C}_j \mathbf{Y}_i^b$ and the k-dimensional vector $\mathbf{C}_j (\mathbf{y}_i^o - \bar{\mathbf{y}}_i^b)$, where $\mathbf{C}_j = (\mathbf{Y}_j^b)^T \mathbf{R}_j^{-1}$ as in Step 4. (This can be done separately for each batch, in parallel, and the total number of operations is roughly $2k^2\ell_{[g]}$.) Then do Steps 5–8 separately for each local analysis; when $\mathbf{C}\mathbf{Y}^b$ and $\mathbf{C}(\mathbf{y}^o - \bar{\mathbf{y}}^b)$ are required in Steps 5 and 7, compute them by summing the corresponding arrays $C_j Y_j^b$ and $C_j (y_j^o - \bar{y}_j^b)$ over the batches j of observations that are used in the local analysis. To avoid redundant addition in these steps, batches that are used in exactly the same subset of the local analyses should be combined into a single batch. The total number of operations required by the summations over batches roughly k^2ps , where s is the average number of batches used in each local analysis. Both this and the $2k^2\ell_{[q]}$ operations described before are smaller than

the roughly $2k^2p\bar{\ell}=2k^2q\ell_{[g]}$ operations they combine to replace.

This approach has similarities with the "sequential" approach of [15] and [38], in which observations are divided into uncorrelated batches and a separate analysis is done for each batch; the analysis is done in the observation space whose dimension is the number of observations in a batch. However, in the sequential approach, the analysis ensemble for one batch of observations is used as the background ensemble for the next batch of observations. Since batches with disjoint local regions of influence can be analyzed separately, some parallelization is possible, though the LETKF approach described above is more easily distributed over a large number of processors. For a serial implementation, either approach may be faster depending on the application and the ensemble size.

4 Asynchronous Observations: 4D-LETKF

In theory, one can perform a new analysis each time new observations are made. In practice, this is a good approach if observations are made at regular and sufficiently infrequent time intervals. However, in many applications, such as weather forecasting, observations are much too frequent for this approach. Imagine, for example, a 6 hour interval between analyses, like at the National Weather Service. Since weather can change significantly over such a time interval, it is important to consider observations taken at intermediate times in a more sophisticated manner than to pretend they occurred at the analysis time (or to simply ignore them). Operational versions of 3D-VAR and 4D-VAR (see Section 2.2) do take into account the timing of the observations, and one of the primary strengths of 4D-VAR is that it does so in a precise manner, by considering which forecast model trajectory best fits the observations over a given time interval (together with assumed background statistics at the start of this interval).

We have seen that the analysis step in an ensemble Kalman filter considers model states that are linear combinations of the background ensemble states at the analysis time, and compares these model states to observations taken at the analysis time. Similarly, we can consider approximate model trajectories that are linear combinations of the background ensemble trajectories over an interval of time, and compare these approximate trajectories with the observations taken over that time interval. Instead of asking which model trajectory best fits the observations, we ask which linear combination of the background ensemble trajectories best fits the observations. As before, this is relatively a low-dimensional optimization problem that is much more computationally tractable than the full nonlinear problem.

This approach is similar to that of an ensemble Kalman smoother [7, 8], but over a much shorter time interval. As compared to a "filter", which estimates the state of a system at time t using observations made up to time t, a "smoother" estimates the system state at time t using observations made before and after time t. Over a long time interval, one must generally take a more sophisticated

approach to smoothing than to simply consider linear combinations of an ensemble of trajectories generated over the entire interval, both because the trajectories may diverge enough that linear combinations of them will not approximate model trajectories, and because in the presence of model error there may be no model trajectory that fits the observations over the entire interval. Over a short enough time interval however, the approximation of true system trajectories by linear combinations of model trajectories with similar initial conditions is quite reasonable.

