

# Pollution of Earth's Stratosphere: Ozone Recovery and Chemistry/Climate Coupling

AOSC / CHEM 433 & AOSC / CHEM 633

Ross Salawitch

Class Web Sites:

<http://www2.atmos.umd.edu/~rjs/class/fall2020>  
<https://myelms.umd.edu/courses/1291919>

## Motivating questions:

- a) How might climate change (future variations in temperature *and* / or circulation) driven by rising GHGs affect stratospheric ozone?
- b) Might climate at the surface be affected by stratospheric ozone?

Lecture 17  
10 November 2020

# Announcements: Class

- Problem Set #3:

[https://www2.atmos.umd.edu/~rjs/class/fall2020/problem\\_sets/ACC\\_2020\\_problem\\_set\\_03.pdf](https://www2.atmos.umd.edu/~rjs/class/fall2020/problem_sets/ACC_2020_problem_set_03.pdf)

and is due on ~~Tuesday, 10 November~~ Thursday, 12 November.

- Due to “popular demand”, this problem set will be completed outside of ELMS: prefer you mail me and Laura McBride ([mcbridel@terpmail.umd.edu](mailto:mcbridel@terpmail.umd.edu)) either one PDF file (entire P Set) or two PDF files (one per question) when the problem set is complete, with an email subject such as:

AOSC 433: Problem Set 3 \*or\* CHEM 633: Problem Set 3 , etc

- Please email me and Laura with questions

*Problem Set #3 due date extended to Thursday, 12 Nov due to issues with the class website on 5 Nov 2020.*

# Announcements: Class

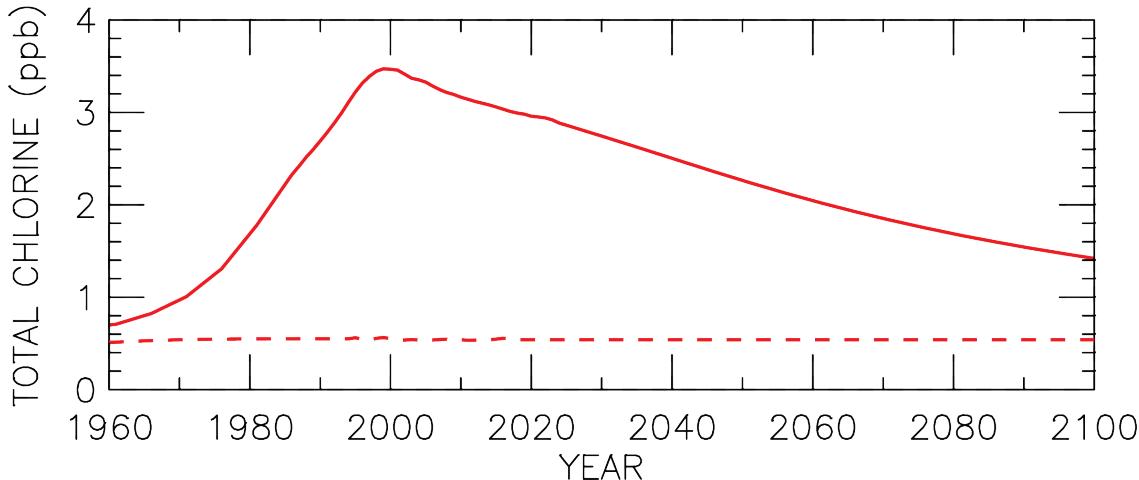
Next Exam:

- week from today (17 Nov), conducted in the same format as first exam:  
i.e., open book, open web, normal class time, via Zoom with camera on
- conceptual focus mainly on material covered in Lectures 11 to 17

Thursday, 12 Nov:

- review of Lectures 11 to 17
- please hold off on questions about exam until Thursday

# Recovery of the Ozone Layer



Time series of **chlorine** content of organic halocarbons that reach the stratosphere. Past values based on direct atmospheric observation. Future values based on projections that include the lifetime for removal of each halocarbon.

Table 6-4, WMO/UNEP 2018

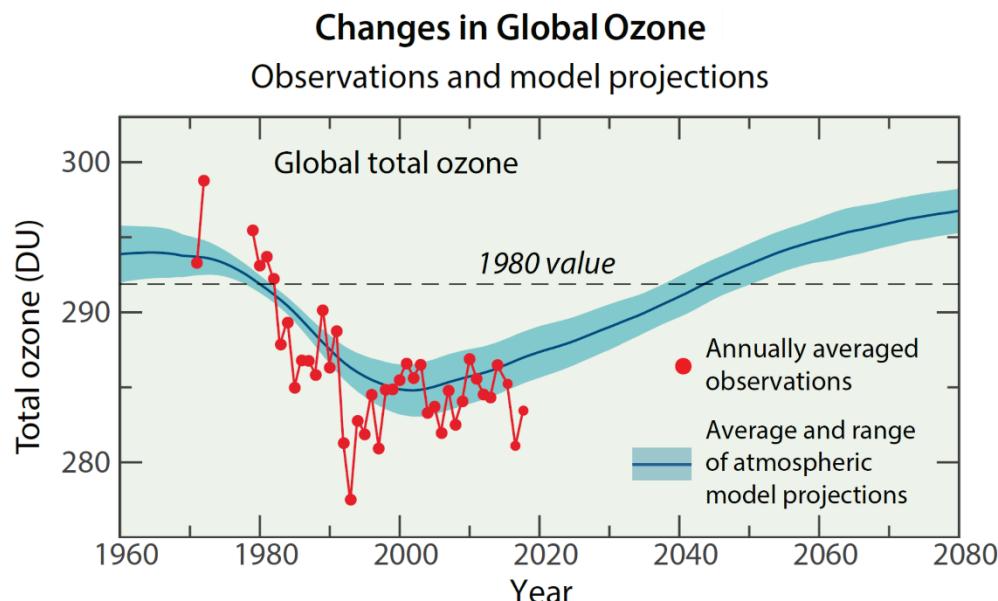
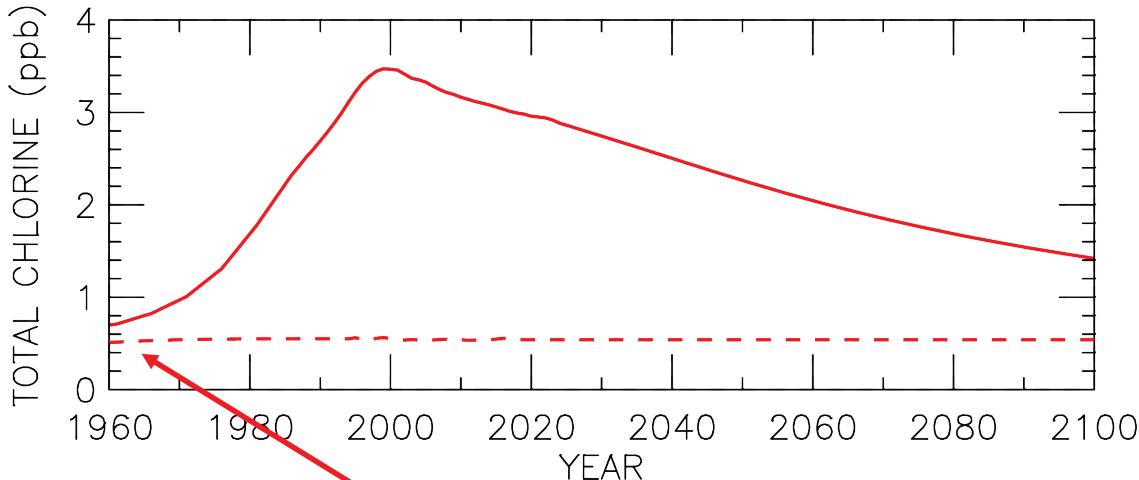


Fig Q20-1, WMO/UNEP Twenty QAs Ozone

# Recovery of the Ozone Layer



Why was total chlorine ~0.6 ppb in 1960?

Time series of **chlorine** content of organic halocarbons that reach the stratosphere. Past values based on direct atmospheric observation. Future values based on projections that include the lifetime for removal of each halocarbon.

Table 6-4, WMO/UNEP 2018

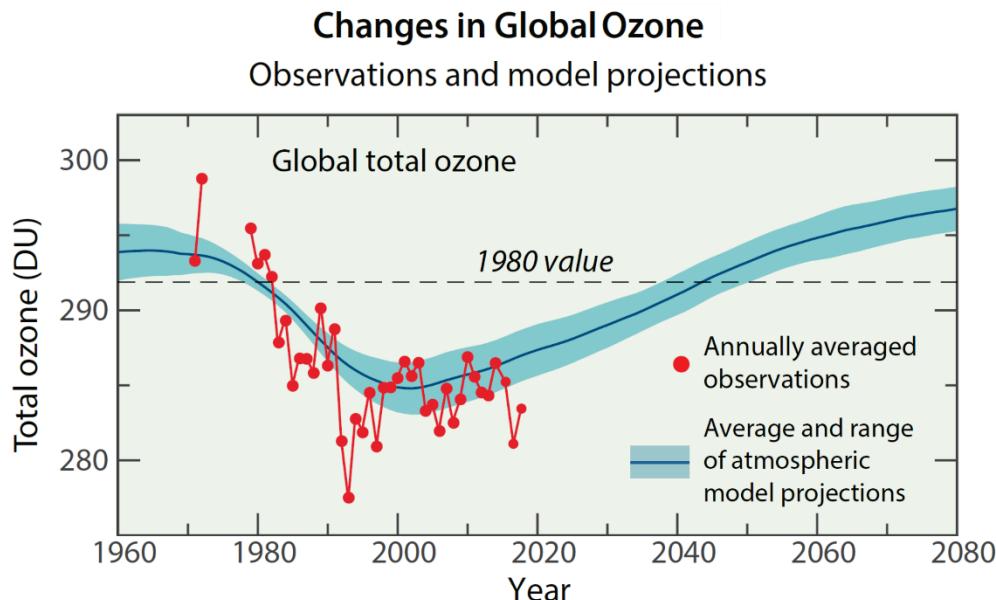
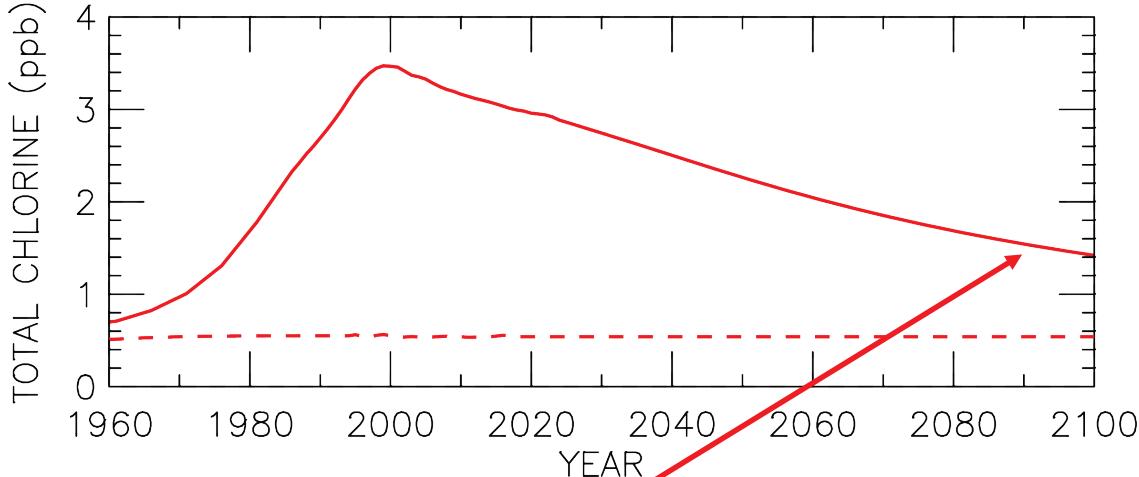


Fig Q20-1, WMO/UNEP Twenty QAs Ozone

# Recovery of the Ozone Layer



What very long lived halocarbons  
are responsible for  $\text{CCl}_y$  in year 2100  
exceeding  $\text{CCl}_y$  in year 1960 ?

Time series of **chlorine** content of organic halocarbons that reach the stratosphere. Past values based on direct atmospheric observation. Future values based on projections that include the lifetime for removal of each halocarbon.

Table 6-4, WMO/UNEP 2018

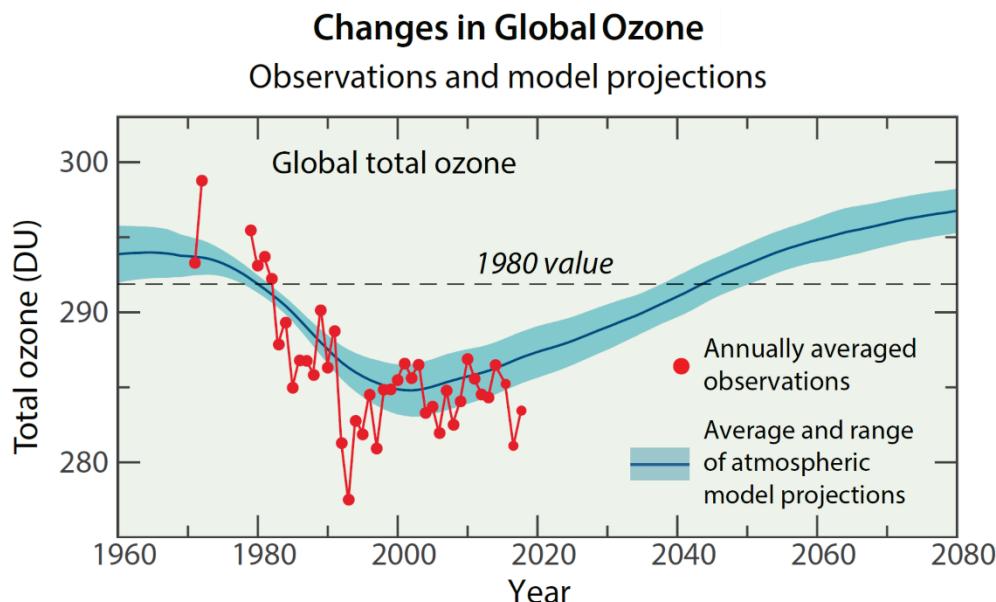
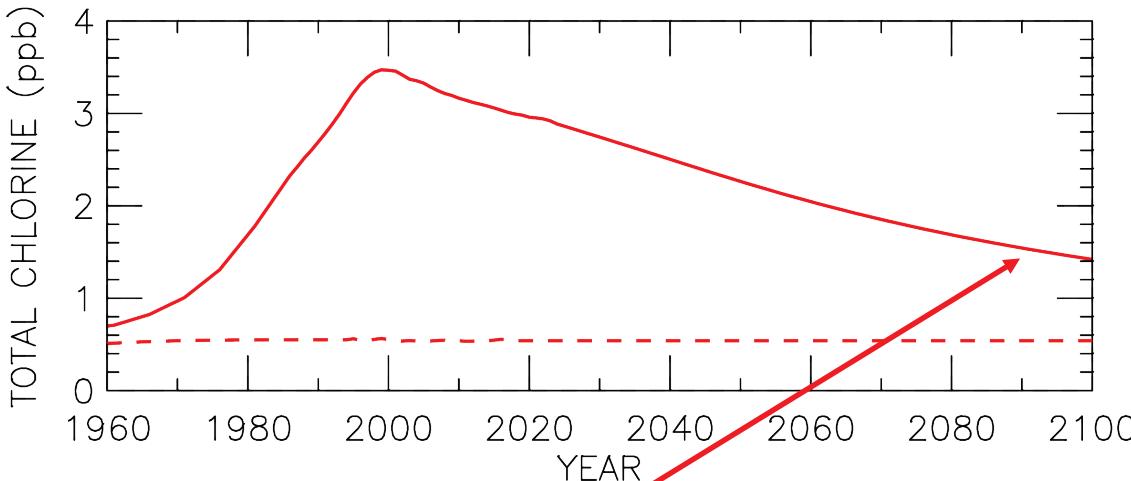


Fig Q20-1, WMO/UNEP Twenty QAs Ozone

# Recovery of the Ozone Layer

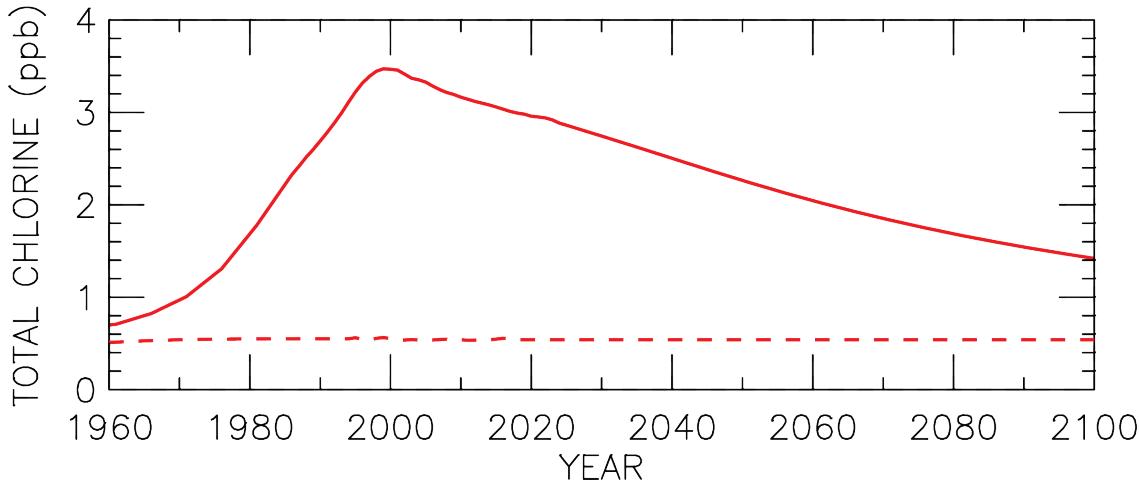


Time series of **chlorine** content of organic halocarbons that reach the stratosphere. Past values based on direct atmospheric observation. Future values based on projections that include the lifetime for removal of each halocarbon.

