

Introduction to Photolysis

AOSC / CHEM 433 & AOSC / CHEM 633

Ross Salawitch

Class Web Sites:

<http://www2.atmos.umd.edu/~rjs/class/fall2020>

<https://myelms.umd.edu/courses/1291919>

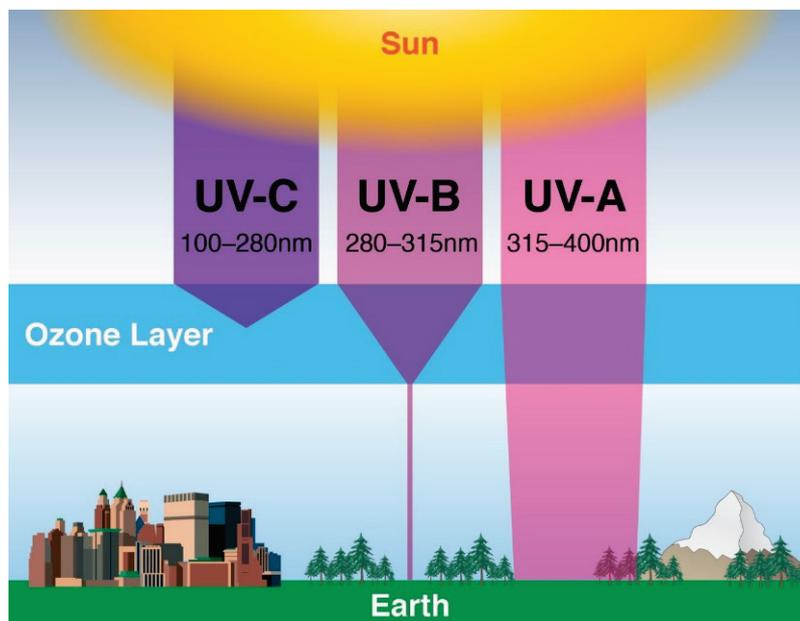


Figure 3.3.2, Wilmoth, Salawitch, and Canty, 2018

https://www2.atmos.umd.edu/~rjs/class/spr2020/readings/green_chemistry_chapter_3.3.pdf

Lecture 11 Catch-Up

21 October 2020

Biological Effects of UV Radiation

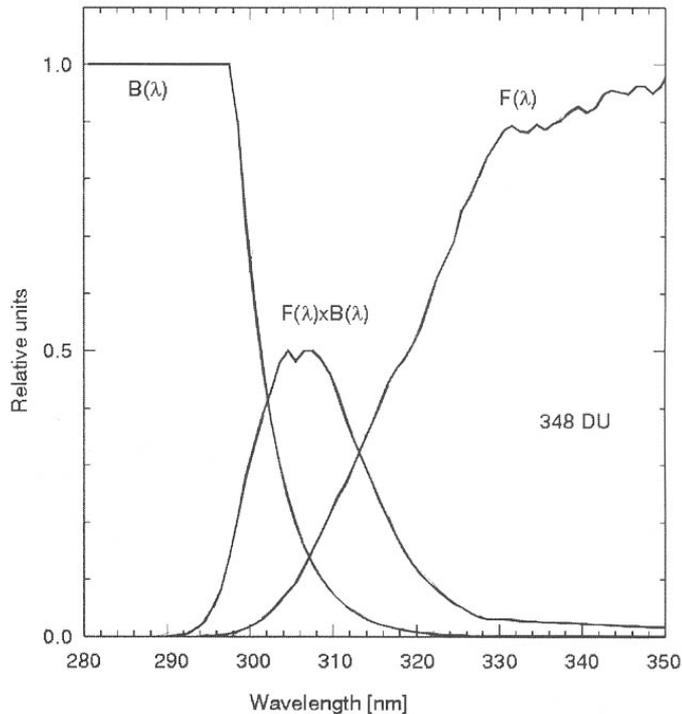


Fig. 1. Biologically active UV radiation. The overlap between the spectral irradiance $F(\lambda)$ and the erythemal action spectrum $B(\lambda)$ given by McKinlay and Diffey [6] shows the spectrum of biologically active radiation, $F(\lambda)B(\lambda)$. The area under the product function $F(\lambda)B(\lambda)$ is the biologically active dose rate. For a total ozone column of 348 DU.

Humans are:

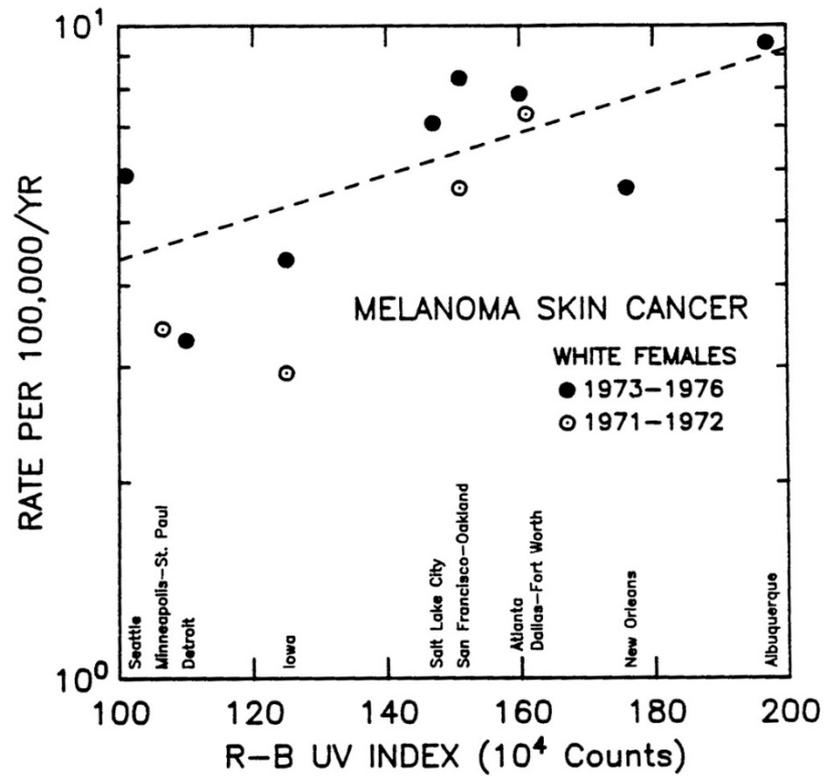
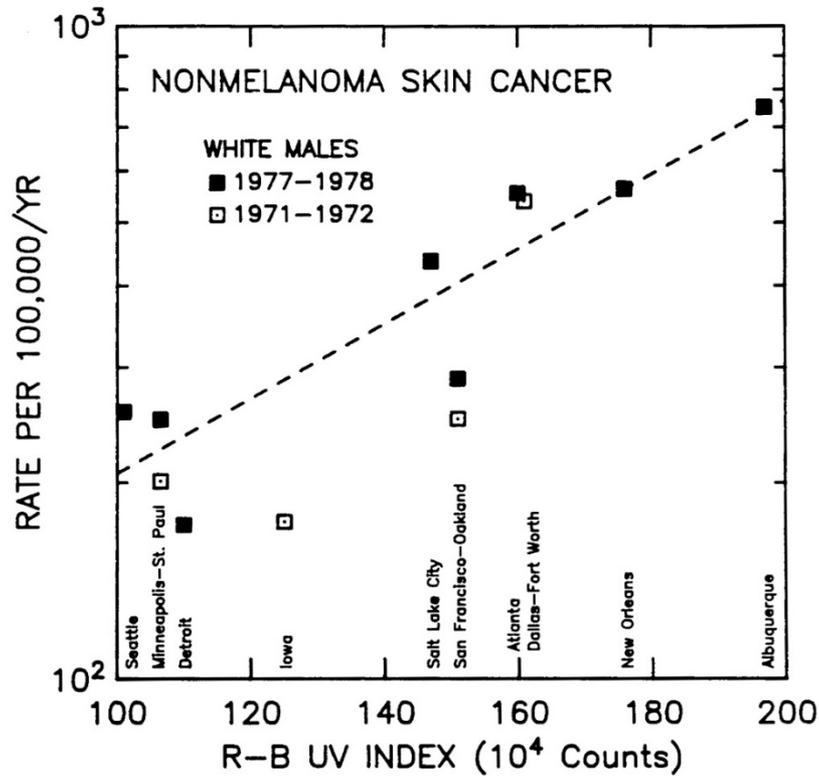
- strongly affected by exposure to UV-C radiation (100 to 280 nm)
- moderately affected by exposure to UV-B radiation (280 to 315 nm)
- weakly affected by exposure to UV-A radiation (315 to 400 nm)

http://www.who.int/uv/uv_and_health/en

⇐ From Mandronich et al., *J. Photochemistry and Photobiology*, vol. 46, pg. 5, 1998

The “biologically active dose rate” maximizes in the UV-B region at ~305 nm

Relationship Between Cancer and UV



Scotto and Fraumeni, *Cancer Epidemiology*, W. B. Saunders and Co, Philadelphia, 1982.