While this approach to assimilating asynchronous observations is suitable for any ensemble Kalman filter [16], it is particularly simple to implement in the LETKF framework; I will call this extension 4D-LETKF. To be specific, suppose that we have observations $(\tau_j, \mathbf{y}_{\tau_j}^o)$ taken at various times τ_j since the previous analysis. Let H_{τ_j} be the observation operator for time τ_j and let \mathbf{R}_{τ_j} be the error covariance matrix for these observations. In Section 2.3, we mapped a vector \mathbf{w} in the k-dimensional space \tilde{S} into observation space using the formula $\bar{\mathbf{y}}^b + \mathbf{Y}^b \mathbf{w}$, where the background observation operator H to the background ensemble at the analysis time. So now, for each time τ_j , we apply H_{τ_j} to the background ensemble at time τ_j , calling the mean of the resulting vectors $\bar{\mathbf{y}}_{\tau_j}^b$ and forming their differences from the mean into the matrix $\mathbf{Y}_{\tau_j}^b$.

We now form a combined observation vector \mathbf{y}^o by concatenating (vertically) the (column) vectors $\mathbf{y}_{\tau_j}^o$, and similarly by vertical concatenation of the vectors $\bar{\mathbf{y}}_{\tau_j}^b$ and matrices $\mathbf{Y}_{\tau_j}^b$ respectively, we form the combined background observation mean $\bar{\mathbf{y}}^b$ and perturbation matrix \mathbf{Y}^b . We form the corresponding error covariance matrix \mathbf{R} as a block diagonal matrix with blocks \mathbf{R}_{τ_j} (this assumes that observations taken at different times have uncorrelated errors, though such correlations if present could be included in \mathbf{R}).

Given this notation, we can then use the same analysis equations as in the previous sections, which are based on minimizing the cost function \tilde{J}^* given by (19). (We could instead write down the appropriate analogue to (15) and minimize the resulting nonlinear cost function \tilde{J} ; this would be no harder than in the case of synchronous observations.) Referring to Section 3, the only change is in Step 1, which one should perform for each observation time τ_j (using the background ensemble and observation operator for that time) and then concatenate the results as described above. Step 2 still only needs to be done at the analysis time, since its output is used only in Step 8 to form the analysis ensemble in model space. All of the intermediate steps work exactly the same, in terms of the output of Step 1.

In practice the model will be integrated with a discrete time step that in general will not coincide with the observation times τ_j . One should either interpolate the background ensemble trajectories to the observation times, or simply round the observation times off to the nearest model integration time. In either case, one must either store the background ensemble trajectories until the analysis time, or perform Step 1 of Section 3 during the model integration and store its results. The latter approach will require less storage if the number of observations per model integration time step is less than the number of model variables.

5 Summary, Results, and Acknowledgments

In this article I have described a general framework for data assimilation in spatiotemporally chaotic systems using an ensemble Kalman filter that in its basic version (Section 3) is relatively efficient and simple to implement. In a particular application, one may be able to improve accuracy by experimenting with different approaches to localization (see the discussion in Sections 2.2 and 2.3), covariance inflation (see the end of Section 2.3), and/or asynchronous observations (Section 4). For very large systems and/or when maximum efficiency is important, one should consider carefully the comments about implementation in Section 3 (and at the end of Section 4, if applicable). One can also apply this method to low-dimensional chaotic systems, without using localization.

Results using the LETKF approach will be reported elsewhere. The quality of these results is similar to other published results for square-root ensemble Kalman filters [38, 30, 34]. In particular, J. Harlim has obtained results [13] comparable to those in [38, 30] for a perfect model scenario, using a system with one spatial dimension proposed by E. Lorenz in 1995 [26, 27]. Also, E. Kostelich and I. Szunyogh have obtained preliminary results comparable to those in [34] using the LEKF approach of [29, 30] for the National Weather Service's global forecast model, again in a perfect model scenario, with the LETKF approach running several times faster. Thus, this article does not describe a fundamentally new method for data assimilation, but rather a refinement of existing approaches that combines simplicity with the flexibility to adapt to a variety of applications.