Table 6-4, WMO/UNEP 2018

Gas	Atmospheric Lifetime (years)	Ozone Depletion Potential (ODP) <sup>b</sup>
<b>Halogen Source Gases</b>		
<b>Chlorine Gases</b>		
CFC-11 ( $\text{CCl}_3\text{F}$ )	52	1
Carbon tetrachloride ( $\text{CCl}_4$ )	32	0.87
CFC-113 ( $\text{CCl}_2\text{FCClF}_2$ )	93	0.81
CFC-12 ( $\text{CCl}_2\text{F}_2$ )	102	0.73
Methyl chloroform ( $\text{CH}_3\text{CCl}_3$ )	5.0	0.14
HCFC-141b ( $\text{CH}_3\text{CCl}_2\text{F}$ )	9.4	0.102
HCFC-142b ( $\text{CH}_3\text{CClF}_2$ )	18	0.057
HCFC-22 ( $\text{CHF}_2\text{Cl}$ )	12	0.034
Methyl chloride ( $\text{CH}_3\text{Cl}$ )	0.9	0.015
<b>Bromine Gases</b>		
Halon-1301 ( $\text{CBrF}_3$ )	65	15.2
Halon-1211 ( $\text{CBrClF}_2$ )	16	6.9
Methyl bromide ( $\text{CH}_3\text{Br}$ )	0.8	0.57
<b>Hydrofluorocarbons (HFCs)</b>		
HFC-23 ( $\text{CHF}_3$ )	228	0
HFC-143a ( $\text{CH}_3\text{CF}_3$ )	51	0
HFC-125 ( $\text{CHF}_2\text{CF}_3$ )	30	0
HFC-134a ( $\text{CH}_2\text{FCF}_3$ )	14	0
HFC-32 ( $\text{CH}_2\text{F}_2$ )	5.4	0
HFC-152a ( $\text{CH}_3\text{CHF}_2$ )	1.6	0
HFO-1234yf ( $\text{CF}_3\text{CF}=\text{CH}_2$ )	0.03	0

# Recovery of the Ozone Layer



Time series of **chlorine** content of organic halocarbons that reach the stratosphere. Past values based on direct atmospheric observation. Future values based on projections that include the lifetime for removal of each halocarbon.

Table 6-4, WMO/UNEP 2018

What is especially “odd” about these graphs when scrutinized in tandem ?

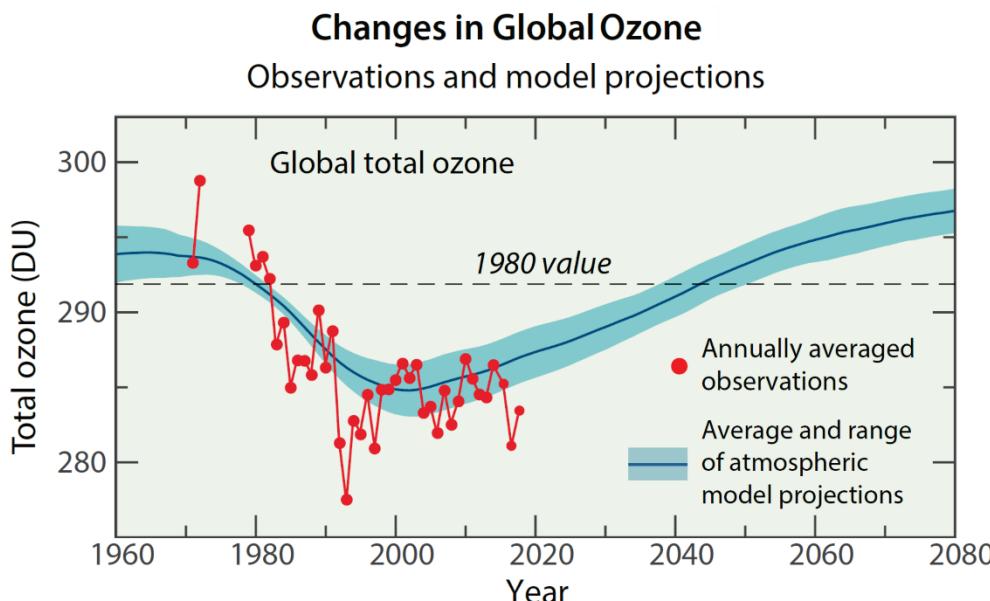
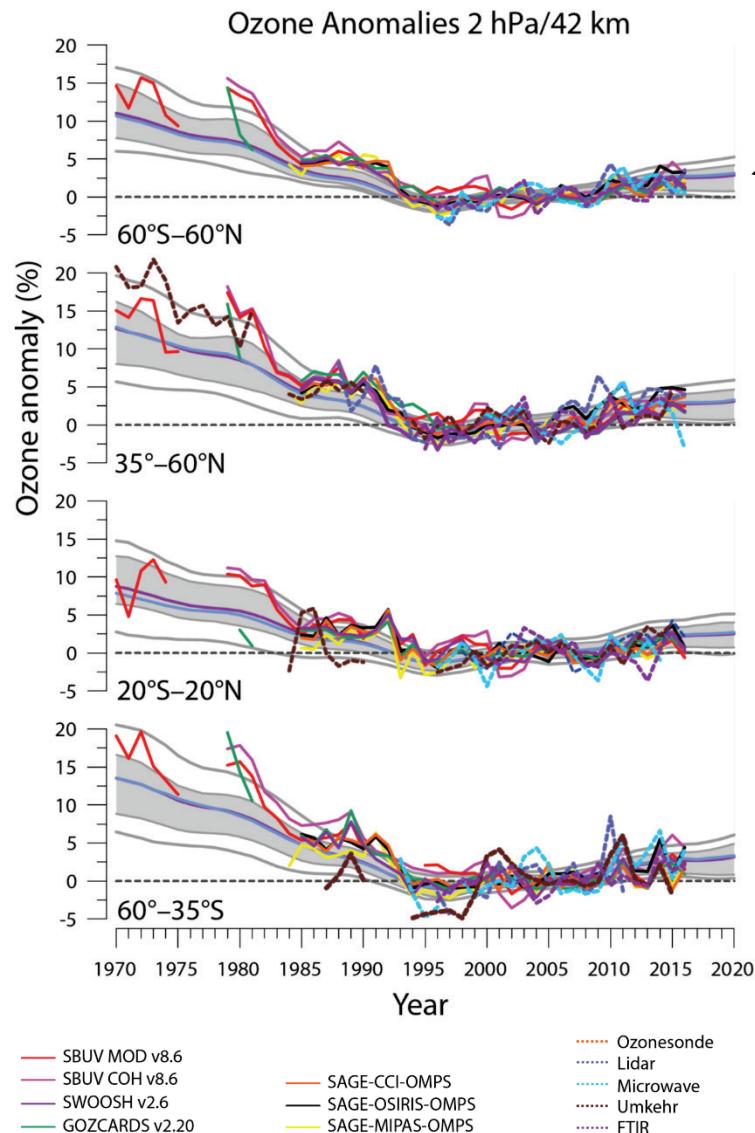


Fig Q20-1, WMO/UNEP Twenty QAs Ozone

# Past Trends, Upper Stratospheric Ozone



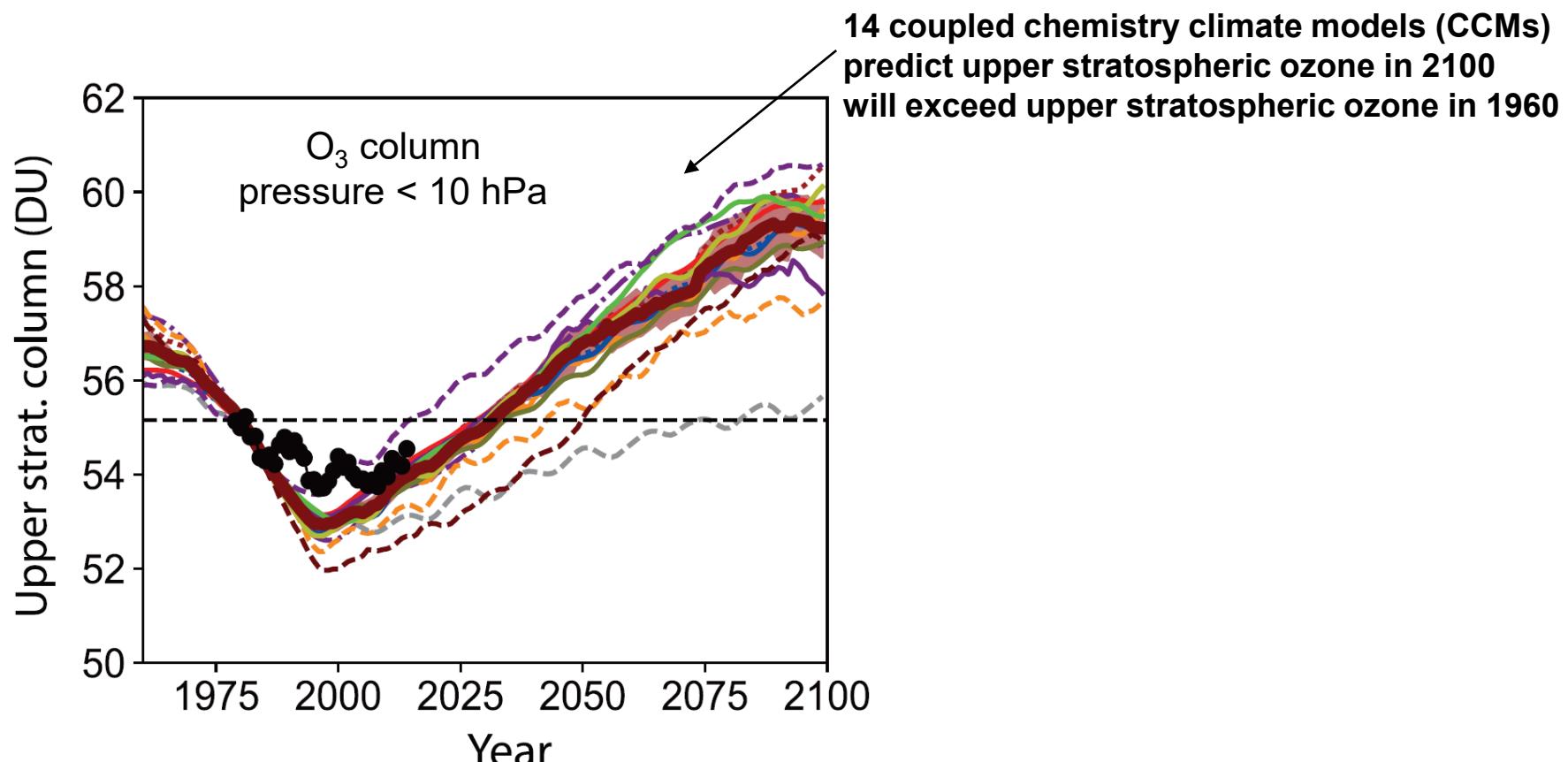
Grey: range of model calculations,  
where models are forced by  
rising levels of stratospheric  
halogens

Trends in ozone at 40 km are “well understood”  
and generally follow track time history of  
stratospheric chlorine loading.