Factor of 2 rise in UV Index leads to factor of 4 rise in Non-Melanoma Skin Cancer:

i.e., Non-Melanoma Skin Cancer rises about twice as fast as incident solar ultraviolet (UV) radiation

Factor of 2 rise in UV Index leads to factor of 2 rise in Melanoma Skin Cancer:

i.e., Melanoma Skin Cancer rises at about the same rate as incident solar ultraviolet (UV) radiation

Relationship Between UV and Column Ozone

50% drop in O₃ leads to 100% rise in UV
i.e., UV rises about twice as fast as
O₃ declines

Therefore, Melanoma Skin Cancer
should rise by 2 % for every
1 % drop in O₃

and Non-Melanoma Skin Cancer
should rise by 4 % for every
1 % drop in O₃

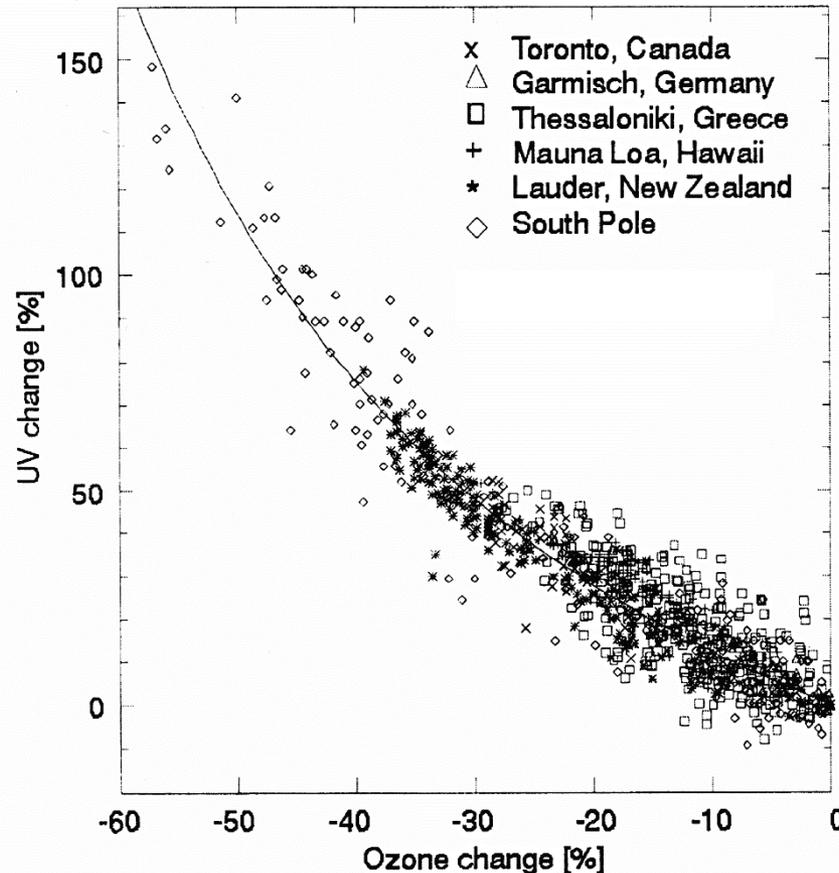
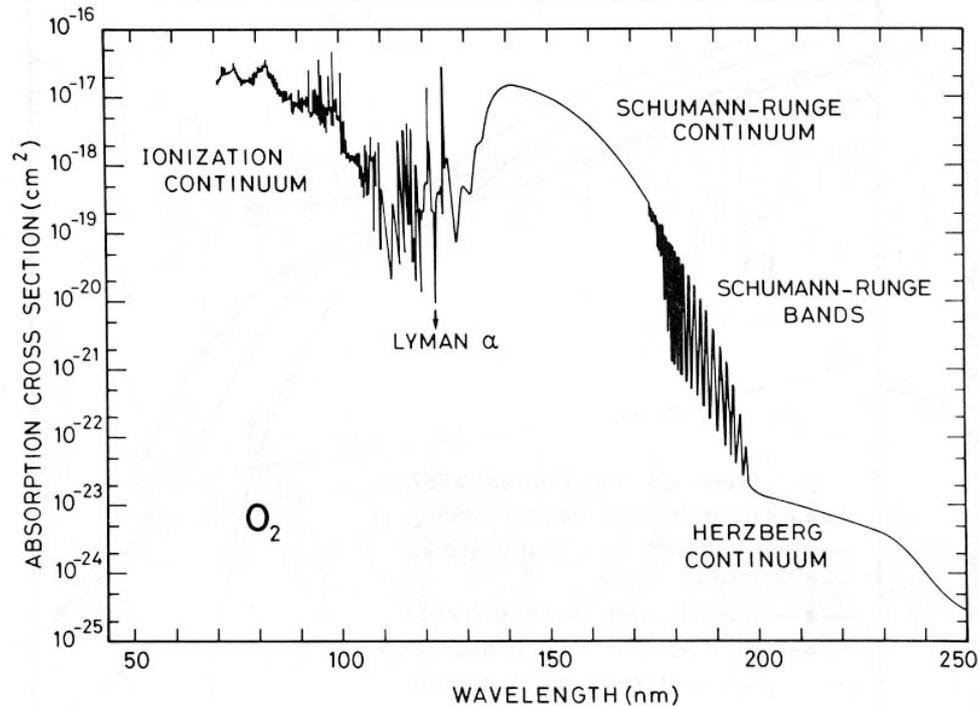


Fig. 2. Dependence of erythemal ultraviolet (UV) radiation at the Earth's surface on atmospheric ozone, measured on cloud-free days at various locations, at fixed solar zenith angles. Legend: South Pole [8]; Mauna Loa, Hawaii [9]; Lauder, New Zealand [10]; Thessaloniki, Greece (updated from Ref. [11]); Garmisch, Germany [12]; and Toronto, Canada (updated from Ref. [13]).

Madronich et al., *J. of Photochemistry and Photobiology B*, Vol. 46, 5–19, 1998.

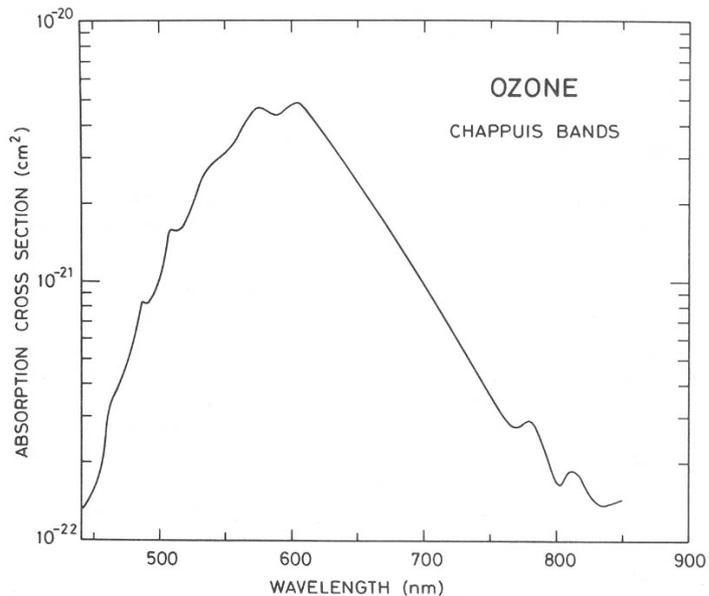
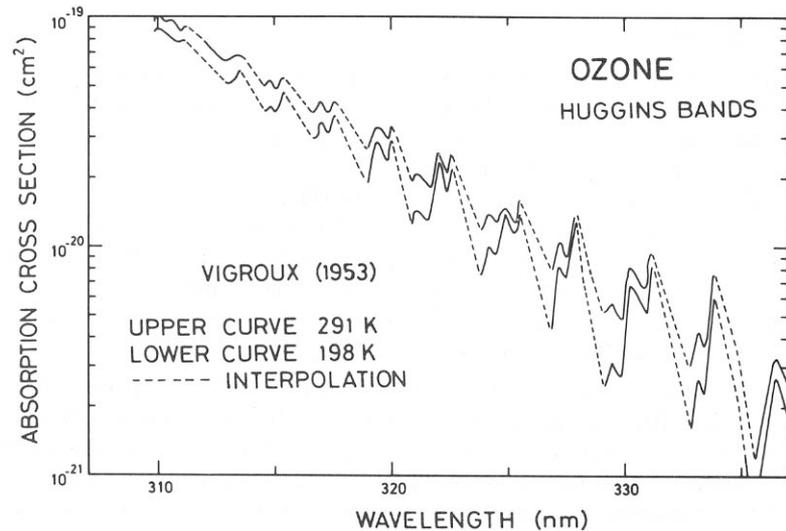
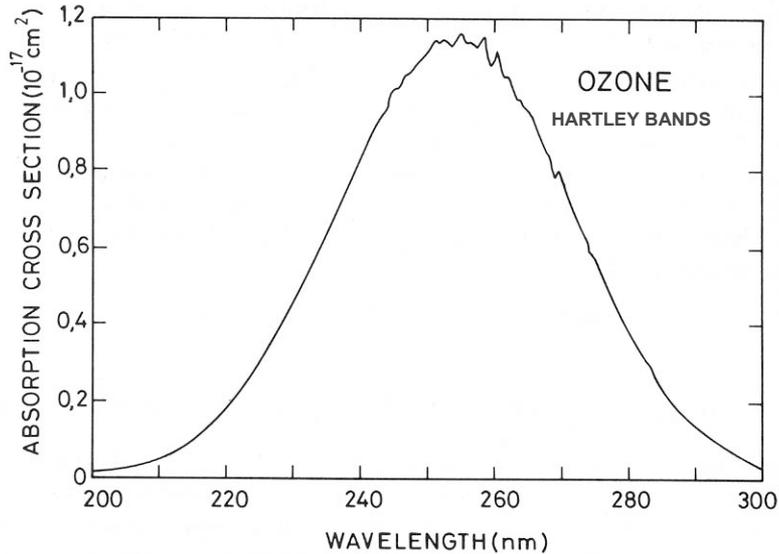
Absorption Cross Section of O₂



From Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986

- O₂ can not dissociate longward of ~250 nm
- All absorption shown above is dissociative (e.g., leads to production of two O atoms)
- Structure in the O₂ cross section is related to whether the initial transition involves an unbound electronic state (smooth) or involves a specific vibrational level of an electronic state (banded, due to requirement of specific quanta of energy)

Absorption Cross Section of O₃



Quantum Mechanically Allowed Transitions

Table 4.6 Theoretical limits corresponding to different photolysis products (nm).

| | $\text{O}_2(^3\Sigma_g^-)$ | $\text{O}_2(^1\Delta_g)$ | $\text{O}_2(^1\Sigma_g^+)$ | $\text{O}_2(^3\Sigma_u^-)$ | $\text{O}_2(^3\Sigma_u^+)$ |
|------------------------|----------------------------|--------------------------|----------------------------|----------------------------|----------------------------|
| $\text{O}(^3\text{P})$ | 1180 | 590 | 460 | 230 | 170 |
| $\text{O}(^1\text{D})$ | 410 | 310 | 260 | 167 | 150 |
| $\text{O}(^1\text{S})$ | 234 | 196 | 179 | 129 | 108 |

From Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986

Solar Spectral Actinic Flux

130 ATMOSPHERIC PHOTOCHEMISTRY AND CHEMICAL KINETICS

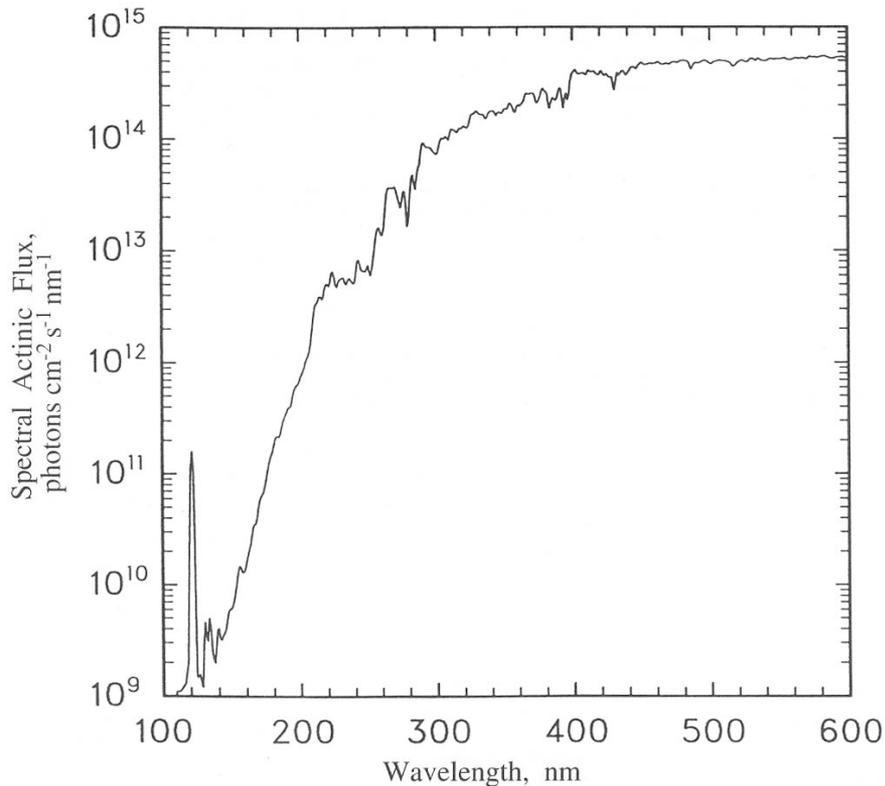


FIGURE 6. Solar spectral actinic flux (photons $\text{cm}^{-2} \text{s}^{-1} \text{nm}^{-1}$) at the top of Earth's atmosphere.

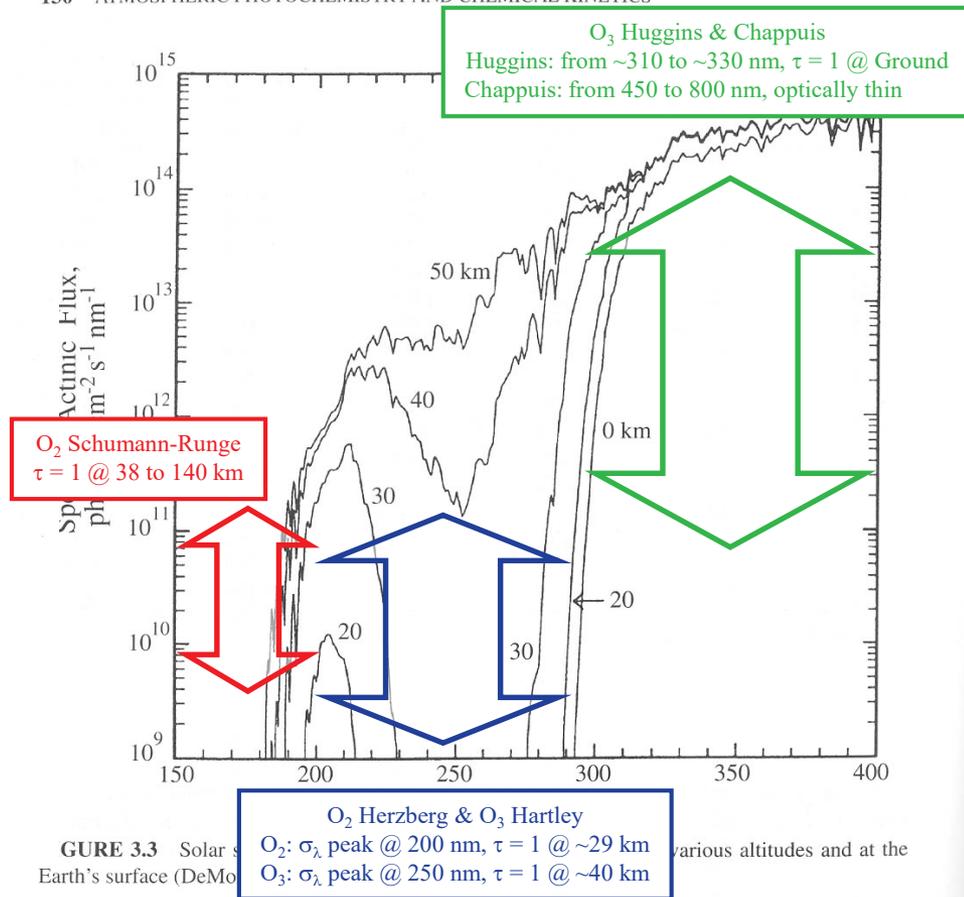


FIGURE 3.3 Solar spectral actinic flux at various altitudes and at the Earth's surface (DeMore et al., 1994).

From DeMore et al., *Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling*, Evaluation No. 11, 1994.

From Seinfeld and Pandis, *Atmospheric Chemistry and Physics*, 1998.

Photodissociation Frequencies

Next goal is to understand:

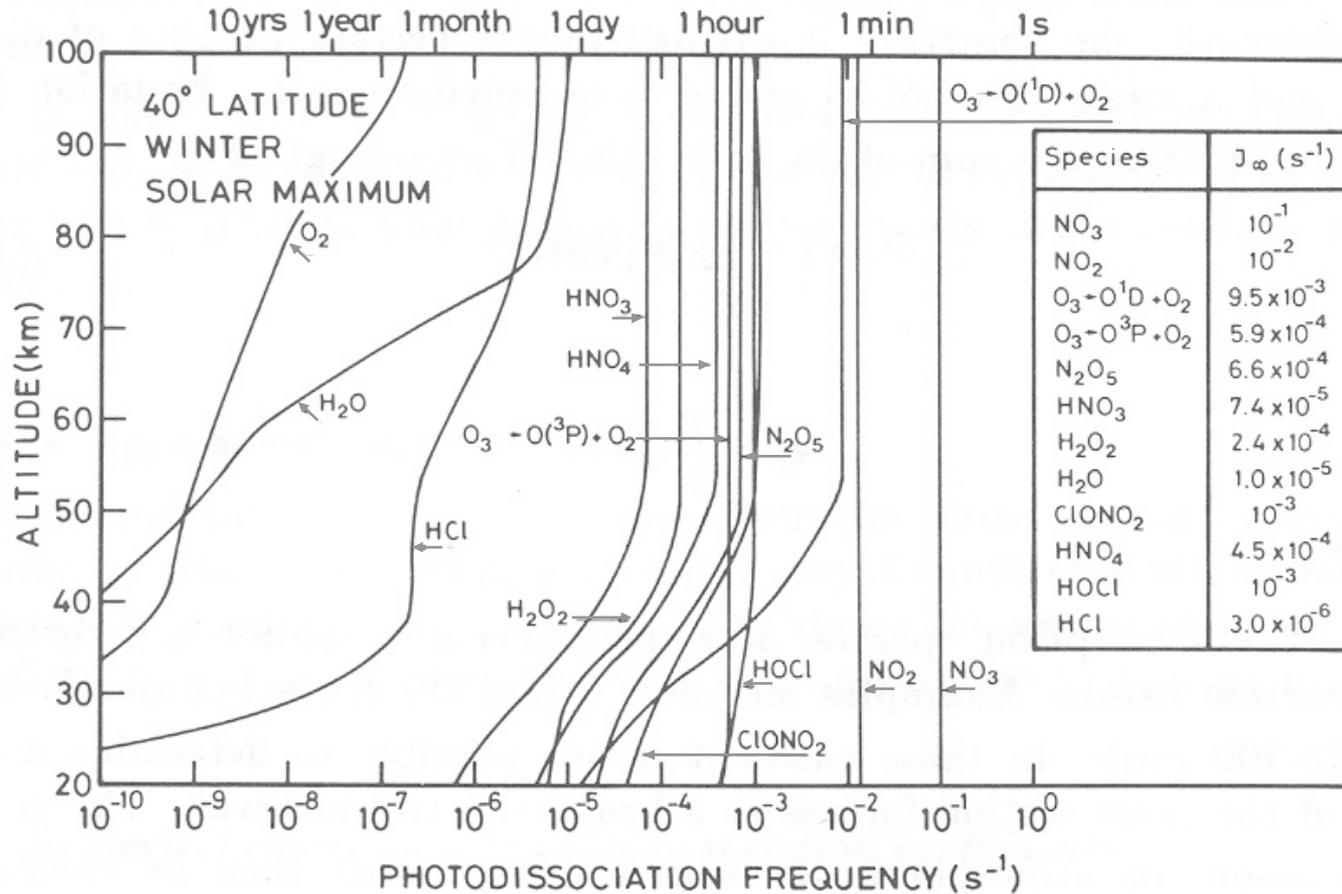


FIGURE 4.58 Photodissociation frequencies for numerous important atmospheric species.

From Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986

Photodissociation Frequencies

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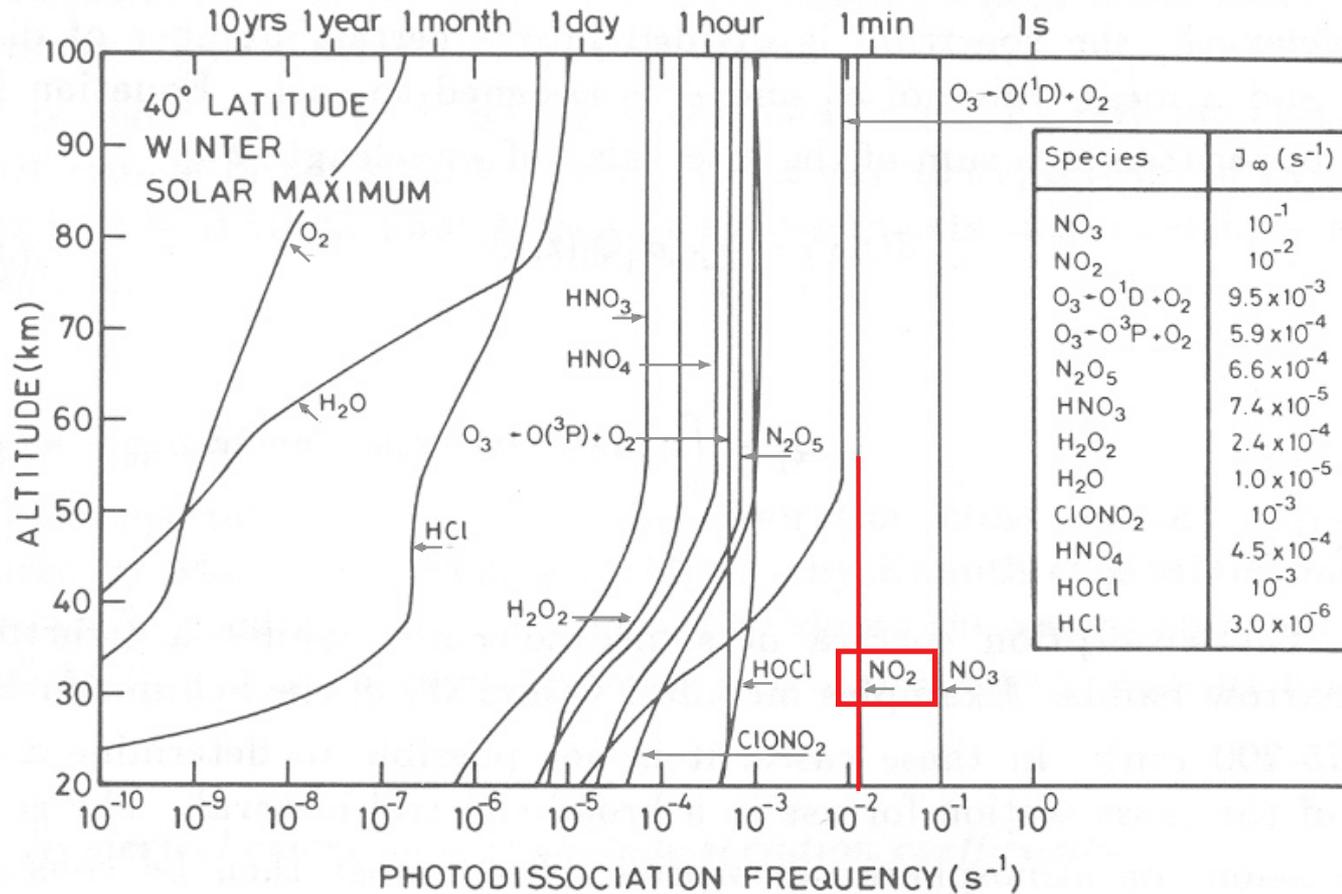


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From Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986

Photodissociation Frequencies

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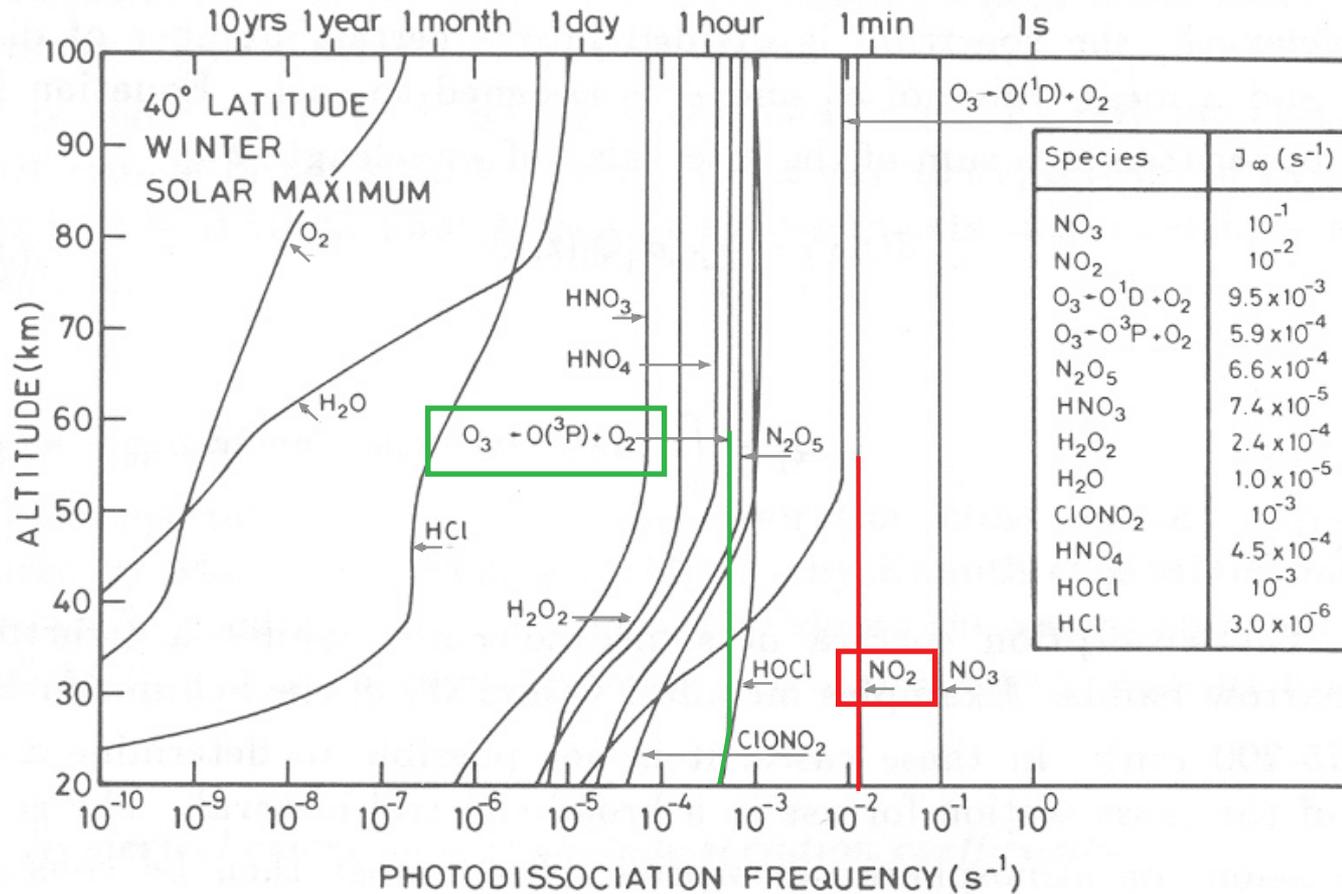


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Photodissociation Frequencies

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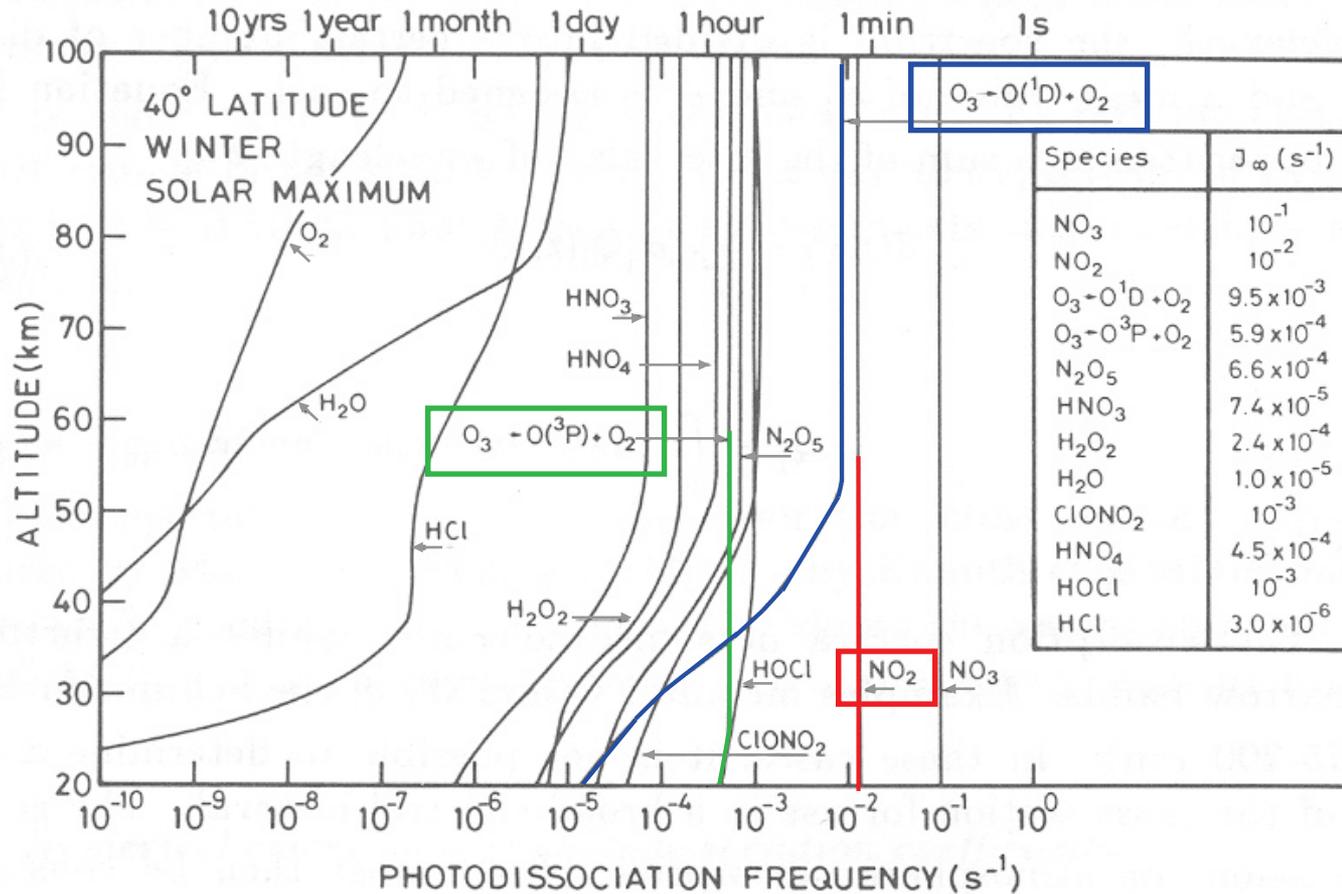
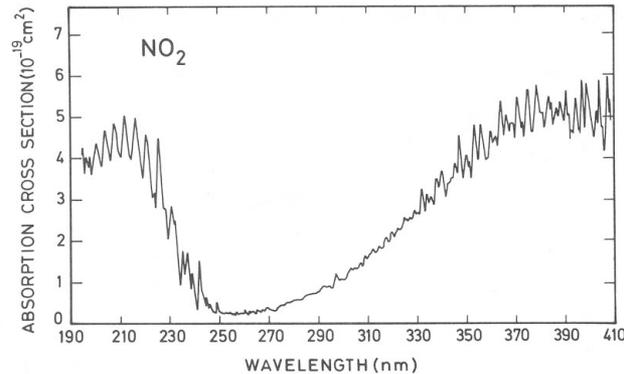


FIGURE 4.58 Photodissociation frequencies for numerous important atmospheric species.