I thank my collaborators for their ideas and encouragement. In particular, E. Kostelich, I. Szunyogh, and G. Gyarmati have put considerable effort into a parallel LETKF implementation for global forecast models. J. Harlim has performed a 4D-LETKF implementation and done extensive testing on several models. E. Fertig has tested the approach described here for a nonlinear observation operator, and is working with H. Li and J. Liu on applying our parallel LETKF implementation to the NASA global forecast model. M. Cornick has done an LETKF implementation for Rayleigh-Bénard convection. Using the LEKF of [29, 30], T. Miyoshi has tested the smoothed localization described in Section 2.3 [28], and T. Sauer helped develop and test the 4D extension described here [16]. E. Kalnay, E. Ott, D. Patil, and J. Yorke have provided numerous insights that form the basis for this work. I am grateful to all for their feedback, which has helped to refine the approach described in this paper.

References

- J. L. Anderson, An ensemble adjustment Kalman filter for data assimilation, Mon. Wea. Rev. 129, 2884–2903 (2001).
- [2] J. L. Anderson & S. L. Anderson, A Monte Carlo implementation of the nonlinear filtering problem to produce ensemble assimilations and forecasts, Mon. Wea. Rev. 127, 2741–2758 (1999).
- [3] C. H. Bishop, B. J. Etherton, S. J. Majumdar, Adaptive sampling with the ensemble transform Kalman filter. Part I: Theoretical aspects, Mon. Wea. Rev. 129, 420–436 (2001).
- [4] G. Burgers, P. J. van Leeuwen, G. Evensen, Analysis scheme in the ensemble Kalman filter, Mon. Wea. Rev. 126, 1719–1724 (1998).
- [5] A. Doucet, N. de Freitas, N. Gordon, eds., *Sequential Monte Carlo Methods in Practice*, Springer-Verlag, 2001.
- [6] G. Evensen, Sequential data assimilation with a nonlinear quasi-geostrophic model using Monte Carlo methods to forecast error statistics, J. Geophys. Res. 99, 10143–10162 (1994).
- [7] G. Evensen & P. J. van Leeuwen, An ensemble Kalman smoother for nonlinear dynamics, Mon. Wea. Rev. 128, 1852–1867 (2000).
- [8] G. Evensen, The ensemble Kalman filter: theoretical formulation and practical implementation, Ocean Dynam. 53, 343–367 (2003).
- [9] I. Fukumori, A partitioned Kalman filter and smoother, Mon. Wea. Rev. 130, 1370–1383 (2002).
- [10] M. Ghil & P. Malanotte-Rizzoli, Data assimilation in meteorology and oceanography, Adv. Geophys. 33, 141–266 (1991).
- [11] T. M. Hamill & C. Snyder, A hybrid ensemble Kalman filter–3D variational analysis scheme Mon. Wea. Rev. 128 2905–2919 (2000).
- [12] T. M. Hamill, J. S. Whitaker, and C. Snyder, Distance-dependent filtering of background error covariance estimates in an ensemble Kalman filter, Mon. Wea. Rev. 129, 2776–2790 (2001).
- [13] J. Harlim & B. R. Hunt, A local ensemble transform Kalman filter: an efficient scheme for assimilating atmospheric data, preprint.
- [14] P. L. Houtekamer & H. L. Mitchell, Data assimilation using an ensemble Kalman filter technique, Mon. Wea. Rev. 126, 796–811 (1998).