Fig 3-15, WMO/UNEP Ozone Report

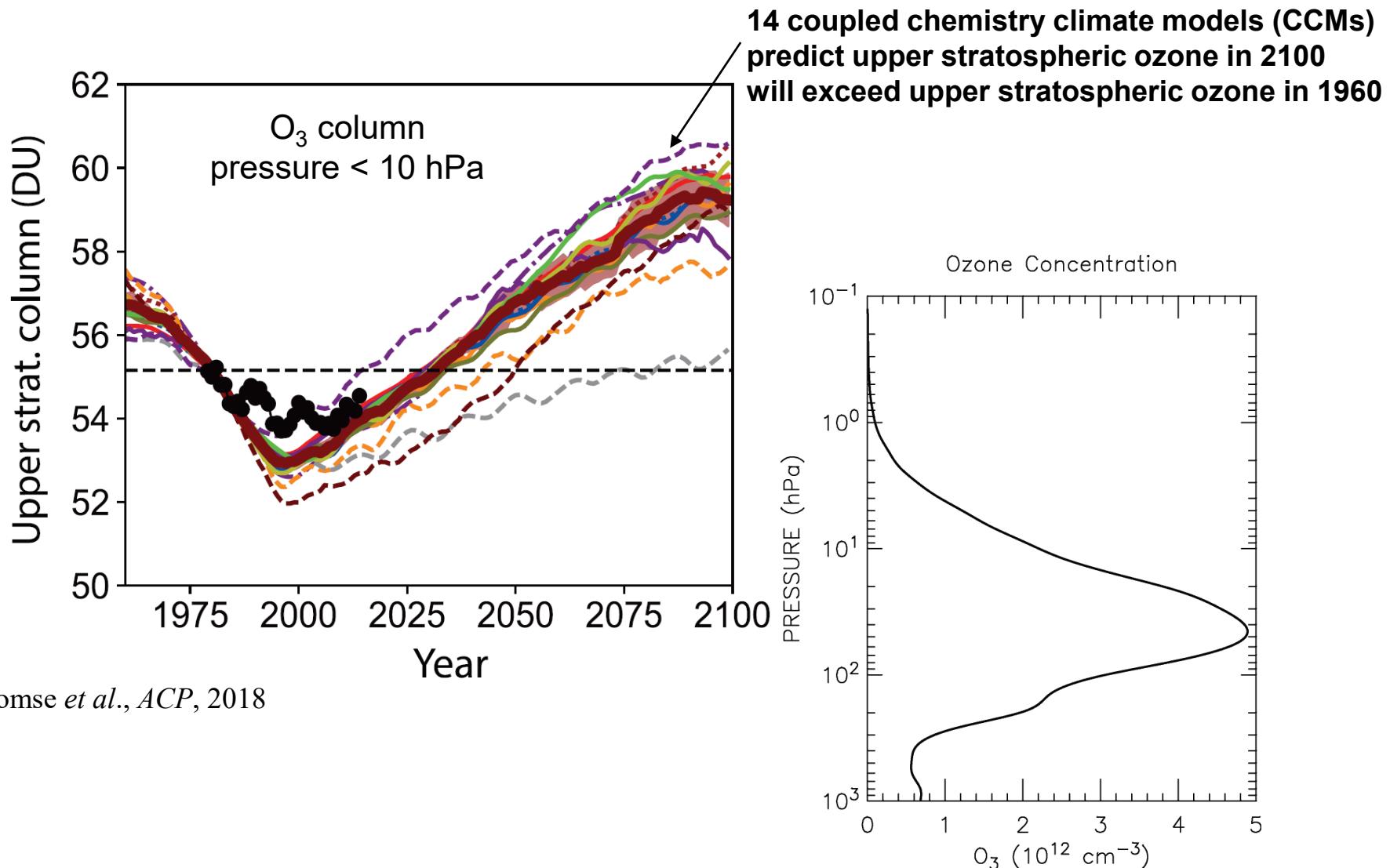
Lecture 15, Slides 50 & 51

# Future Trends, Upper Stratospheric Ozone

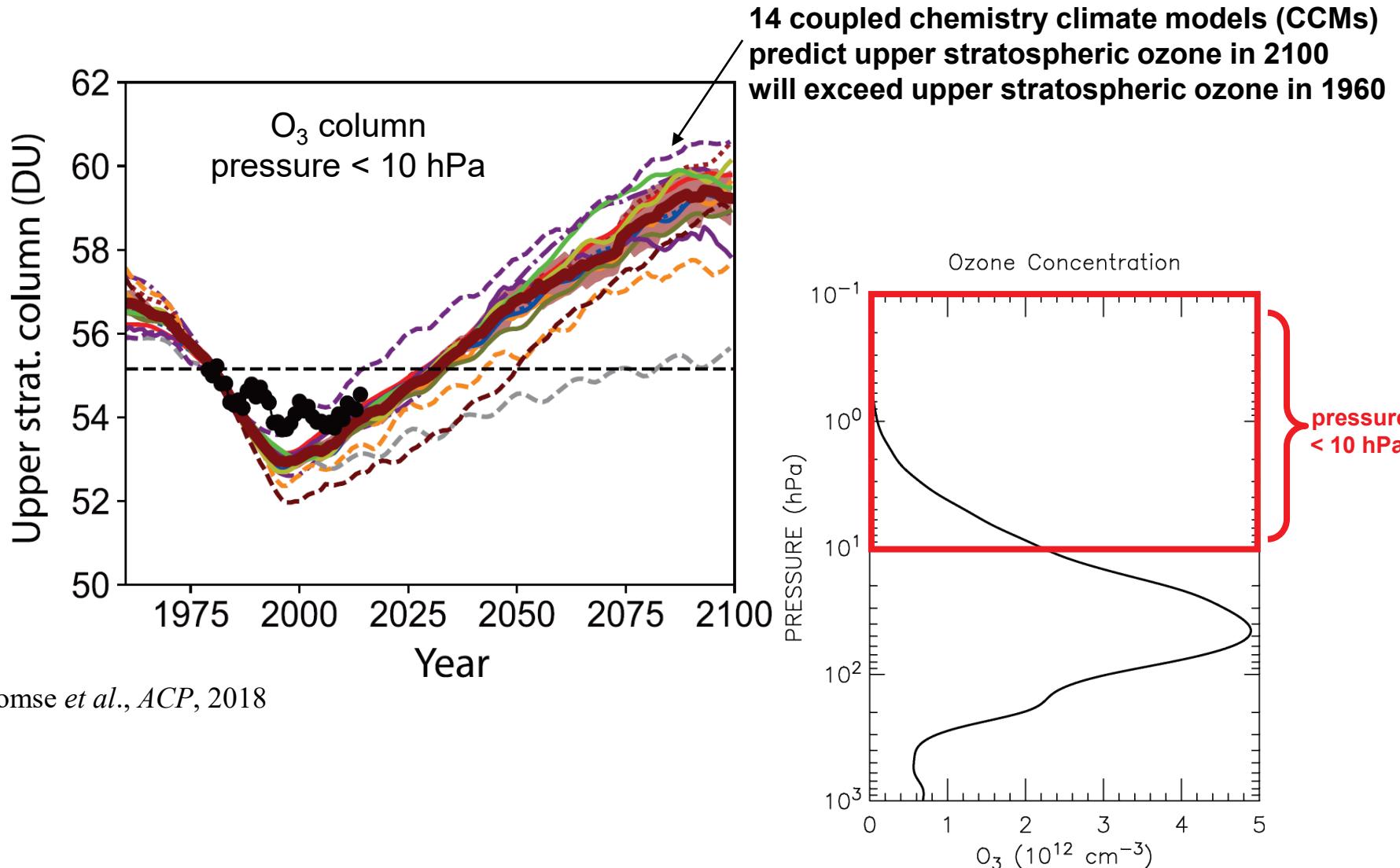


Dhomse *et al.*, ACP, 2018

# Future Trends, Upper Stratospheric Ozone



# Future Trends, Upper Stratospheric Ozone



# Climate and Chemistry Coupling

**Scientists have long known that rising GHGs leads to cooling of the stratosphere, due to direct radiative effects**

**The stratosphere has been cooling past several decades in a manner broadly consistent with theory:**

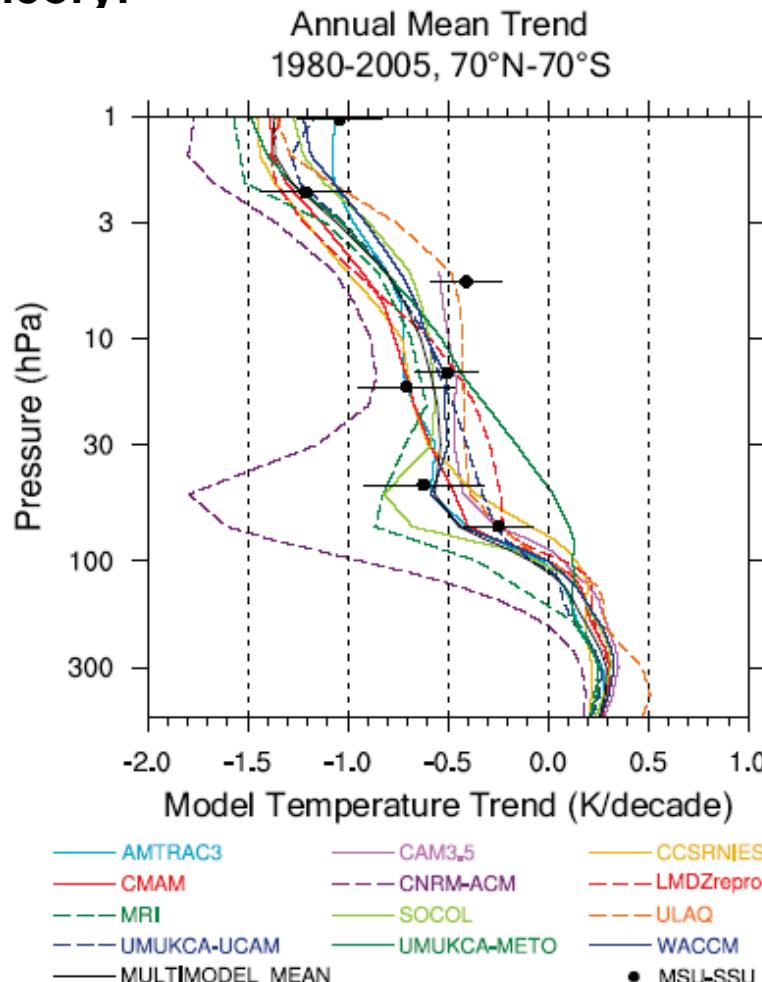
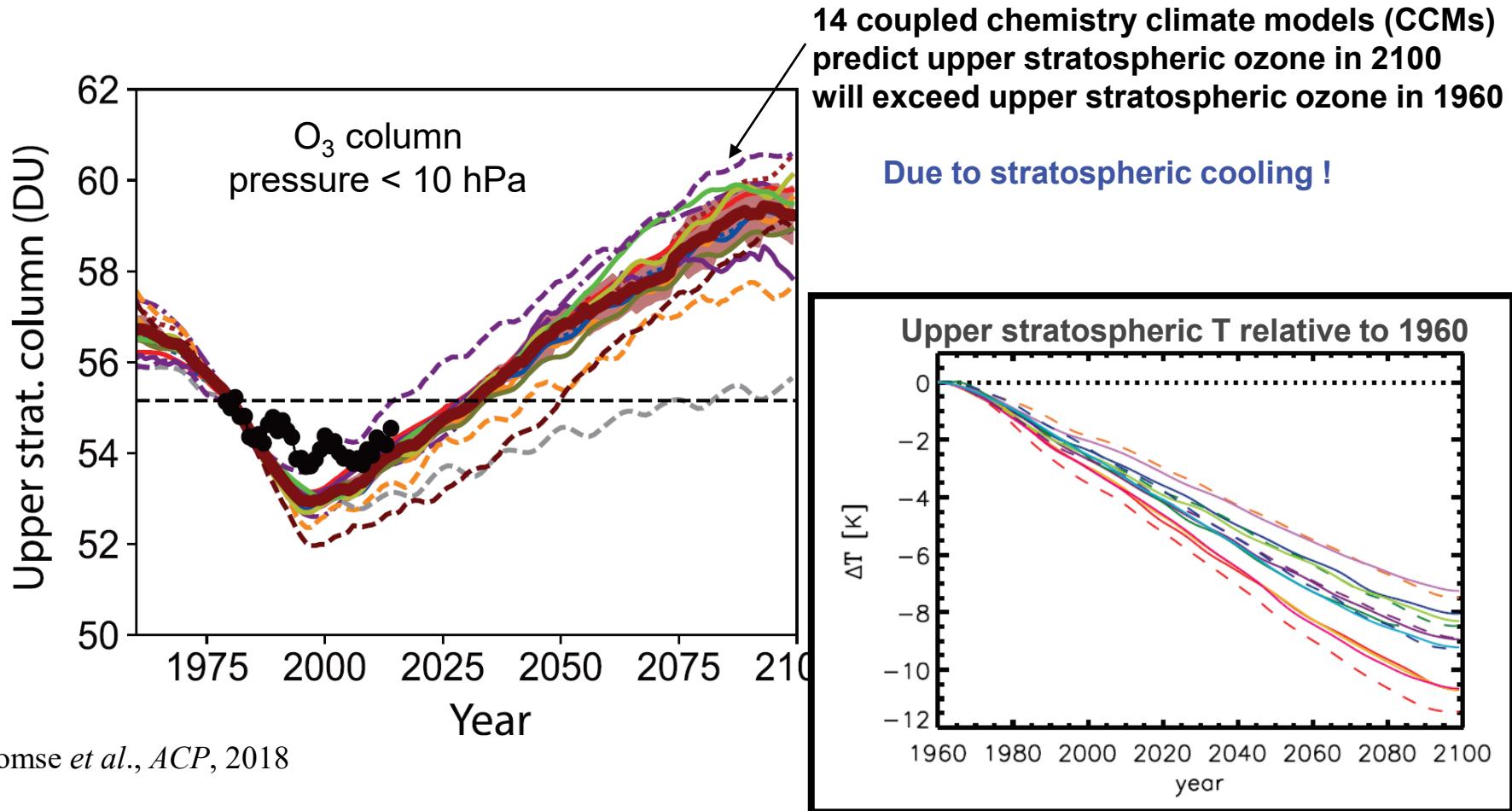


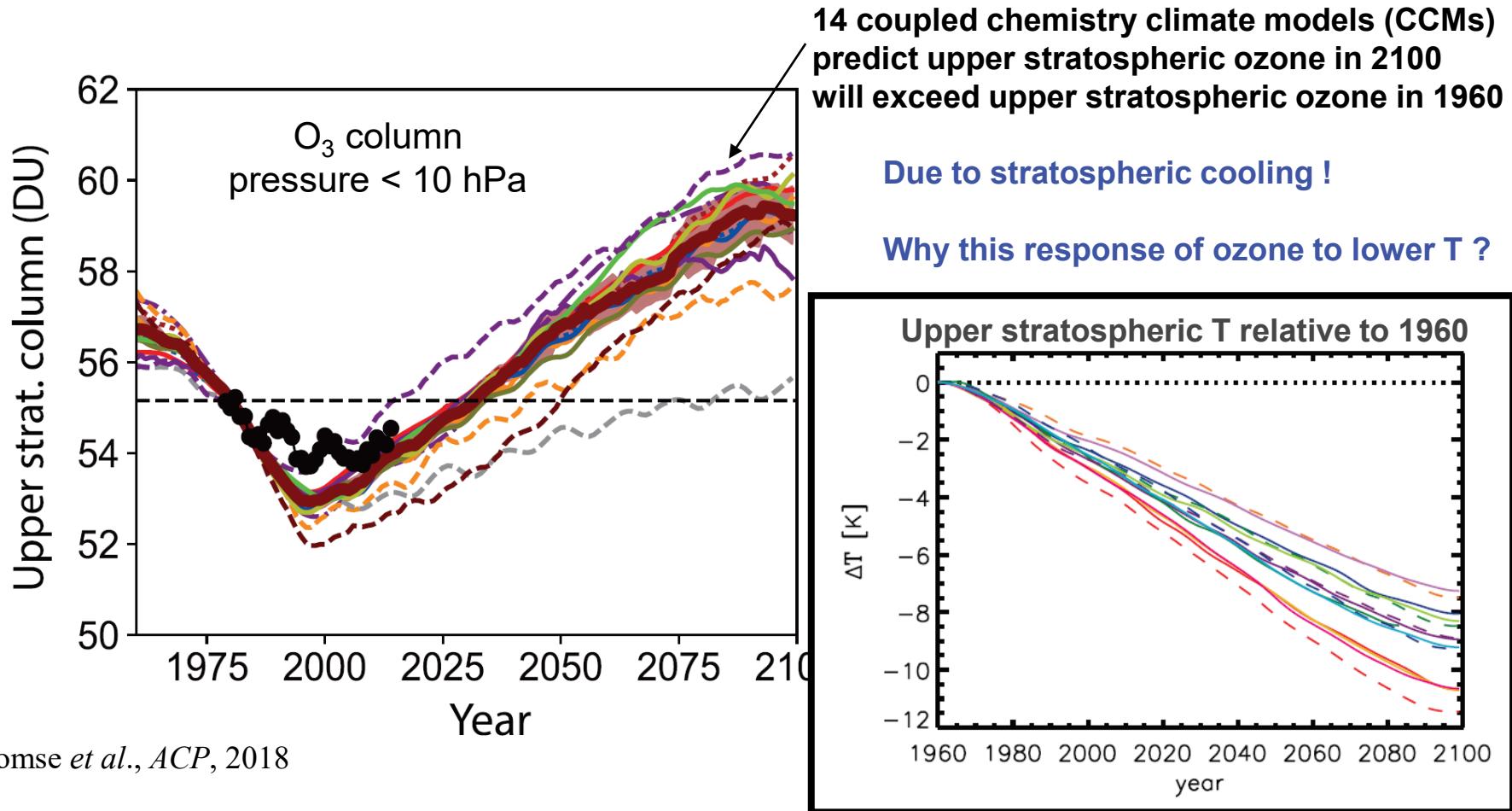
Figure 4-11  
WMO/UNEP (2011)

Lecture 16, Slide 54

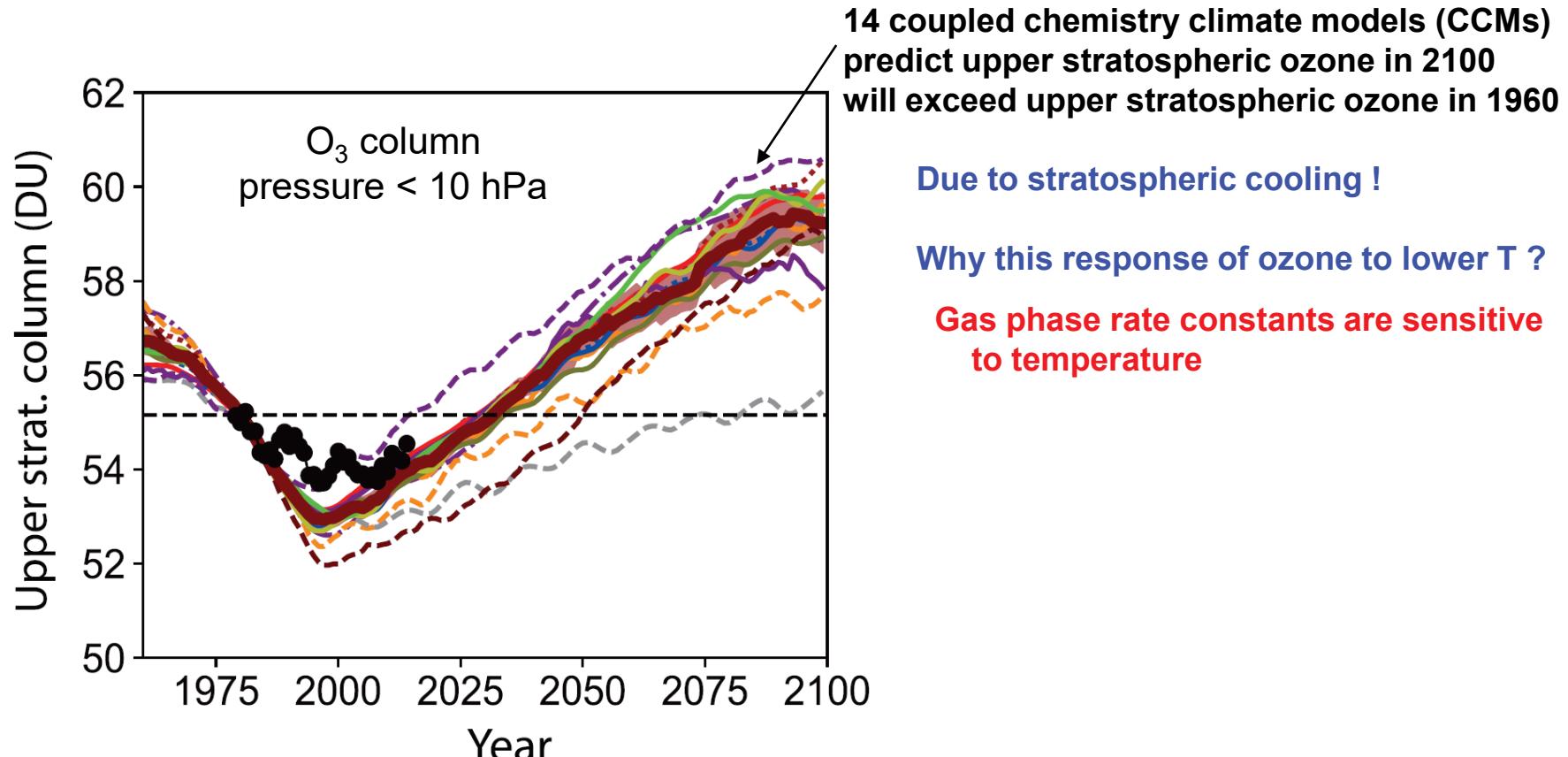
# Future Trends, Upper Stratospheric Ozone