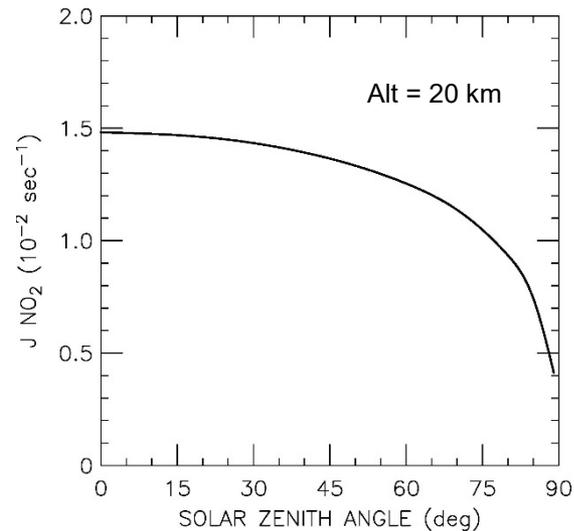
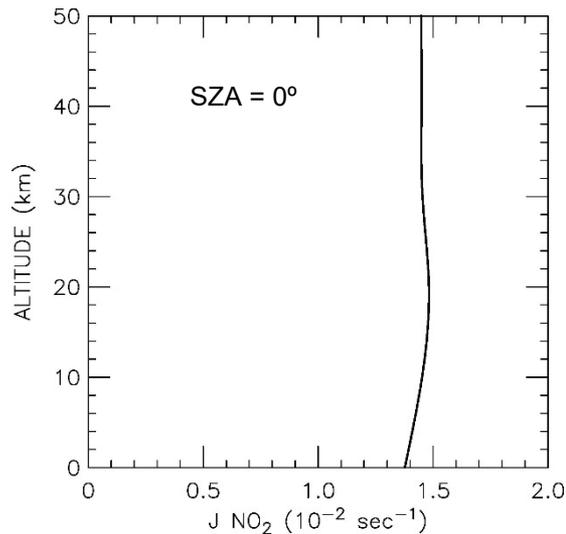
From Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986

NO₂ Photolysis

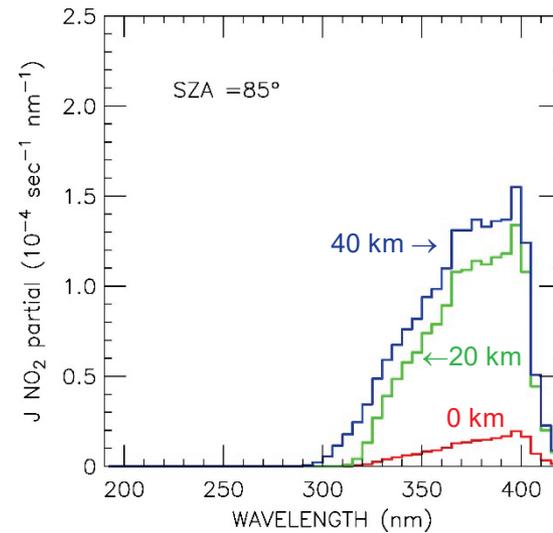
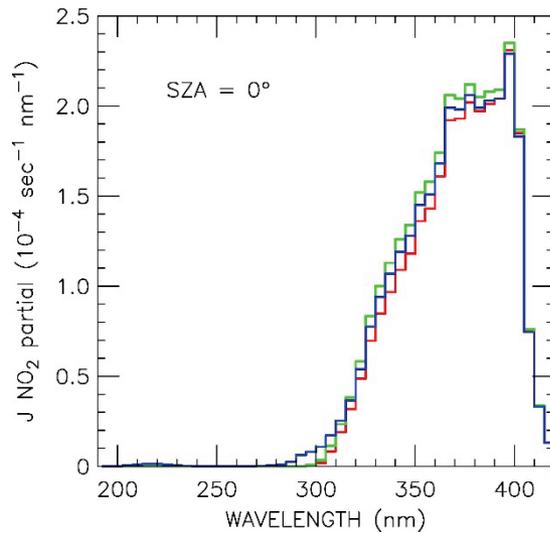
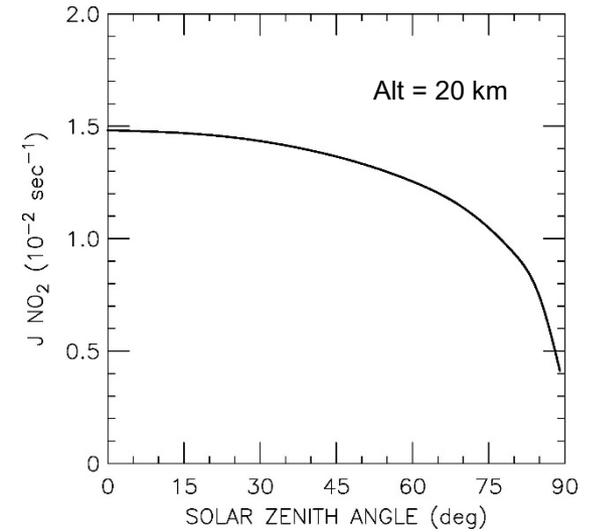
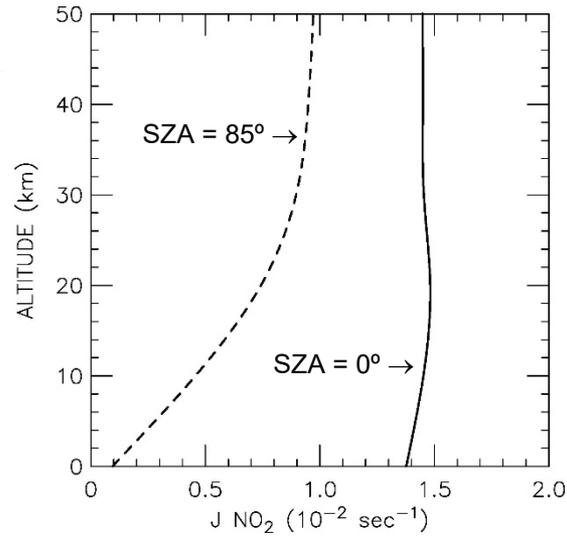
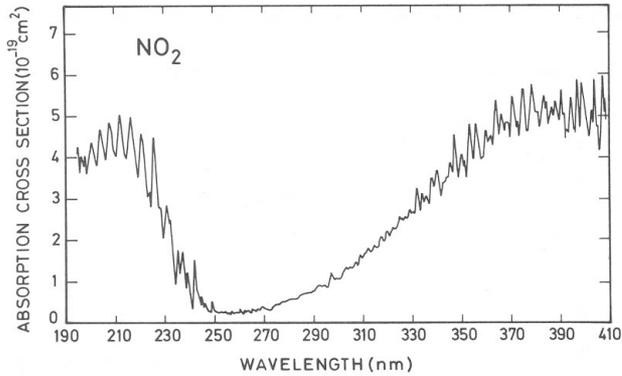
The majority of NO₂ photolysis occurs longward of 300 nm, where the atmosphere is optically thin with respect to absorption by O₃ and O₂:



leading to a value for J_{NO₂} that is nearly independent of height and SZA:

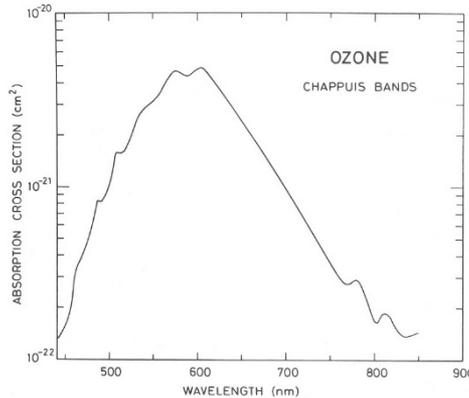


NO₂ Photolysis

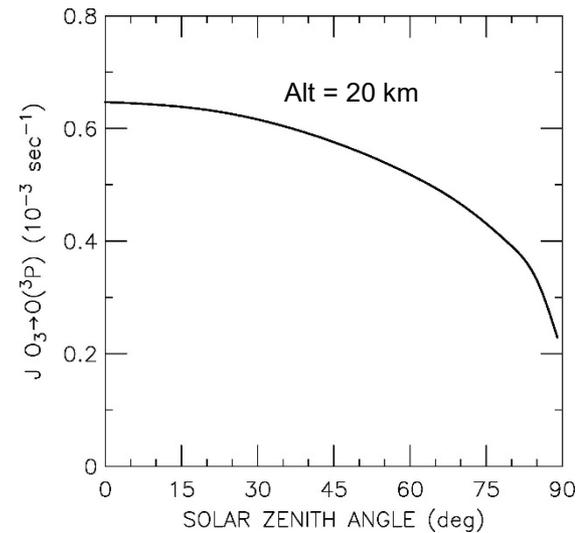
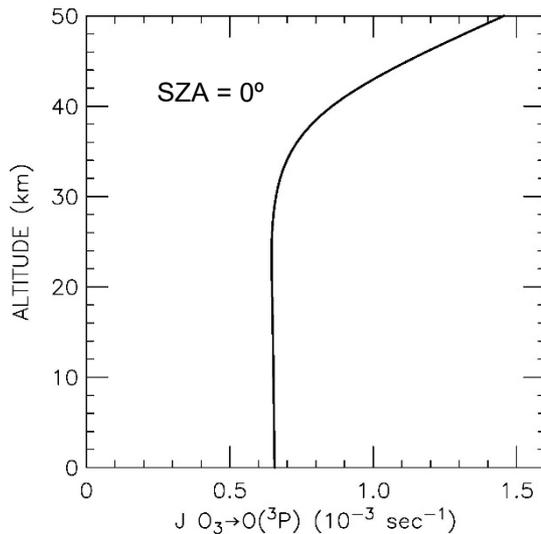