- [15] P. L. Houtekamer & H. L. Mitchell, A sequential ensemble Kalman filter for atmospheric data assimilation, Mon. Wea. Rev. 129, 123–137 (2001).
- [16] B. R. Hunt, E. Kalnay, E. J. Kostelich, E. Ott, D. J. Patil, T. Sauer, I. Szunyogh, J. A. Yorke, A. V. Zimin, Four-dimensional ensemble Kalman filtering, Tellus A 56, 273–277 (2004).
- [17] K. Ide, P. Courtier, M. Ghil, A. C. Lorenc, Unified notation for data assimilation: operational, sequential, and variational, J. Meteo. Soc. Japan 75, 181–189 (1997).
- [18] A. H. Jazwinski, Stochastic Processes and Filtering Theory, Academic Press (1970).
- [19] R. E. Kalman, A new approach to linear filtering and prediction problems, Trans. ASME, Ser. D: J. Basic Eng. 82, 35–45 (1960).
- [20] R. E. Kalman & R. S. Bucy, New results in linear filtering and prediction theory, Trans. ASME, Ser. D: J. Basic Eng. 83, 95–108 (1961).
- [21] E. Kalnay, *Atmospheric Modeling*, *Data Assimilation and Predictability*, Cambridge Univ. Press (2002).
- [22] E. Kalnay & Z. Toth, Removing growing errors in the analysis, Proc. of the 10th Conf. on Numerical Weather Prediction, Amer. Meteo. Soc., Portland, Oregon (1994).
- [23] C. L. Keppenne, Data assimilation into a primitive-equation model with a parallel ensemble Kalman filter, Mon. Wea. Rev. 128, 1971–1981 (2000).
- [24] F.-X. Le Dimet & O. Talagrand, Variational algorithms for analysis and assimilation of meteorological observations: theoretical aspects, Tellus A 38, 97–110 (1986).
- [25] A. C. Lorenc, A global three-dimensional multivariate statistical interpolation scheme, Mon. Wea. Rev. 109, 701–721 (1981).
- [26] E. N. Lorenz, Predictability: a problem partly solved, in *Proc. of the ECMWF Seminar on Predictability*, vol. 1, Reading, UK (1996).
- [27] E. N. Lorenz & K. A. Emanuel, Optimal sites for supplementary weather observations: simulation with a small model, J. Atmos. Sci. 55, 399–414 (1998).
- [28] T. Miyoshi, Ensemble Kalman filter experiments with a primitive-equation global model, Ph.D. dissertation, University of Maryland (2005).

- [29] E. Ott, B. R. Hunt, I. Szunyogh, M. Corazza, E. Kalnay, D. J. Patil, J. A. Yorke, A. V. Zimin, E. J. Kostelich, Exploiting Local Low Dimensionality of the Atmospheric Dynamics for Efficient Ensemble Kalman Filtering, arXiv:physics/0203058v3 (2002).
- [30] E. Ott, B. R. Hunt, I. Szunyogh, A. V. Zimin, E. J. Kostelich, M. Corazza, E. Kalnay, D. J. Patil, J. A. Yorke, A local ensemble Kalman filter for atmospheric data assimilation, Tellus A 56, 415–428 (2004).
- [31] D. F. Parrish & J. C. Derber, The National Meteorological Center's spectral statisticalinterpolation analysis system, Mon. Wea. Rev. **120**, 1747–1763 (1992).
- [32] D. J. Patil, B. R. Hunt, E. Kalnay, J. A. Yorke, E. Ott, Local low dimensionality of atmospheric dynamics, Phys. Rev. Lett. 86, 5878–5881 (2001).
- [33] L. M. Pecora & T. L. Carroll, Synchronization in chaotic systems, Phys. Rev. Lett. 64, 821–824 (1990).
- [34] I. Szunyogh, E. J. Kostelich, G. Gyarmati, D. J. Patil, B. R. Hunt, E. Kalnay, E. Ott, J. A. Yorke, Assessing a local ensemble Kalman filter: perfect model experiments with the National Centers for Environmental Prediction global model, Tellus A 57, 528–545 (2005).
- [35] O. Talagrand & P. Courtier, Variational assimilation of meteorological observations with the adjoint vorticity equation I: theory, Quart. J. Roy. Meteo. Soc. 113, 1311–1328 (1987).
- [36] M. K. Tippett, J. L. Anderson, C. H. Bishop, T. M. Hamill, J. S. Whitaker, Ensemble squareroot filters, Mon. Wea. Rev. 131, 1485–1490 (2003).
- [37] X. Wang, C. H. Bishop, S. J. Julier, Which is better, an ensemble of positive-negative pairs or a centered spherical simplex ensemble?, Mon. Wea. Rev. 132, 1590–1605 (2004).
- [38] J. S. Whitaker, T. M. Hamill, Ensemble data assimilation without perturbed observations, Mon. Wea. Rev. 130, 1913–1924 (2002).
- [39] F. Zhang, C. Snyder, J. Sun, Impacts of initial estimate and observation availability on convective-scale data assimilation with an ensemble Kalman filter, Mon. Wea. Rev. 132, 1238– 1253 (2004).
- [40] M. Zupanski, Maximum likelihood ensemble filter: theoretical aspects, Mon. Wea. Rev. 133, 1710–1726 (2005).