# Future Trends, Upper Stratospheric Ozone

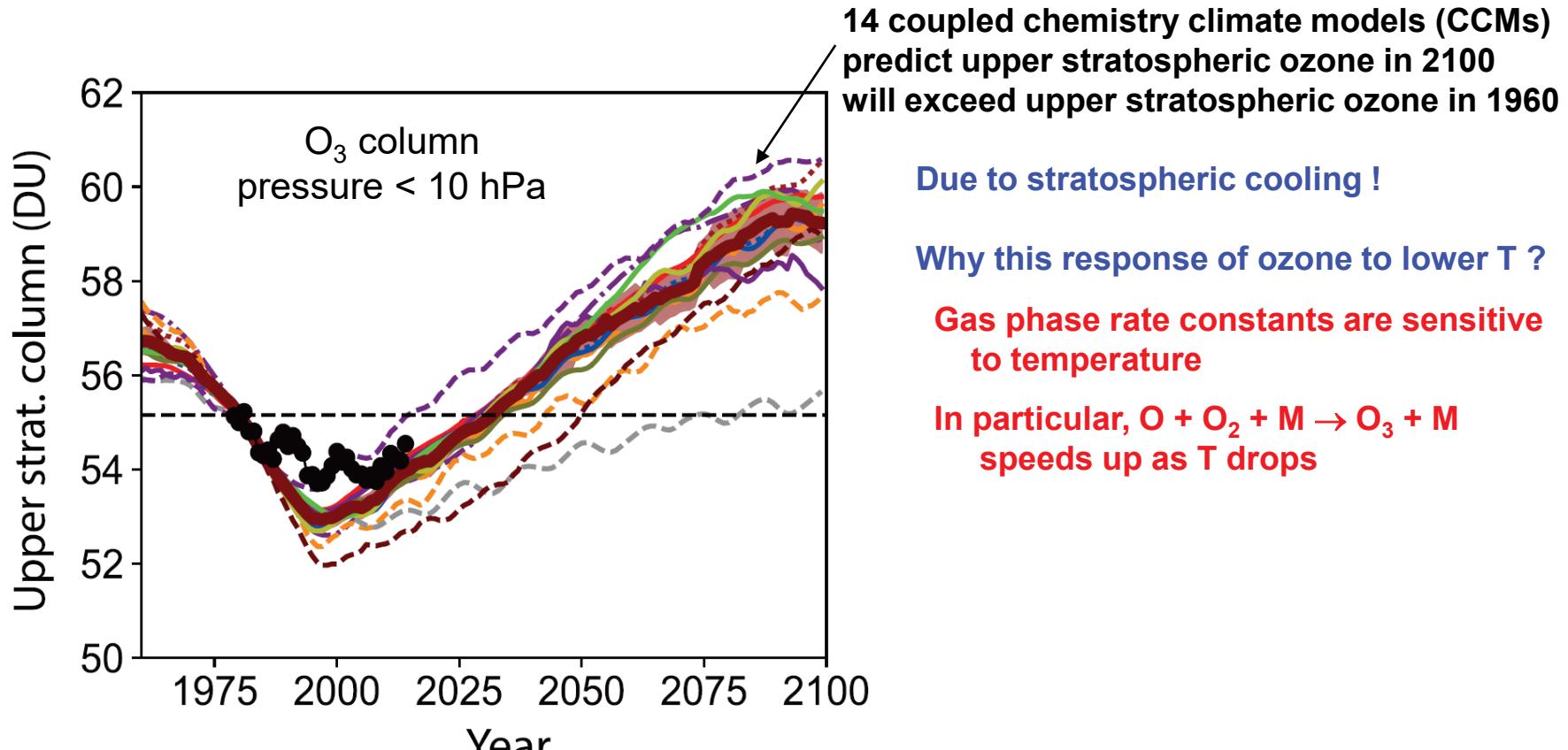


# Future Trends, Upper Stratospheric Ozone



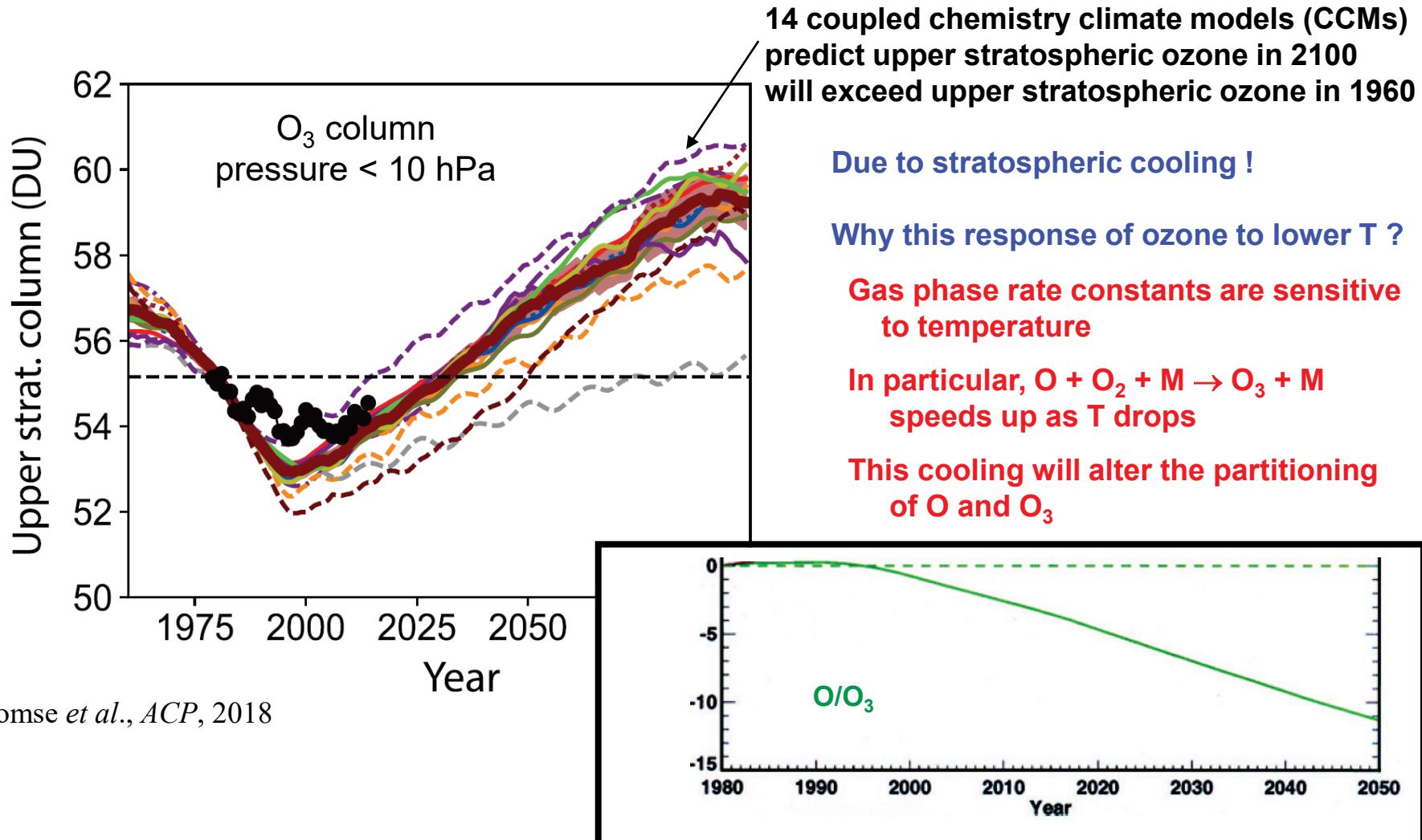
Dhomse *et al.*, ACP, 2018

# Future Trends, Upper Stratospheric Ozone

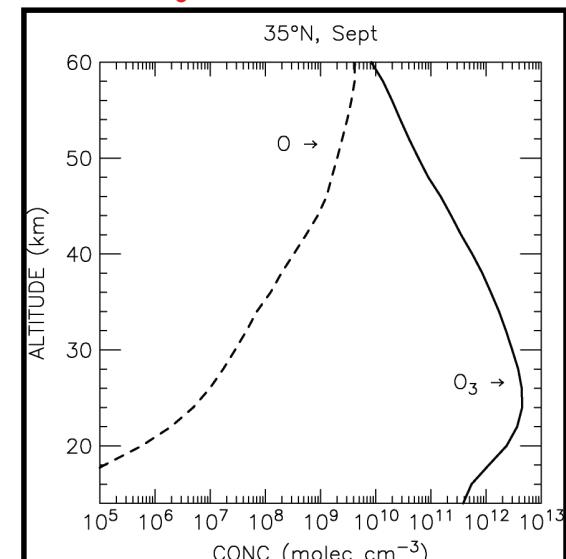
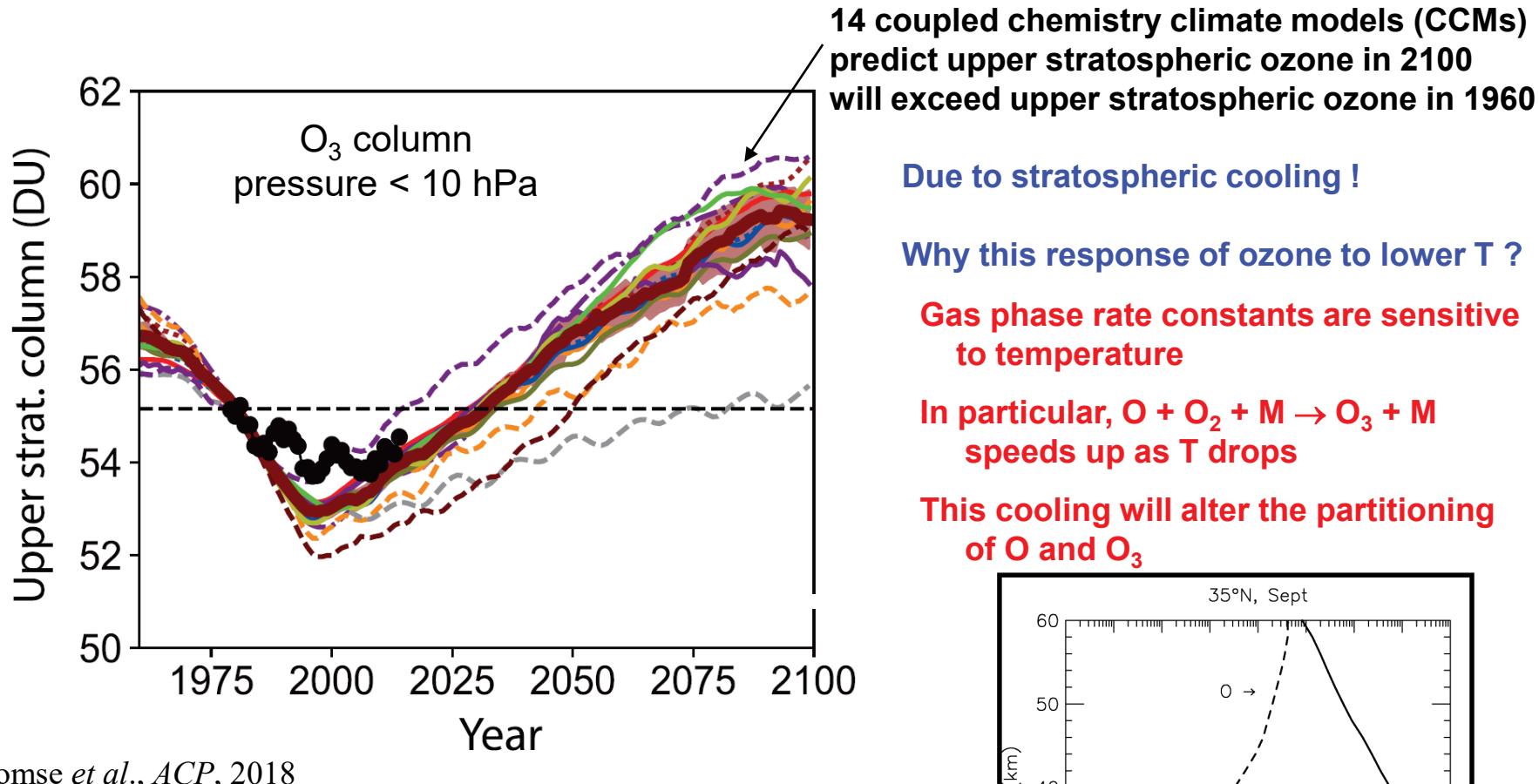


Dhomse *et al.*, ACP, 2018

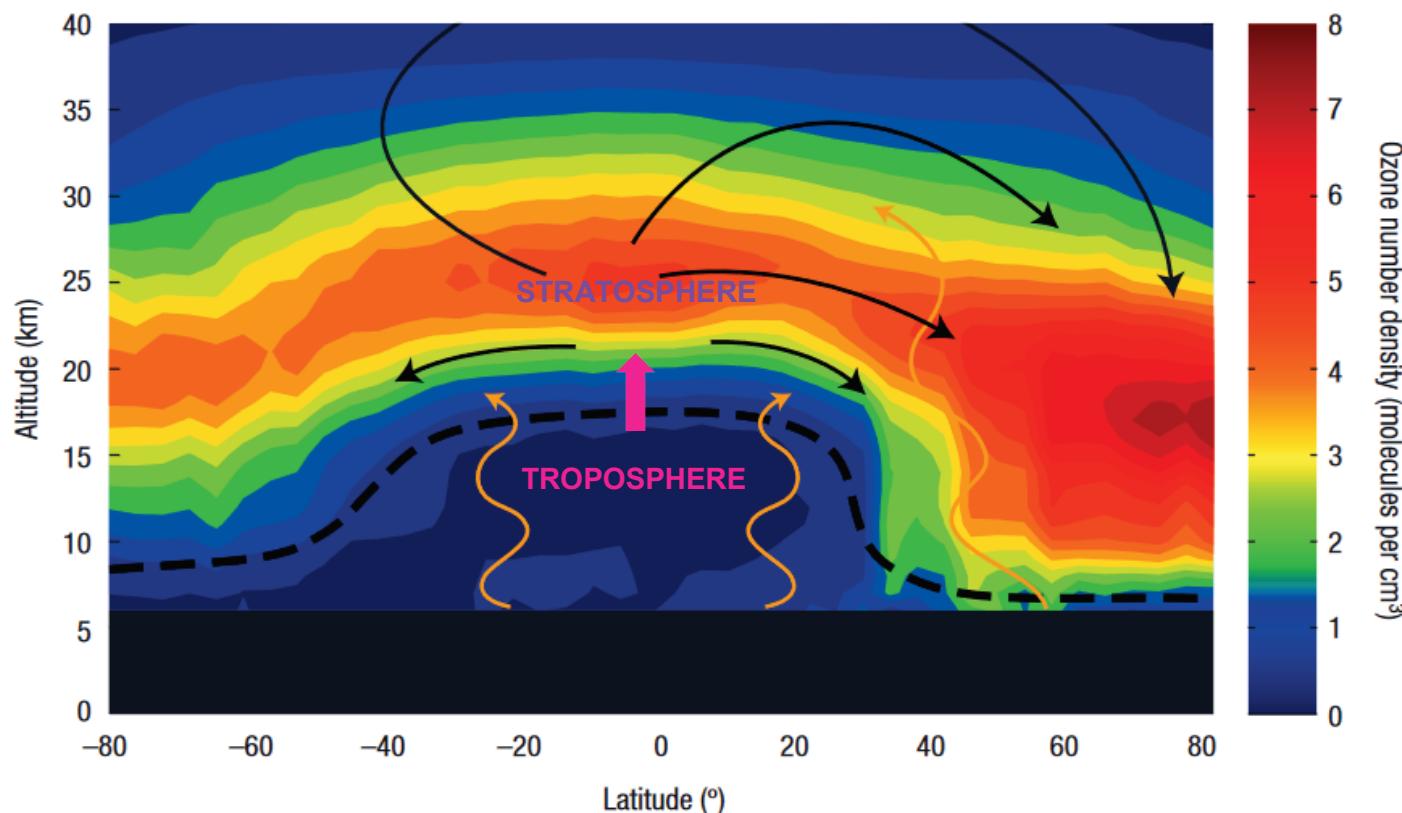
# Future Trends, Upper Stratospheric Ozone



# Future Trends, Upper Stratospheric Ozone



# Tropopause versus Latitude



Brewer–Dobson circulation (arrows), ozone (colors), and tropopause (black dashed line).

Shaw and Shepherd, Nature Geoscience, 2008.