$O_3 \rightarrow O(^3P)$ Photolysis

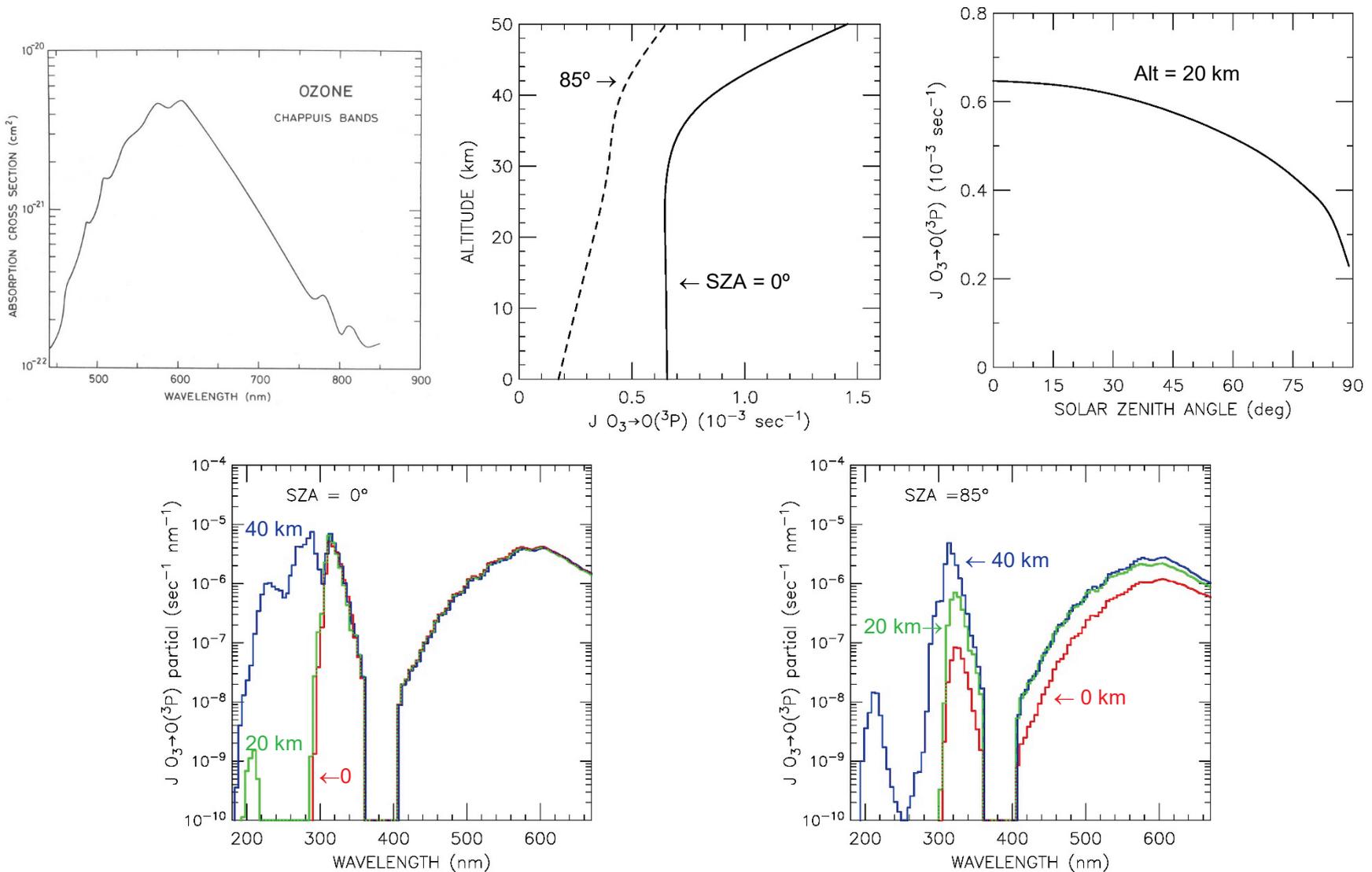
The production of $O(^3P)$ from photolysis of O_3 occurs mainly longward of 500 nm, where the atmosphere is optically thin with respect to absorption by O_3 :



leading to a value for $J_{O_3 \rightarrow O(^3P)}$ that is essentially independent of height and SZA:

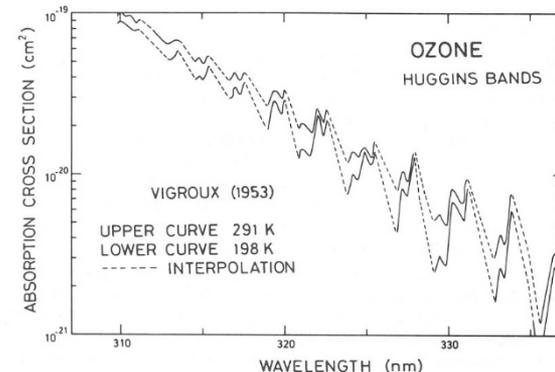
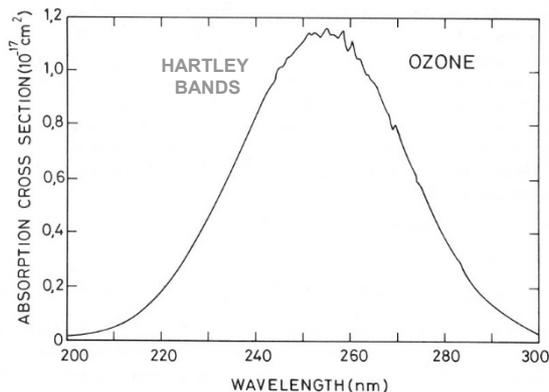
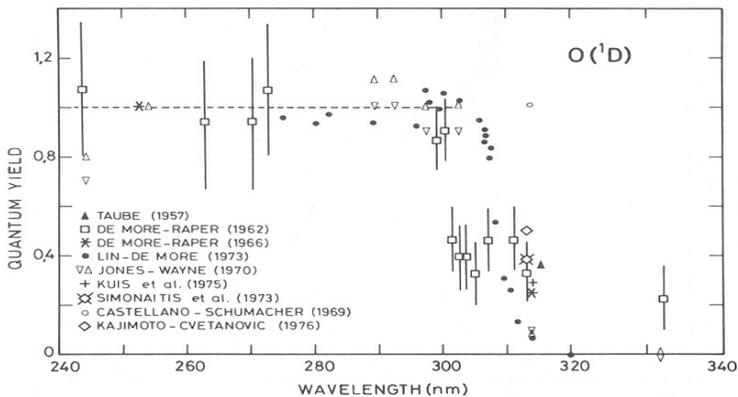


O₃ → O(³P) Photolysis

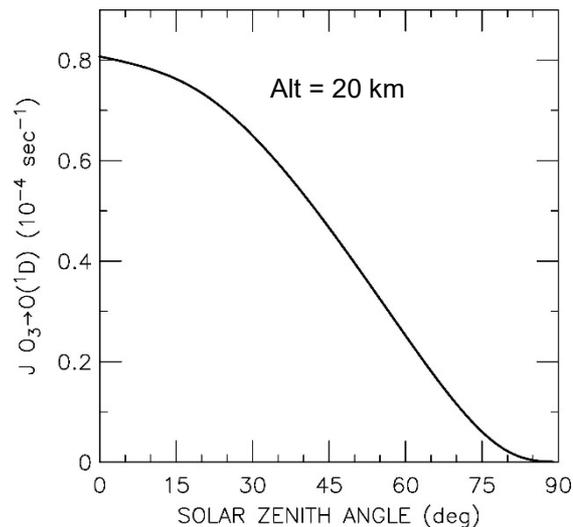
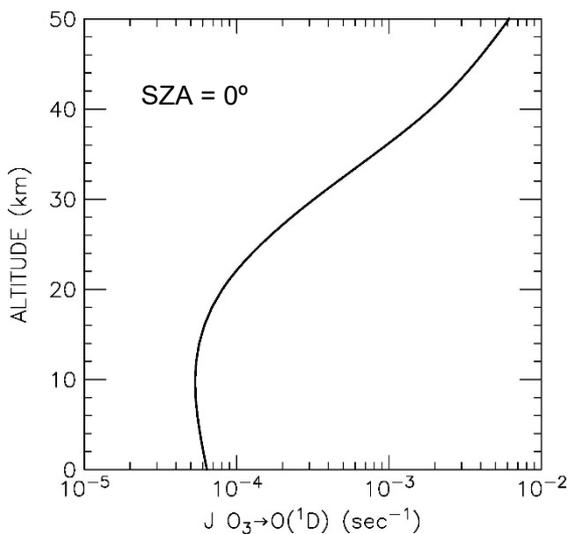