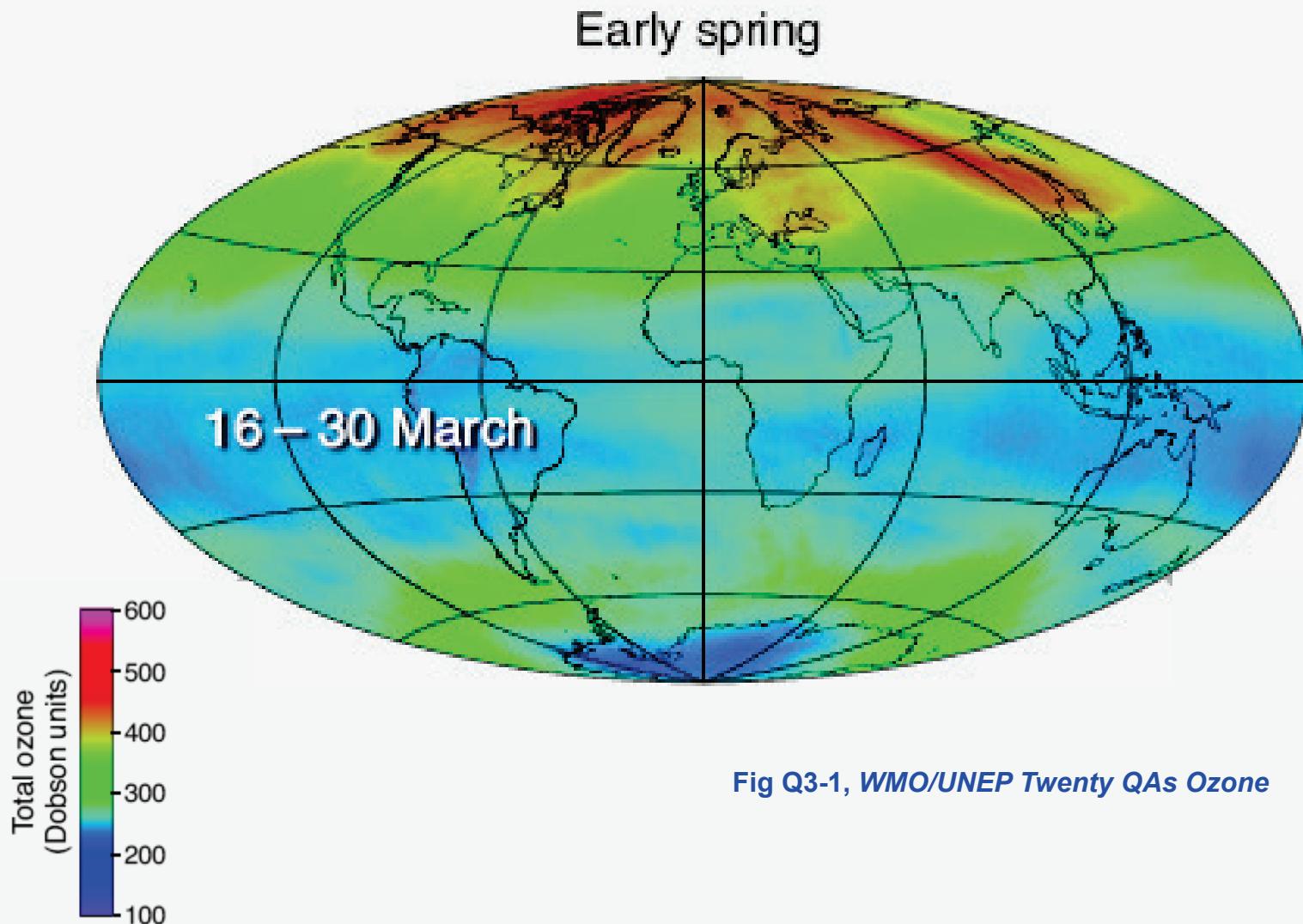
Brewer-Dobson Circulation is a model of atmospheric circulation, proposed by Alan Brewer in 1949 and Gordon Dobson in 1956, that attempts to explain why tropical air has less column ozone than polar air, even though the tropical stratosphere is where most atmospheric ozone is produced

Lecture 3, Slide 43

&

Lecture 12, Slides 46-50

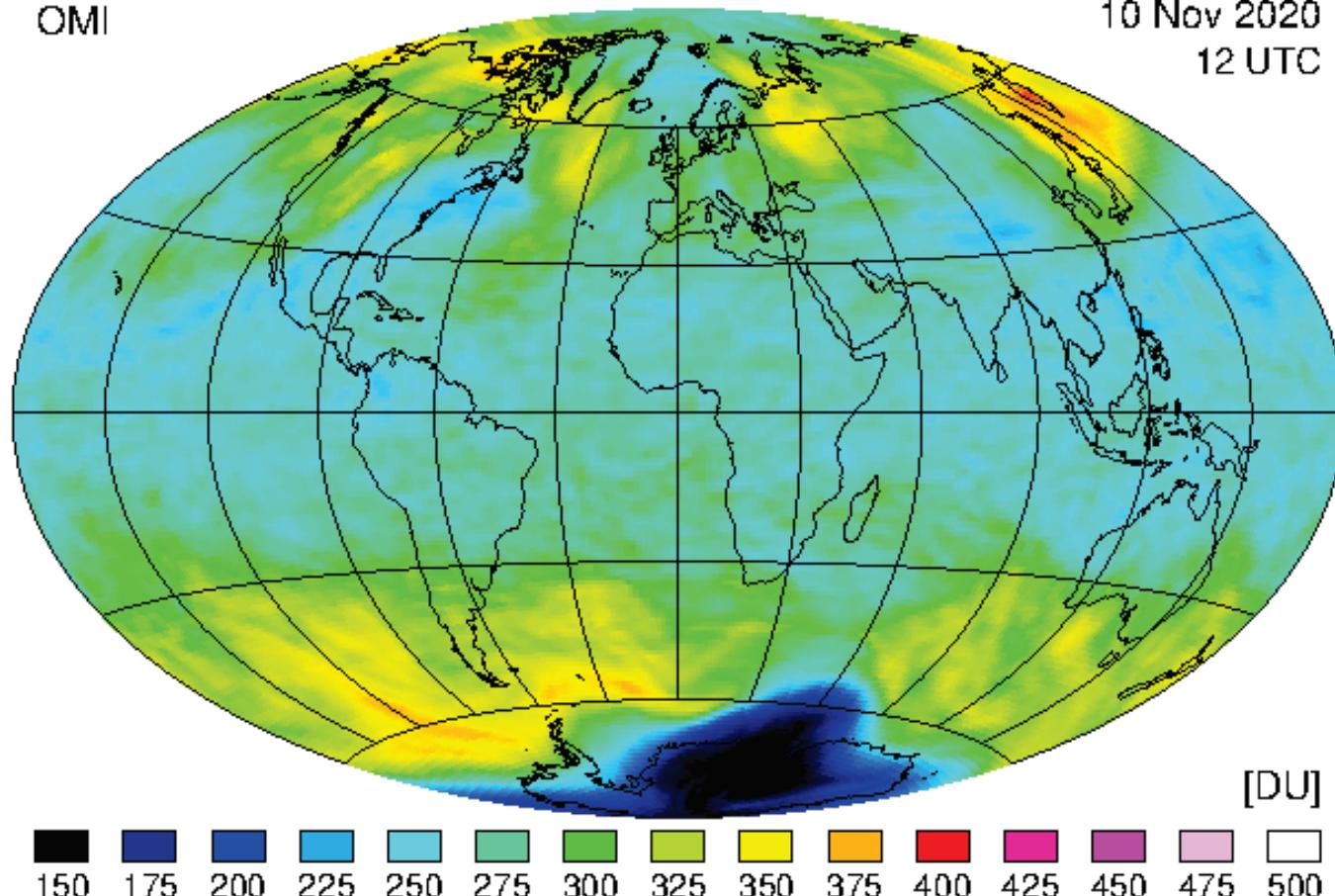
# Global Satellite Maps of Total Ozone in 2009



# Global Satellite Maps of Total Ozone in 2009

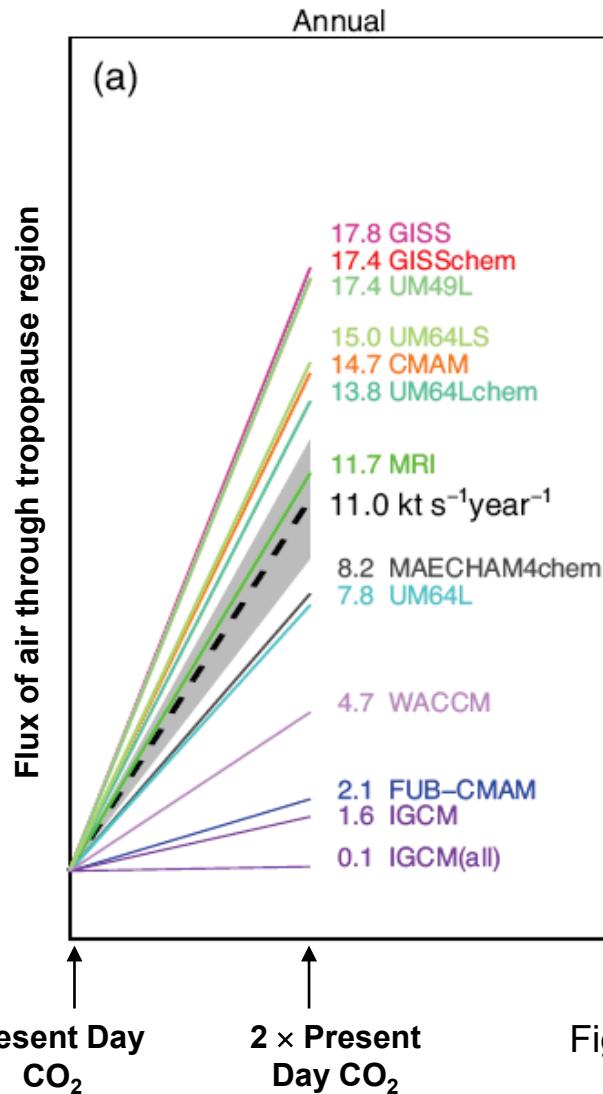
KNMI / NASA  
OMI

Forecast total ozone (D+1)  
10 Nov 2020  
12 UTC



[http://www.temis.nl/protocols/o3field/data/omi/forecast/today\\_wd.gif](http://www.temis.nl/protocols/o3field/data/omi/forecast/today_wd.gif)

# More Chemistry and Climate Coupling



**Figure 5-17.** Trends in exchange of air from troposphere-to-stratosphere computed by 14 CCMs.

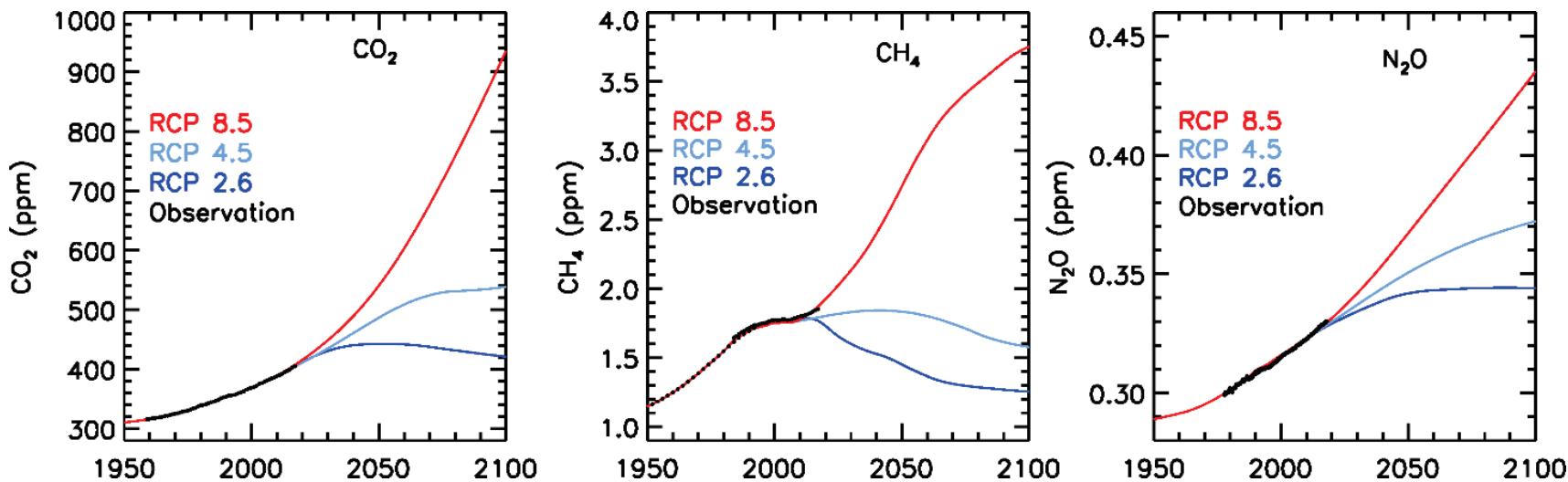
Trends (units of Gg s<sup>-1</sup> year<sup>-1</sup>) are represented by the slope of each line.

Dashed line is the multi-model mean.

After Butchart *et al.*, *Clim. Dyn.*, 2006.

Fig 5.17, WMO/UNEP (2006)

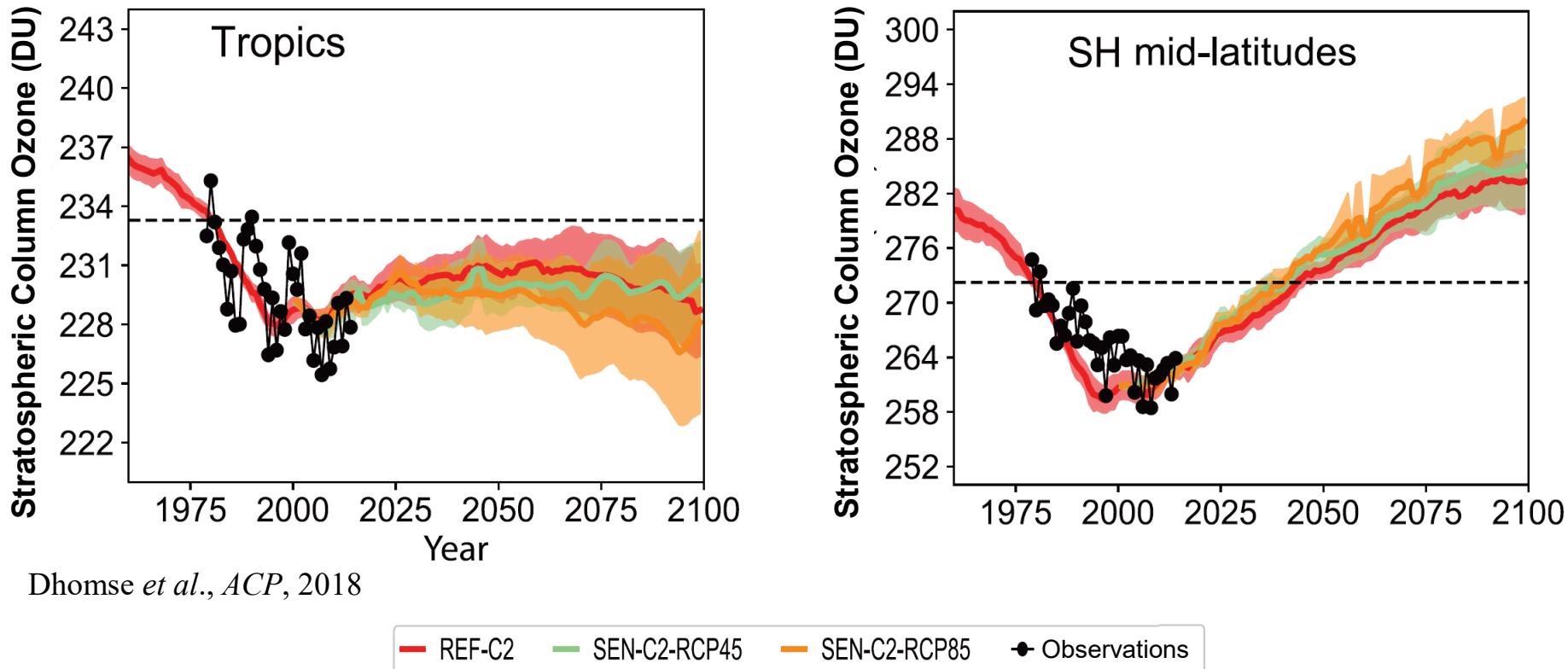
# RCP Scenarios Forecast Wide Range of Possible Futures



- RCP: Representative Concentration Pathway  
Number represents  $\Delta\text{RF}$  of climate ( $\text{W m}^{-2}$ ) at the end of this century
- GHG mixing ratio time series for CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O, as well as CFCs, HCFCs, and HFCs that are provided to climate model groups

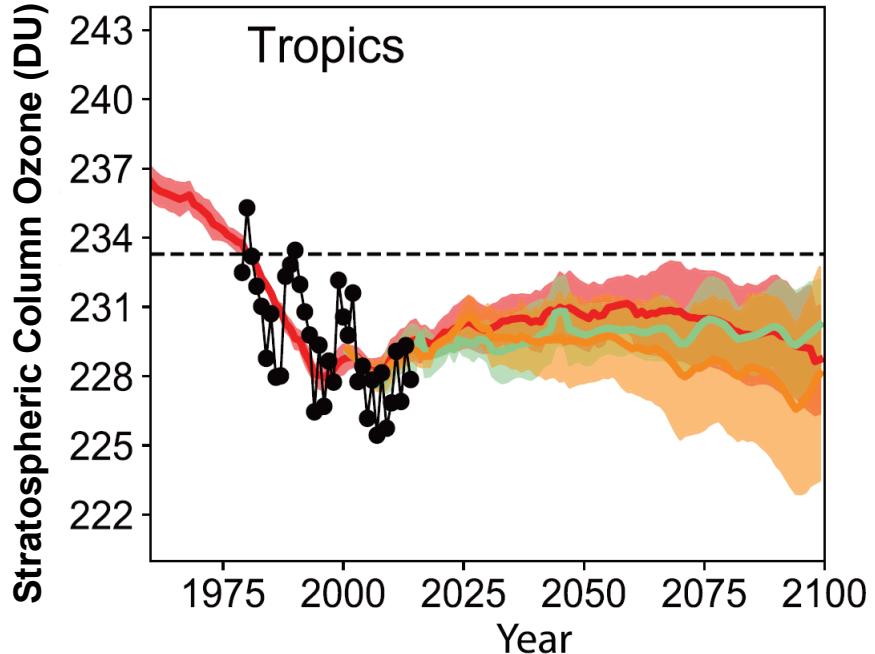
Figure from *Hope et al., 2020*: <https://www.esoar.org/pdfjs/10.1002/essoar.10504179.1>

# More Chemistry and Climate Coupling



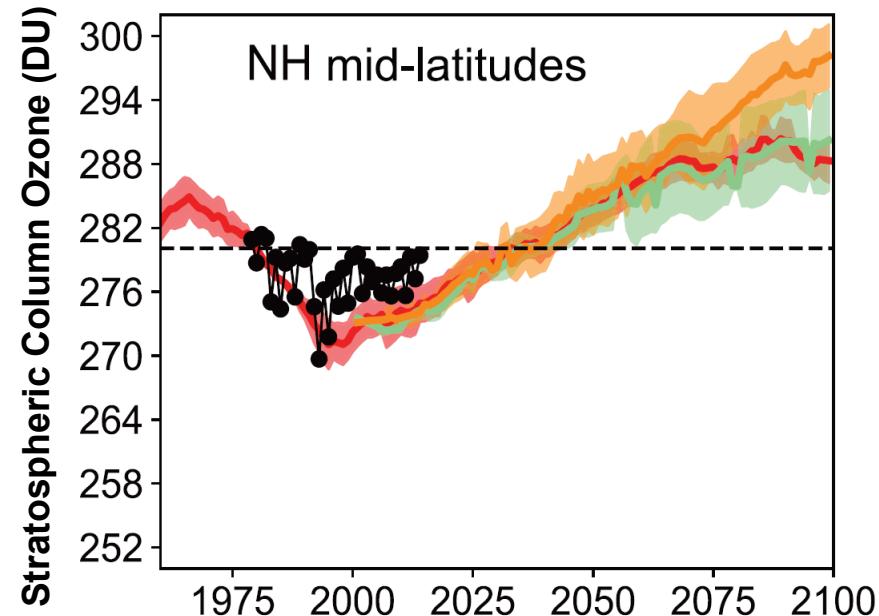
Acceleration of the *Brewer-Dobson Circulation* causes modeled total ozone column in the tropics to exhibit a sustained, long term decline and modeled total ozone column at mid-latitudes to experience a “super recovery”

# More Chemistry and Climate Coupling



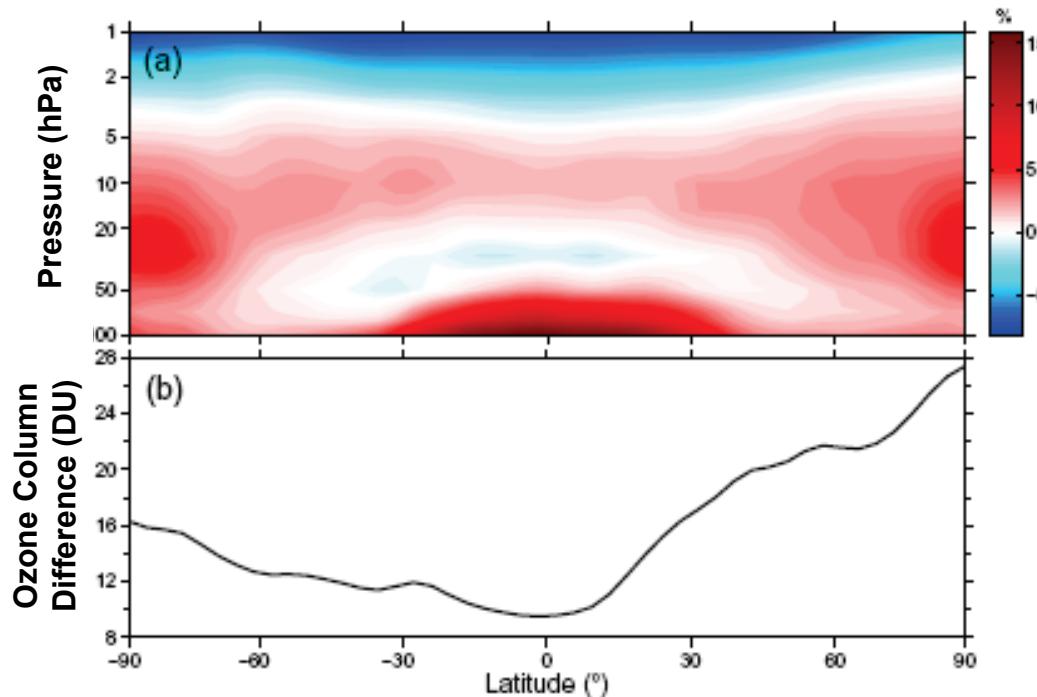
Dhomse *et al.*, *ACP*, 2018

REF-C2    SEN-C2-RCP45    SEN-C2-RCP85    Observations



Acceleration of the *Brewer-Dobson Circulation* causes modeled total ozone column in the tropics to exhibit a sustained, long term decline and modeled total ozone column at mid-latitudes to experience a “super recovery”

# $\text{CH}_4$ and Stratospheric Ozone



Revell *et al.*, ACP, 2012

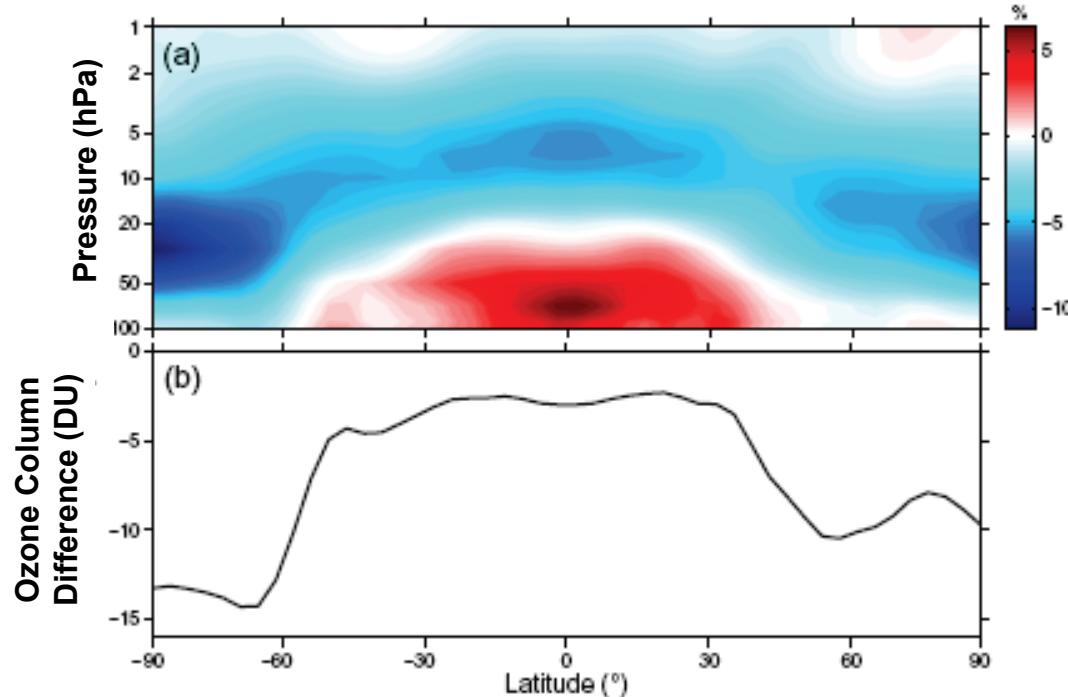
Stratospheric  $\text{O}_3$  difference in the 2090s found for a computer simulation run using  $\text{CH}_4$  from RCP 8.5 minus that of a simulation using  $\text{CH}_4$  from RCP 2.6

Rising  $\text{CH}_4$  leads to:

- a) more ozone loss in the upper stratosphere by increasing the speed of  $\text{OH}$  and  $\text{HO}_2$  ( $\text{HO}_x$ ) mediated loss cycles.
- b) a cooler stratosphere, slowing the rate of all ozone loss cycles
- c) speeding up of the the rate of  $\text{Cl}+\text{CH}_4$ , shifting chlorine from  $\text{ClO}$  into  $\text{HCl}$
- d) more  $\text{HO}_2$  in the lowermost stratosphere where there is sufficient  $\text{CO}$  to result in production of  $\text{O}_3$  by smog chemistry

**Computer models project stratospheric column  $\text{O}_3$  will increase as  $\text{CH}_4$  rises**

# $\text{N}_2\text{O}$ and Stratospheric Ozone



Revell *et al.*, *ACP*, 2012

Difference of stratospheric  $\text{O}_3$  in the 2090s for a computer simulation run using  $\text{N}_2\text{O}$  from RCP 8.5 minus that of a simulation using  $\text{N}_2\text{O}$  from RCP 2.6

Rising  $\text{N}_2\text{O}$  leads to:

- a) ozone loss in the middle & upper stratosphere by increasing the speed of  $\text{NO}$  and  $\text{NO}_2$  mediated loss cycles.
- b) speeds up the rate of  $\text{OH} + \text{NO}_2 + \text{M} \rightarrow \text{HNO}_3$  &  $\text{ClO} + \text{NO}_2 + \text{M} \rightarrow \text{CINO}_3 + \text{M}$  in the lowermost stratosphere, leading to slower ozone loss by these cycles & less  $\text{O}_3$  where these cycles dominate total loss of  $\text{O}_3$

Computer models project stratospheric column  $\text{O}_3$  will decline as  $\text{N}_2\text{O}$  rises