O₃ → O(¹D) Photolysis

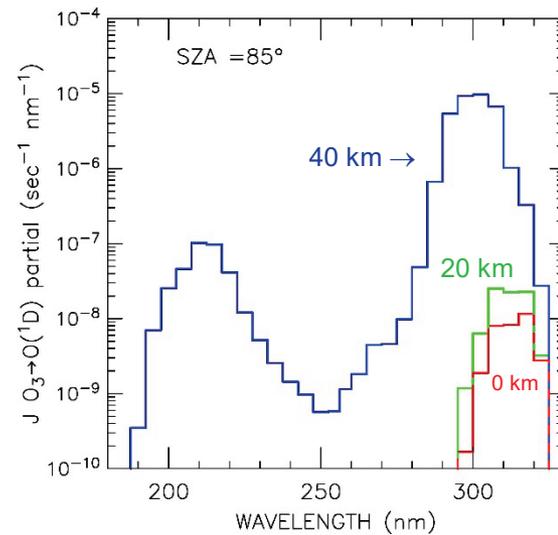
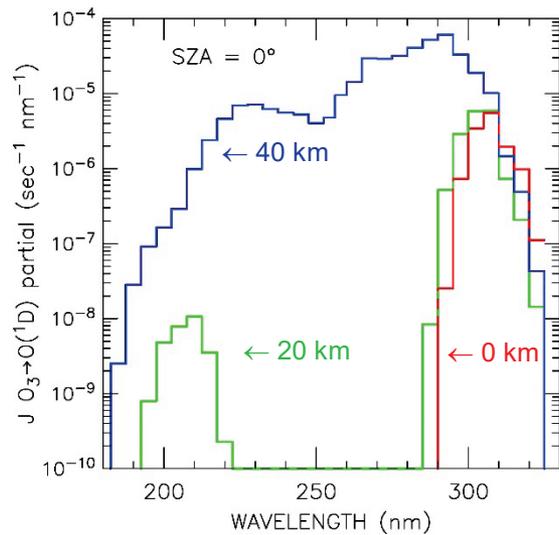
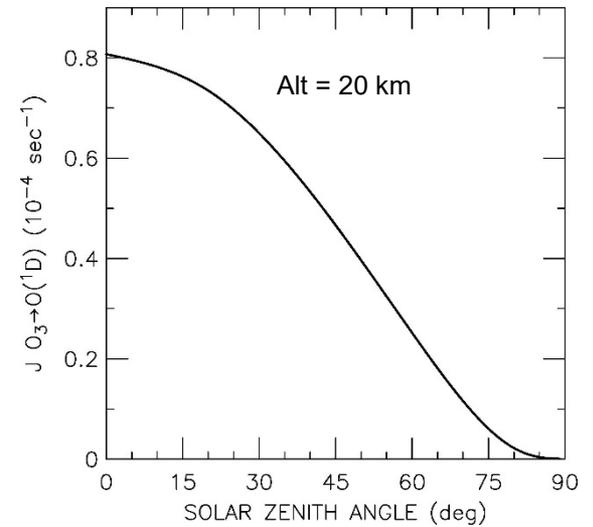
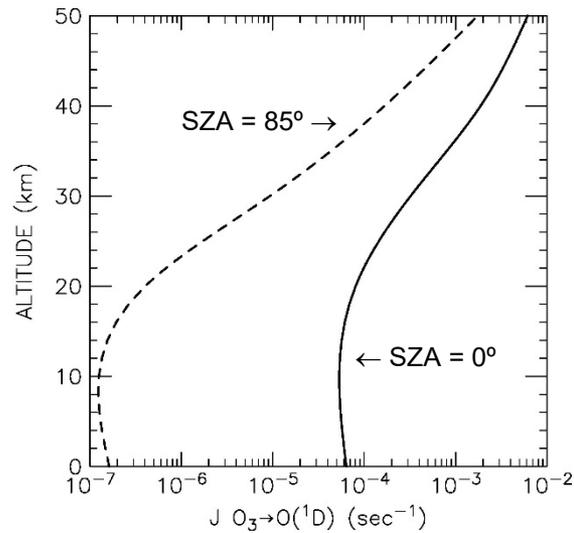
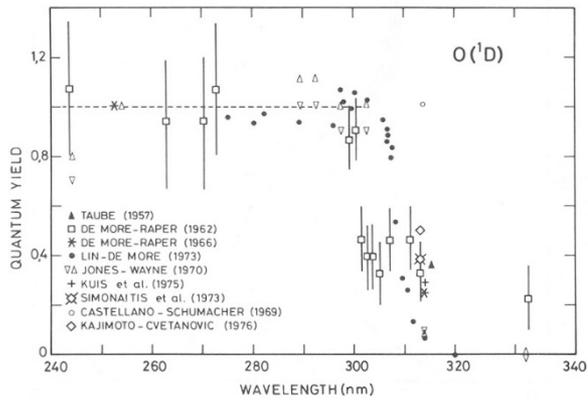
The production of O(¹D) from photolysis of O₃ occurs shortward of 320 nm, where the atmosphere is basically optically thick with respect to absorption by O₃:



leading to a value for $J_{O_3 \rightarrow O(1D)}$ that is dependent on height and SZA:



O₃ → O(¹D) Photolysis



Extra #1: Height and Abundance of Ozone

Chapman expression for $[O_3]$:

$$[O_3] = \left[\frac{f_{O_2} k_2}{J_3 k_4} \right]^{\frac{1}{2}} \left[J_1[O_2] \right]^{\frac{1}{2}} [M]$$

The concentration of O_3 should peak at the altitude where the product of the square-root of the O_2 photolysis rate times the density of air is largest

$\left[J_1[O_2] \right]^{\frac{1}{2}}$ peaks at same altitude as $J_1[O_2]$: ~35 km

$\left[J_1[O_2] \right]^{\frac{1}{2}} [M]$ peaks about a scale height lower: ~28 km

Extra #1: Height and Abundance of Ozone

The total *photolysis rate* is the product of the concentration of a gas and the total photolysis frequency (*J value*):

$$\text{Photolysis Rate Gas (z)} = [\text{Gas}] \times J_{\text{gas}} \quad \text{Units: molecules cm}^{-3} \text{ sec}^{-1}$$

$$= [\text{Gas}] \int_{\lambda_{\min}}^{\lambda_{\max}} J_{\text{gas}}(z, \lambda) d\lambda$$

$$\approx [\text{Gas}] \int_{\lambda_{\min}}^{\lambda_{\max}} \sigma_{\text{gas}}(z, T) F_{\text{TOA}}(\lambda) e^{-\tau(z, \lambda)} d\lambda$$

where:

$$\tau(z, \lambda) = m \int_z^{\infty} \sigma_{\text{O}_2}(\lambda, T) [\text{O}_2(z')] dz' + m \int_z^{\infty} \sigma_{\text{O}_3}(\lambda, T) [\text{O}_3(z')] dz'$$

Extra #1: Height and Abundance of Ozone

The total *photolysis rate* is the product of the concentration of a gas and the total photolysis frequency (*J value*):

$$\text{Photolysis Rate } O_2(z) \approx [O_2(z)] \int_{\lambda_{\min}}^{\lambda_{\max}} \sigma_{O_2}(z, T) F_{\text{TOA}}(\lambda) e^{-\tau(z, \lambda)} d\lambda$$

Assume:

1. O_2 is the only absorber:

$$\tau(z, \lambda) = m \int_z^{\infty} \sigma_{O_2}(\lambda, T) [O_2] dz'$$

2. σ_{O_2} is independent of T:

$$\tau(z, \lambda) = m \sigma_{O_2}(\lambda, T) \int_z^{\infty} [O_2] dz'$$

3. $[O_2]$ falls off exponentially with increasing height:

$$\tau(z, \lambda) = m \sigma_{O_2}(\lambda, T) [O_2]_{\text{ground}} H e^{-z/H}$$

$$\text{Photolysis Rate } O_2(z, \lambda) = J_{O_2} [O_2]$$

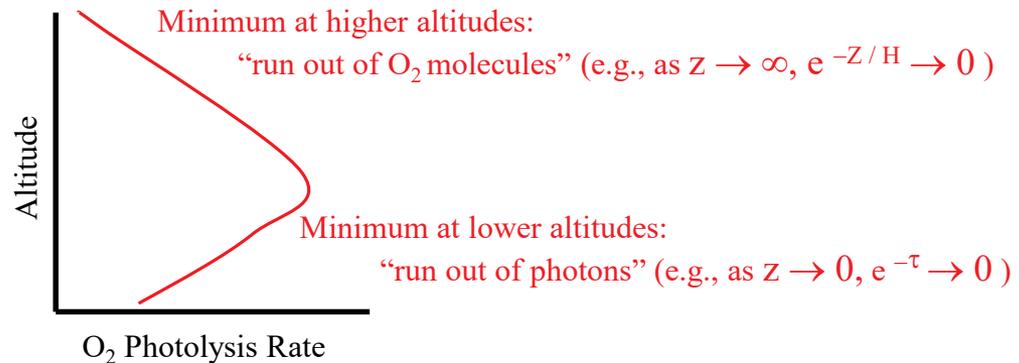
$$[O_2]_{\text{ground}} e^{-z/H} \sigma_{O_2} F_{\text{TOA}} \exp \{-m \sigma_{O_2} [O_2]_{\text{ground}} H e^{-z/H}\}$$

Extra #1: Height and Abundance of Ozone

$$\text{Photolysis Rate } O_2(z, \lambda) = J_{O_2} [O_2] = [O_2]_{\text{ground}} e^{-z/H} \sigma_{O_2} F_{\text{TOA}} \exp \{ -m \sigma_{O_2} [O_2]_{\text{ground}} H e^{-z/H} \}$$

What does this function look like?

Informally-



Formally-

Can show:

$$\frac{d [\text{Photolysis Rate } O_2 (z, \lambda)]}{dz} = 0$$

$$\text{if } m \sigma_{O_2} [O_2]_{\text{ground}} H e^{-z/H} = 1$$

Extra #1: Height and Abundance of Ozone

The *partial photolysis rate* of O₂ maximizes at the altitude where $\tau = 1$

This is true for any gas that is the primary absorber

At what altitude does $\tau = 1$?

$$z \approx H \log [\sigma_{\lambda} n 4 \times 10^{24} (\text{molecules/cm}^2)]$$

What is the value of the partial photolysis rate of O₂ when $\tau = 1$?

$$\text{Photolysis Rate of O}_2 \approx 5 \times 10^{18} (\text{molecules/cm}^3) e^{-z/H} \sigma_{\text{O}_2} F_{\text{TOA}} \frac{1}{e}$$

Let's examine the partial photolysis rate of O₂ in its three absorption regions

Assume $H = 7$ km (realistic for 240 K) & $\theta = 45^\circ$

| | σ_{max} (cm ²) | z ($\tau=1$) (km) | F_{TOA} (#/cm ² /s) | J (sec ⁻¹) | Photolysis Rate (#/cm ³ /sec) |
|--------------------------|---|--------------------------|--|-----------------------------|---|
| Schumann-Runge Continuum | 10^{-17} | 125 | 1×10^{11} | 3.7×10^{-7} | 3.3×10^4 |
| Schumann-Runge Bands | 10^{-20} | 77 | 8×10^{11} | 2.9×10^{-9} | 2.4×10^5 |
| | 3×10^{-23} | 36 | 3×10^{12} | 3.3×10^{-11} | 9.6×10^5 |
| Herzberg Continuum | 10^{-23} | 29 | 2×10^{14} | 7.3×10^{-10} | 5.8×10^7 |