Lecture 6, Slide 69

# Future Ozone: ODSs, CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O

## Global Total Ozone Changes in Response to Ozone Depleting Substances and Greenhouse Gases

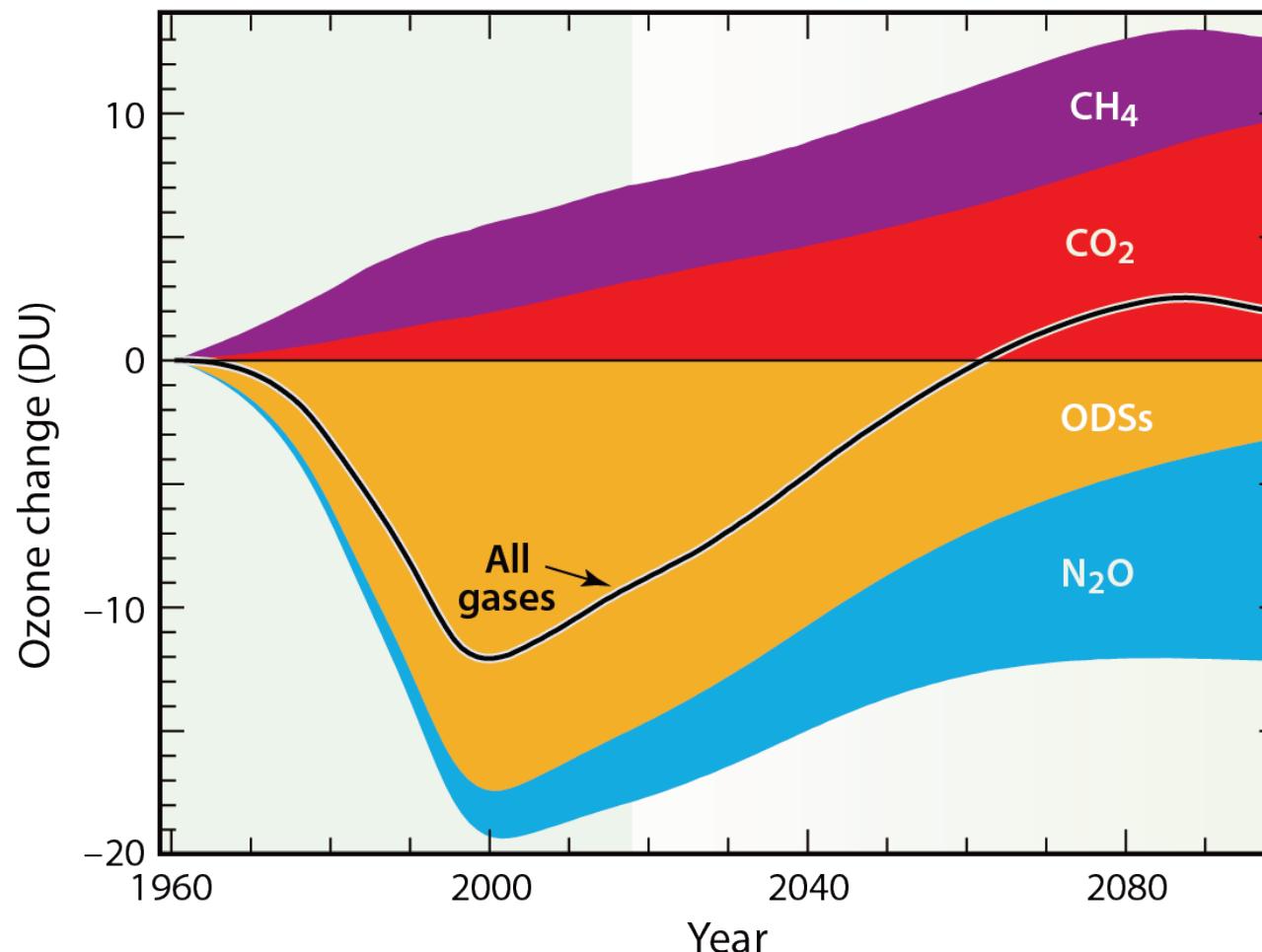


Fig Q20-3, WMO/UNEP Twenty QAs Ozone

# Future Ozone: Regional Variations

## Change in Total Ozone and Equivalent Effective Stratospheric Chlorine Since 1960

Results from atmospheric chemistry-climate models for 1960 to 2100

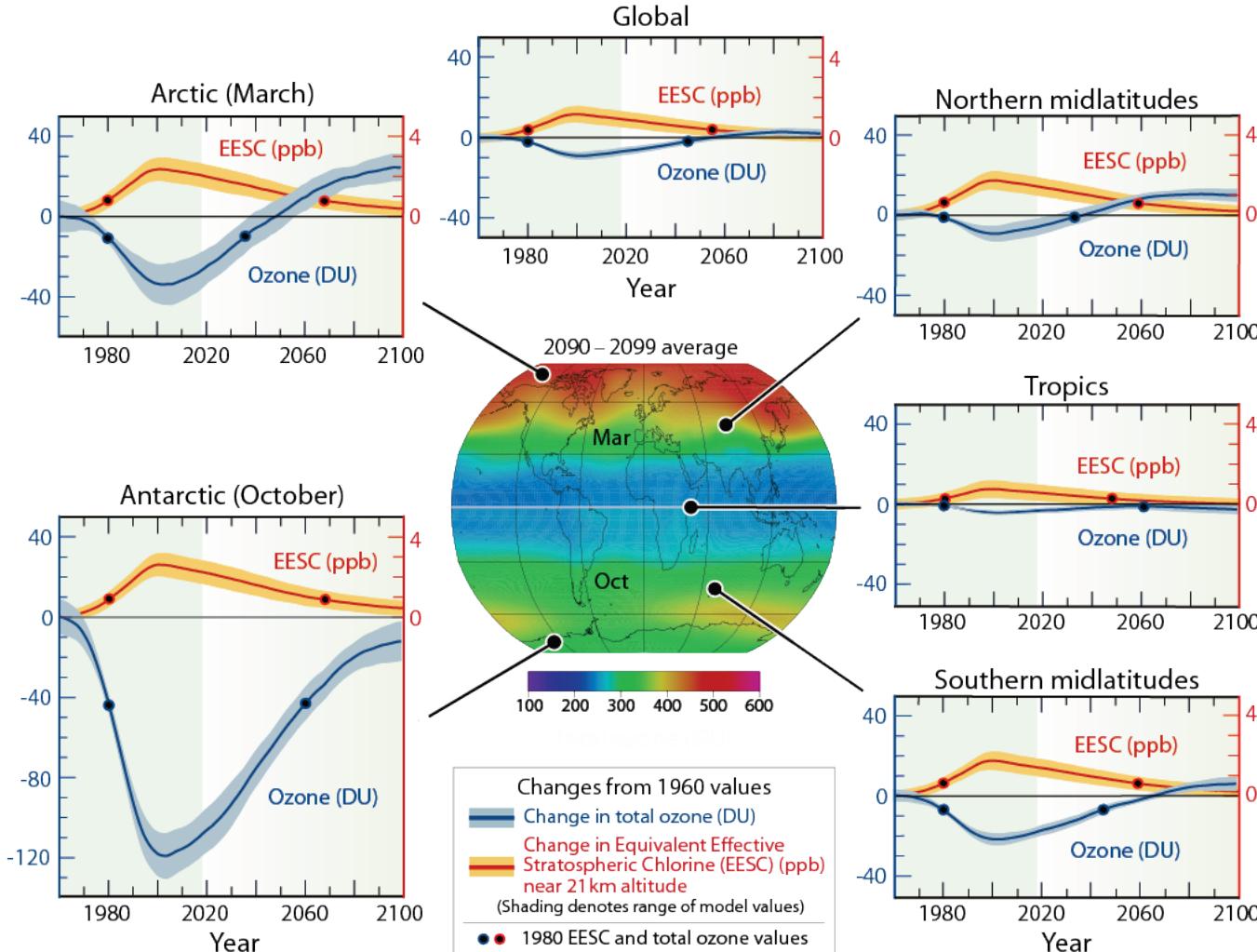
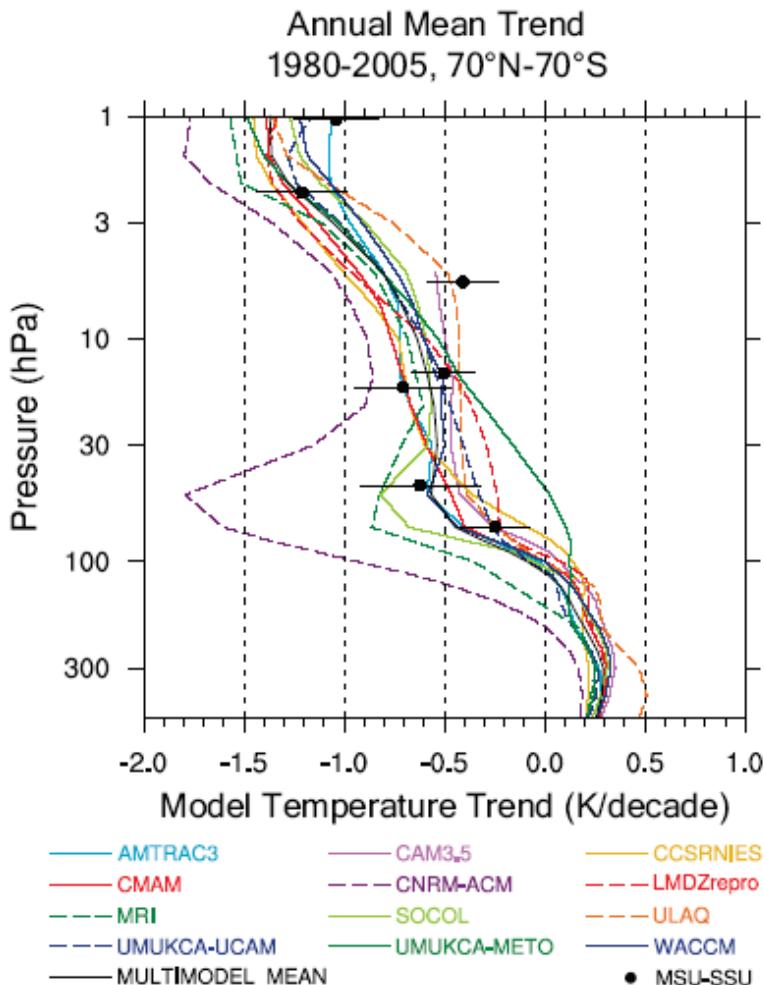


Fig Q20-2, WMO/UNEP Twenty QAs Ozone

# Future Trends, Stratospheric Ozone

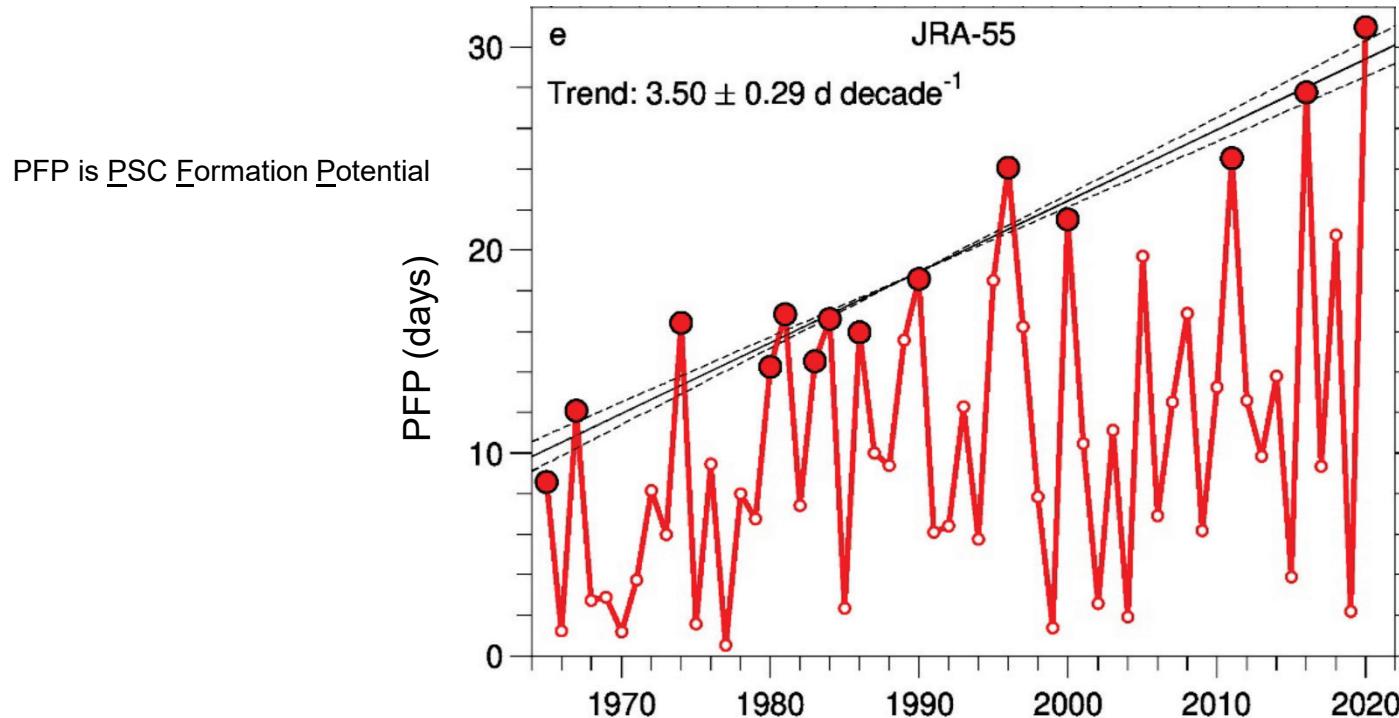


If the stratosphere continues to cool,  
for which region of the stratosphere  
is ozone “most vulnerable”?

Figure 4–11, WMO/UNEP (2011)

# Cold Arctic Winters Tend to Exhibit Larger PFP as a Function of Time

Data:

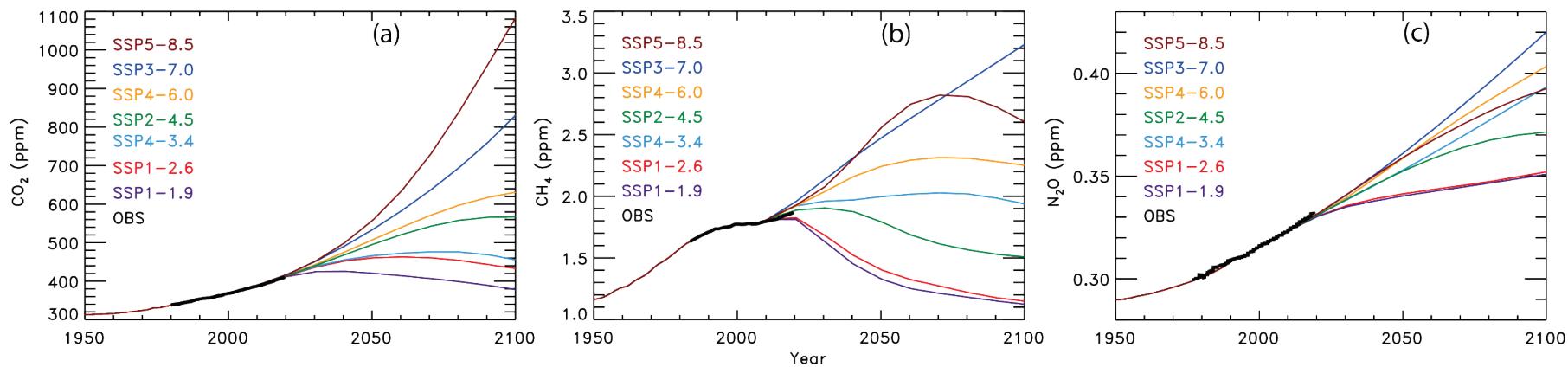


**PSC Formation Potential in Arctic Vortex**  
based on data from the Japanese 55 year re-analysis project  
[https://jra.kishou.go.jp/JRA-55/index\\_en.html](https://jra.kishou.go.jp/JRA-55/index_en.html)

SOLID CIRCLES denote local maxima in PFP relative to a trend line

# SSP: Shared Socioeconomic Pathway Scenarios Will Drive Upcoming IPCC Report

## Climate Model Input



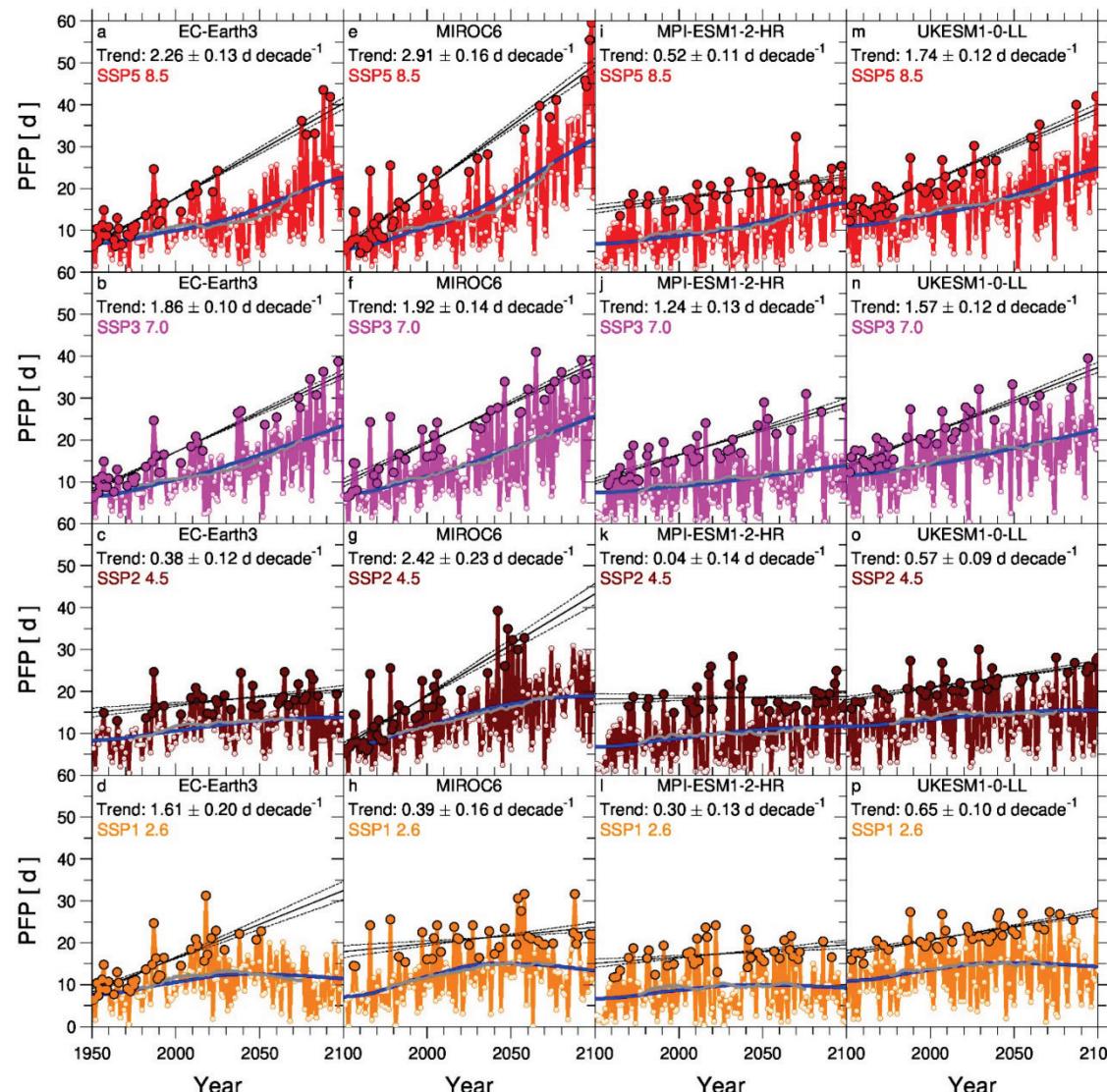
McBride et al., *Earth System Dynamics*, submitted, 2020

Number before dash represents base narrative and number after dash represents  $\text{W m}^{-2}$  RF of climate at end of century

# Tendency for Colder Arctic Winters Getting Colder Drive by Rising GHGs

## Climate Model Output

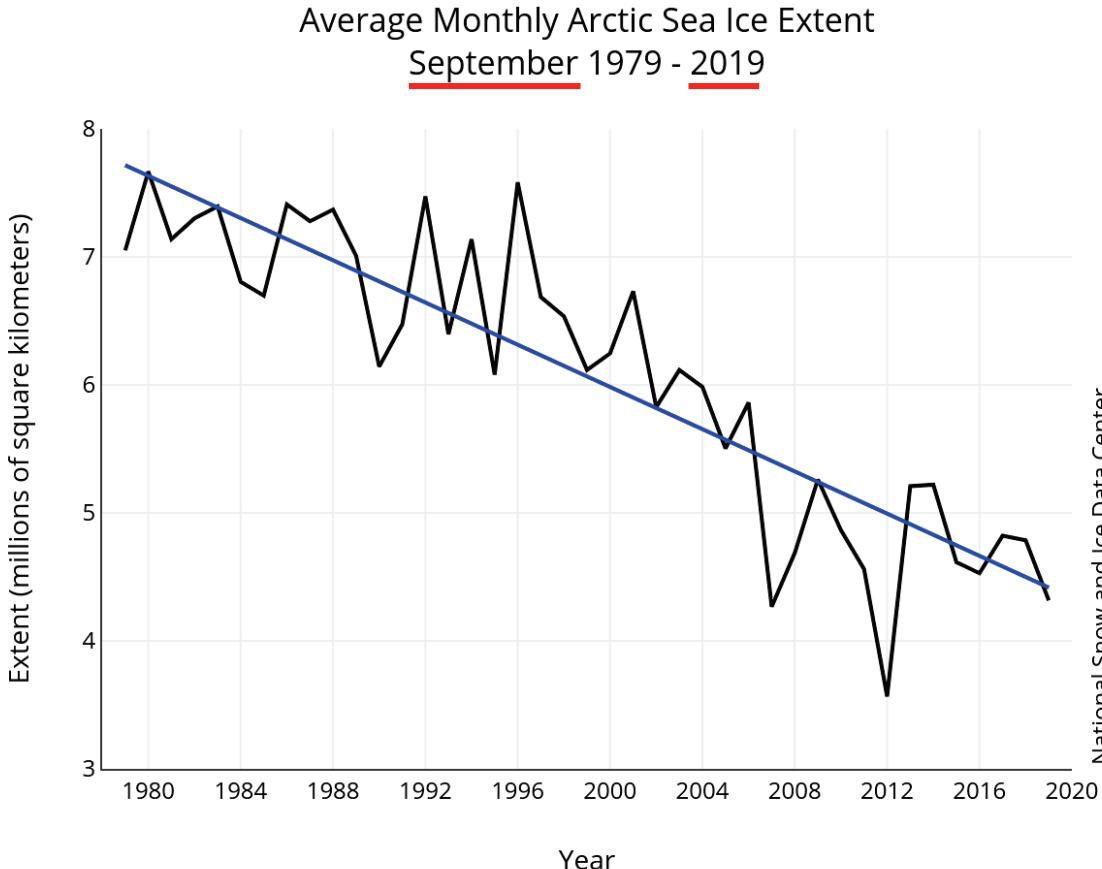
PFP is PSC Formation Potential



von der Gathen, *Nature Communications*, submitted, 2020

Lecture 16, Slide 78

# Arctic Sea-Ice: Canary of Climate Change



- Sea ice: ice overlying ocean
- Annual minimum occurs each September
- Decline of ~13% / decade over satellite era