Extra #1: Height and Abundance of Ozone

The height of the ozone layer (~30 km) and the thickness of the ozone layer (~400 DU) are determined by values of solar actinic flux (F_{TOA}) & O_2 absorption cross section (σ_{max})

Altitude of largest O_2 photolysis rate

Value of O_2 photolysis rate, $J_1[\text{O}_2]$ in Chapman expression at peak altitude

| | σ_{max} (cm^2) | z ($\tau=1$) (km) | F_{TOA} ($\#/\text{cm}^2/\text{s}$) | J (sec^{-1}) | Photolysis Rate ($\#/\text{cm}^3/\text{sec}$) |
|--------------------------|--|--------------------------|---|------------------------------|--|
| Schumann-Runge Continuum | 10^{-17} | 125 | 1×10^{11} | 3.7×10^{-7} | 3.3×10^4 |
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Extra #1: Height and Abundance of Ozone

The height of the ozone layer (~30 km) and the thickness of the ozone layer (~400 DU) are determined by values of solar actinic flux (F_{TOA}) & O_2 absorption cross section (σ_{max})

Suppose the Herzberg Continuum region dominated the photolysis rate of O_2 , but the maximum cross section was different:

$$\sigma_{\text{max}} = 3 \times 10^{-22} \text{ cm}^2 \quad \rightarrow \quad Z_{\text{OZONE LAYER}} = 48 \text{ km}$$

$$\sigma_{\text{max}} = 3 \times 10^{-25} \text{ cm}^2 \quad \rightarrow \quad Z_{\text{OZONE LAYER}} = 2 \text{ km}$$

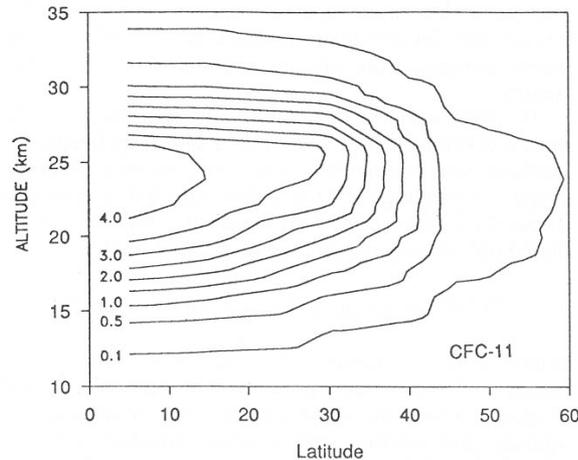
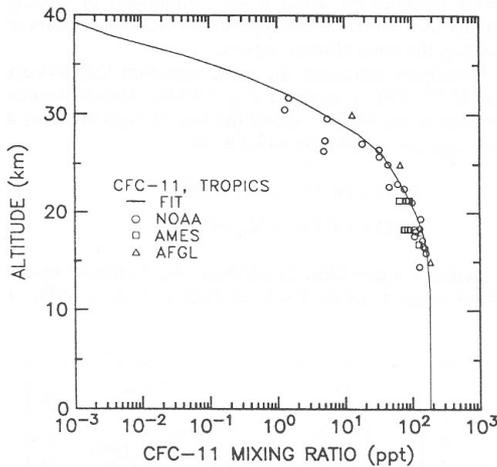
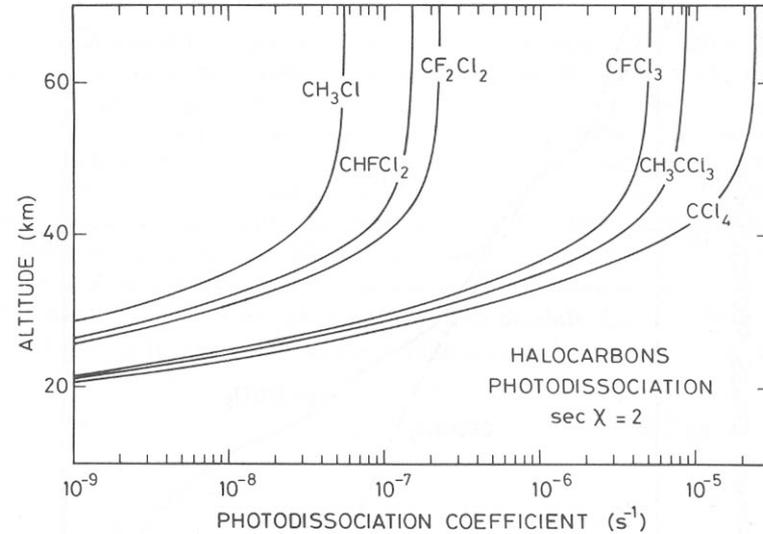
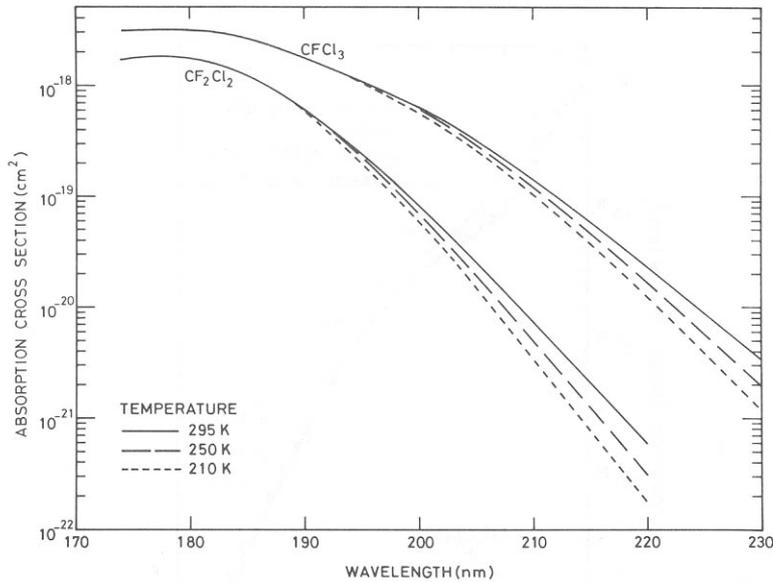
Suppose the Herzberg Continuum region dominated the photolysis rate of O_2 , but the solar irradiance at these wavelengths was different:

$$F_{\text{TOA}} = 2 \times 10^{15} \text{ \#/cm}^2/\text{s} \quad \rightarrow \quad \text{Ozone Column} \approx 900 \text{ DU}$$

$$F_{\text{TOA}} = 2 \times 10^{13} \text{ \#/cm}^2/\text{s} \quad \rightarrow \quad \text{Ozone Column} \approx 100 \text{ DU}$$

Extra #2: CFC Photolysis

Brasseur & Solomon, *Aeronomy of the Middle Atmosphere*, 1986



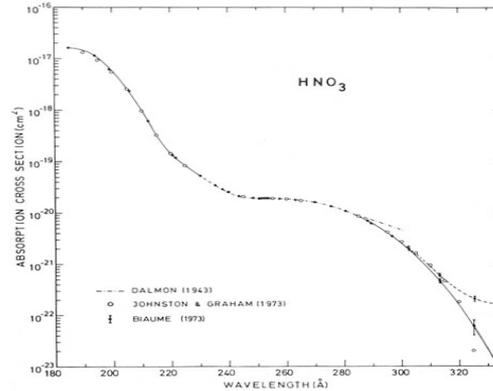
| | Inventory (kg) | Loss Rate (kg/yr) | Lifetime (years) |
|---------------------------------|-----------------------|-----------------------|---------------------|
| CFC ₁₃ | 3.1 × 10 ⁹ | 7.2 × 10 ⁷ | 44 |
| CF ₂ Cl ₂ | 3.5 × 10 ⁹ | 3.0 × 10 ⁷ | 116 |

Minschwaner *et al.*, *JGR*, 98, 10543, 1993

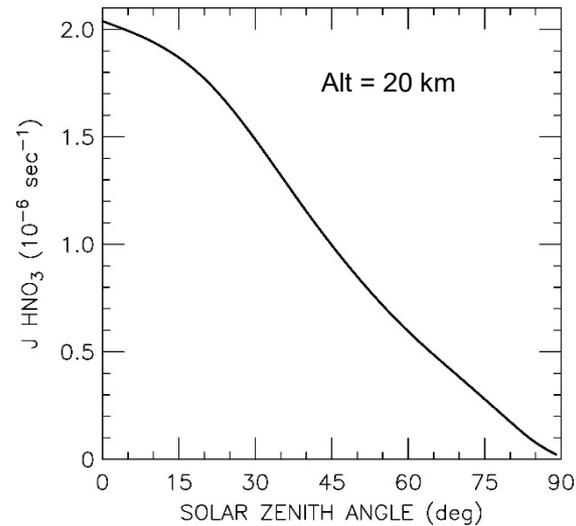
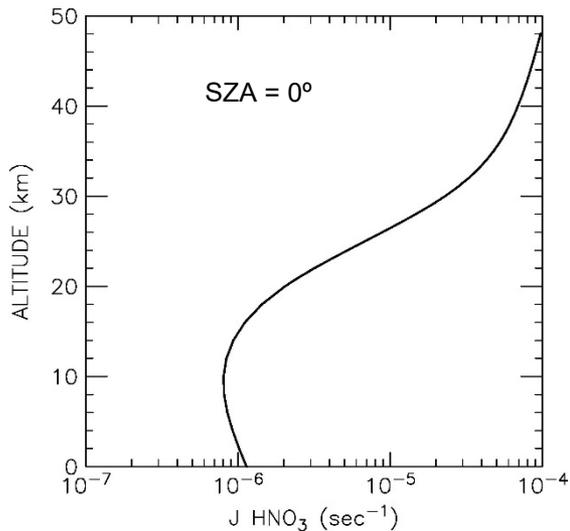
Fig. 12. Diurnally averaged loss rate for CFC₁₃ (molecules cm⁻³ s⁻¹) as a function of altitude and latitude, calculated with the line-by-line model, for equinox. The loss rate was calculated assuming destruction of CFC₁₃ by photolysis only.

Extra #3: HNO₃ Photolysis

The majority of HNO₃ photolysis occurs shortward of 320 nm, where the atmosphere is optically thick with respect to absorption by O₃ and O₂:



leading to a value for J_{HNO_3} that is strongly dependent on height and SZA:



Extra #3: HNO₃ Photolysis