<http://nsidc.org/arcticseaicenews/2019/10/>

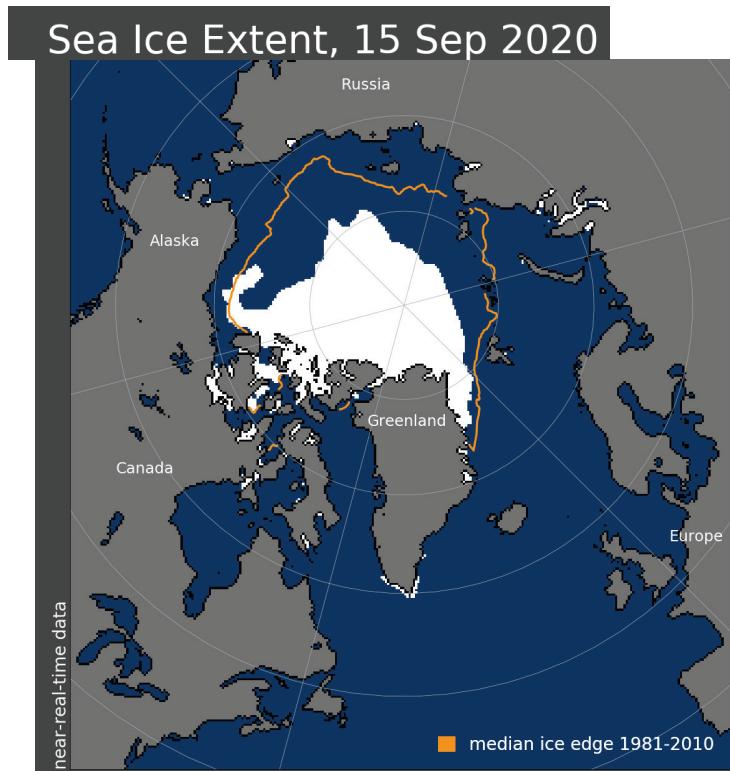
Lecture 8, Slide 28

# Arctic Sea Ice News & Analysis

September  
21, 2020

## Arctic sea ice decline stalls out at second lowest minimum

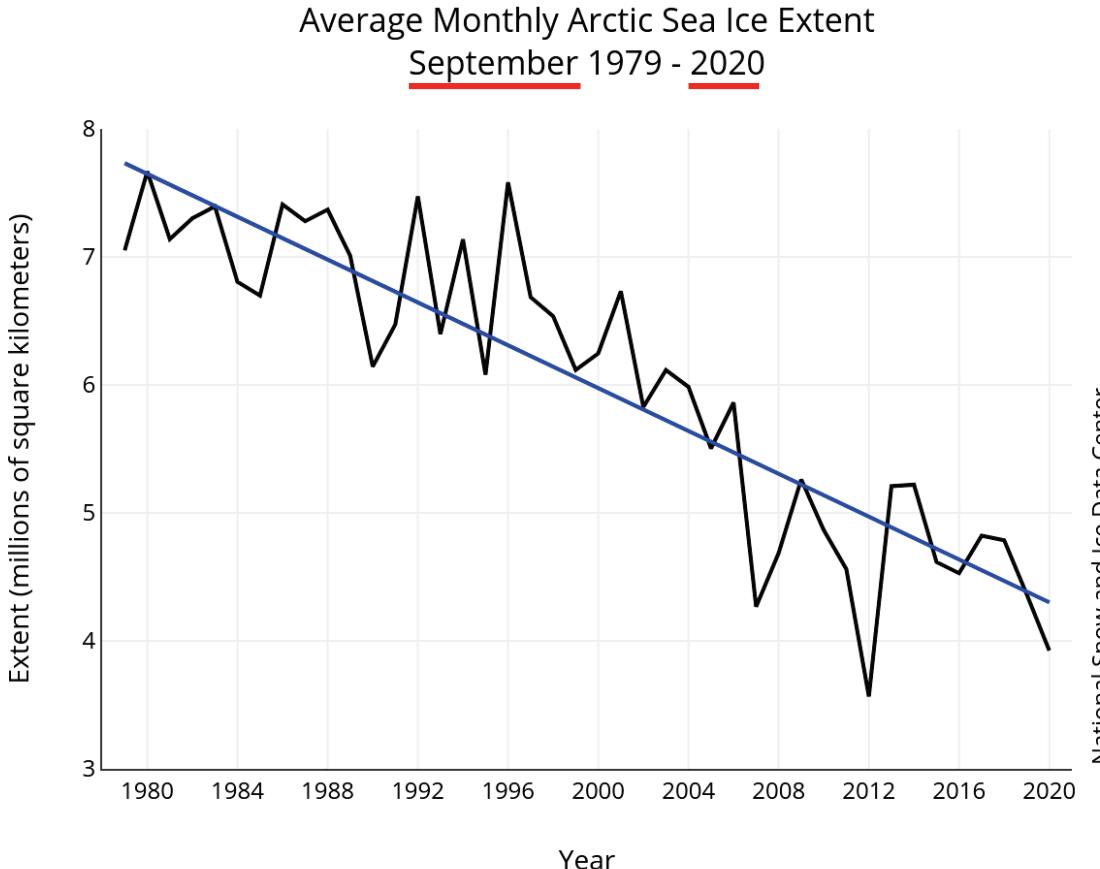
On September 15, Arctic sea ice likely reached its annual minimum extent of 3.74 million square kilometers (1.44 million square miles). The minimum ice extent is the second lowest in the 42-year-old satellite record, reinforcing the long-term downward trend in Arctic ice extent. Sea ice extent will now begin its seasonal increase through autumn and winter. In the Antarctic, sea ice extent is now well above average and within the range of the ten largest ice extents on record, underscoring its high year-to-year variability. The annual maximum for Antarctic sea ice typically occurs in late September or early October.



Lecture 8, Slide 29

<https://nsidc.org/arcticseaincnews/2020/09/arctic-sea-ice-decline-stalls-out-at-second-lowest-minimum/>

# Arctic Sea-Ice: Canary of Climate Change

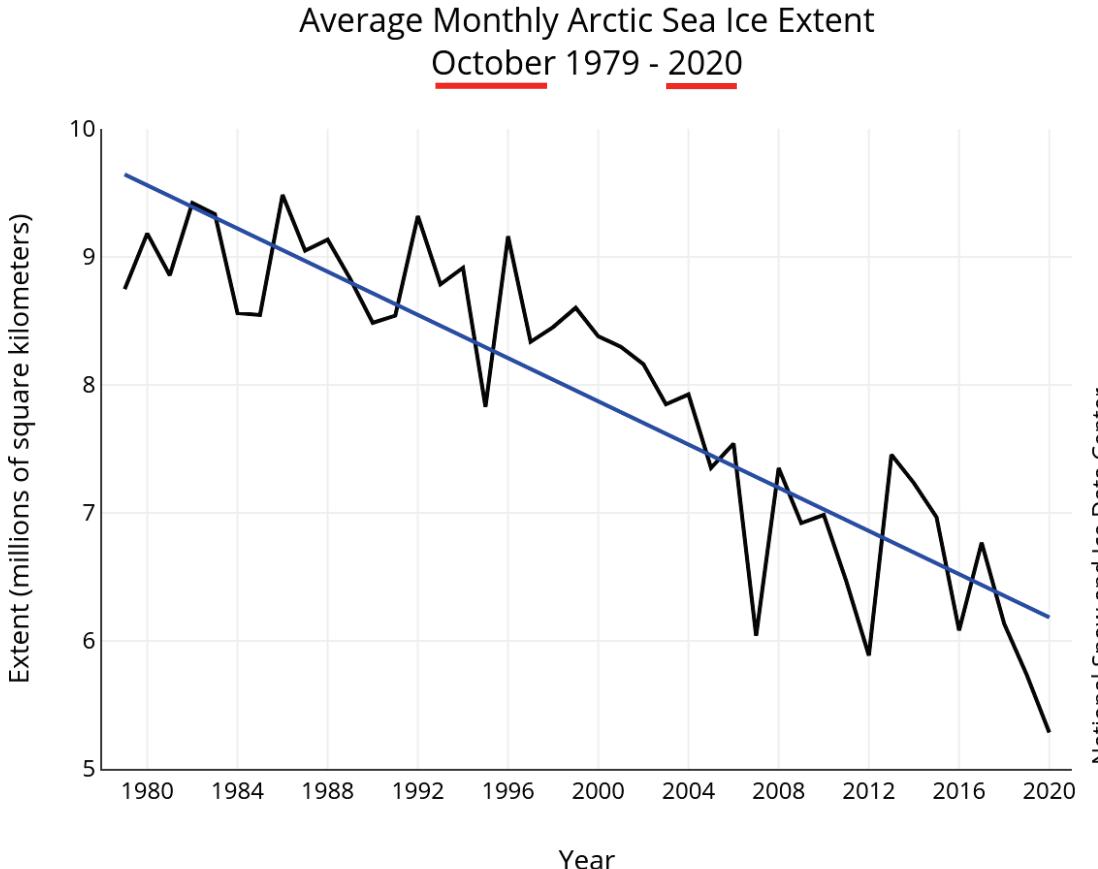


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<http://nsidc.org/arcticseaicenews/2020/10/>

Update to  
Lecture 8, Slide 28

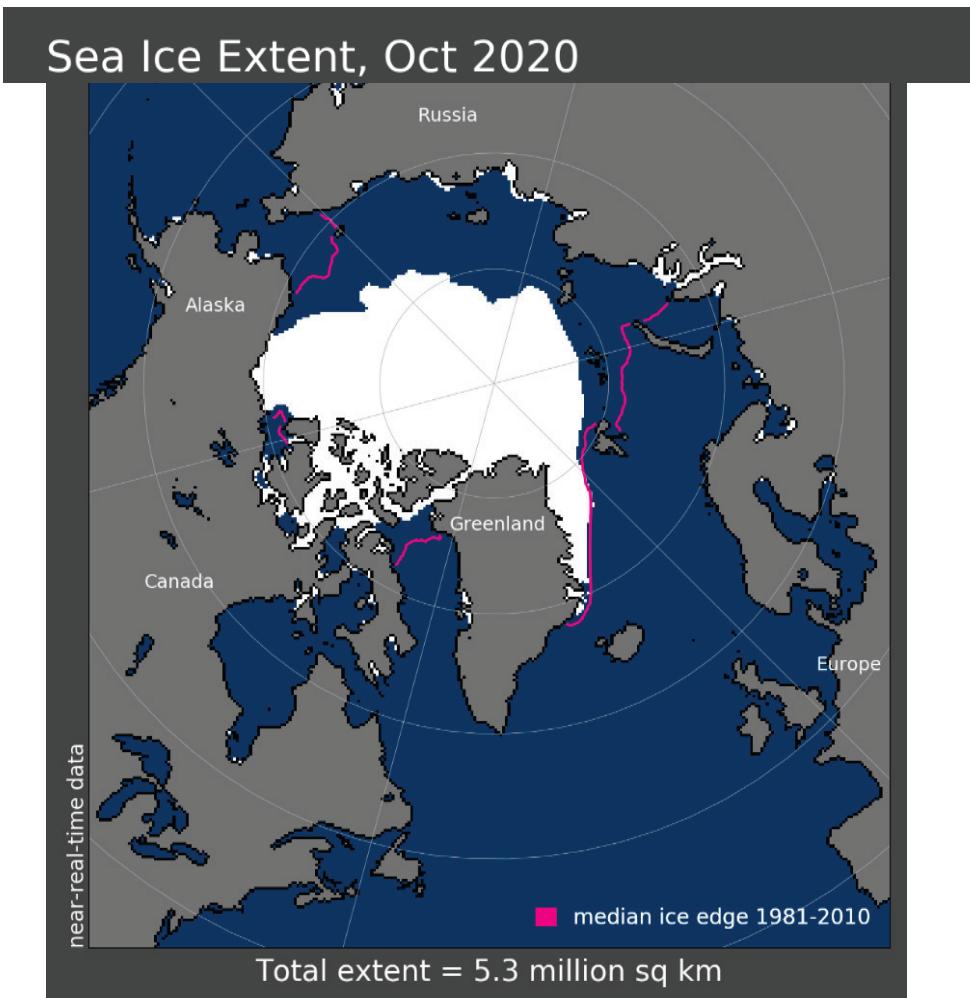
# Arctic Sea-Ice: Canary of Climate Change



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<http://nsidc.org/arcticseainews/files/2020/11/Figure3.png>

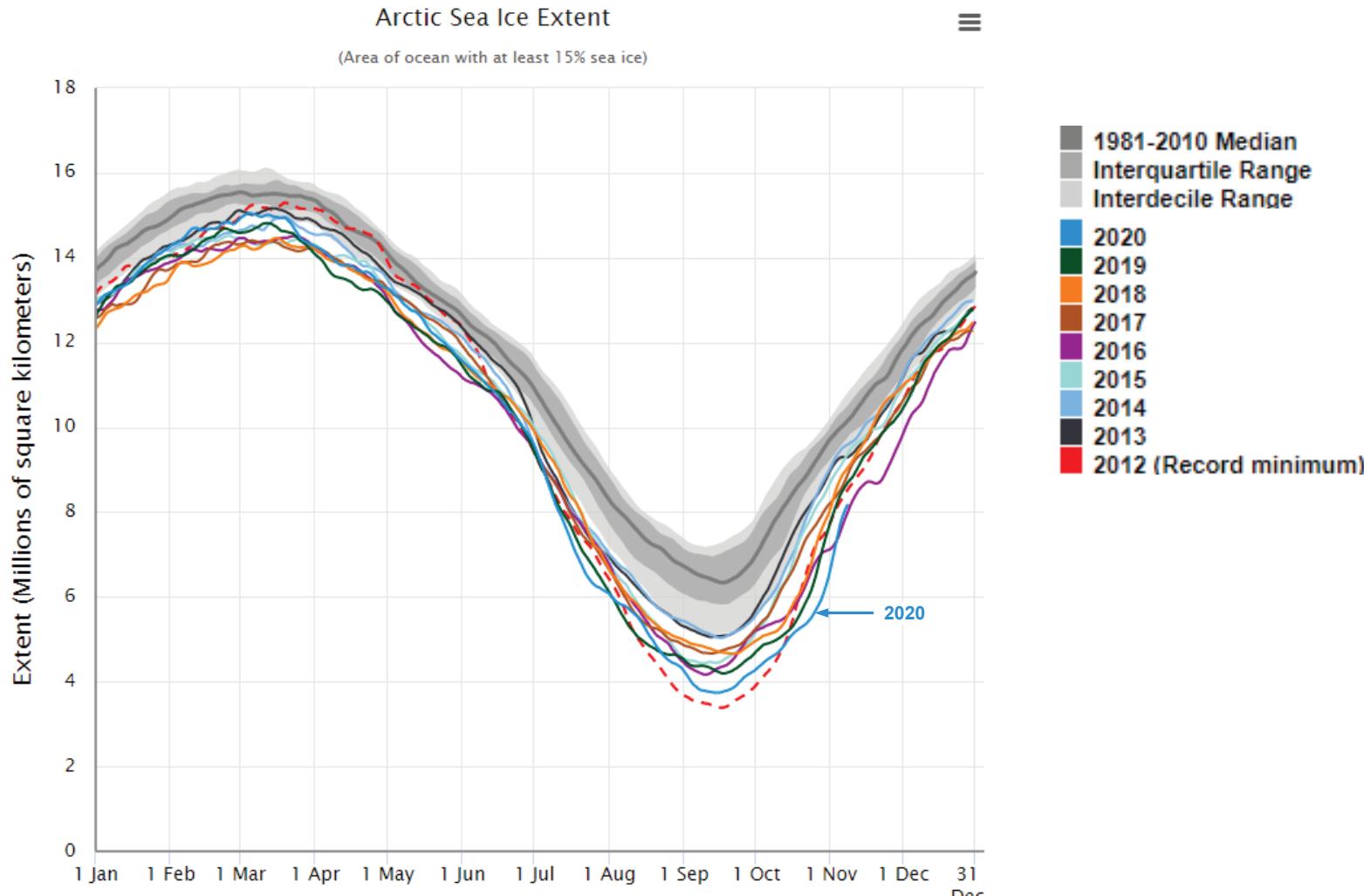
# Declining Arctic Sea Ice: Canary of Climate Change?



Arctic sea ice extent for October 2020 was 5.3 million square kilometers, the largest departure from average conditions seen in any month thus far in the satellite record.

<http://nsidc.org/arcticseaicenews>

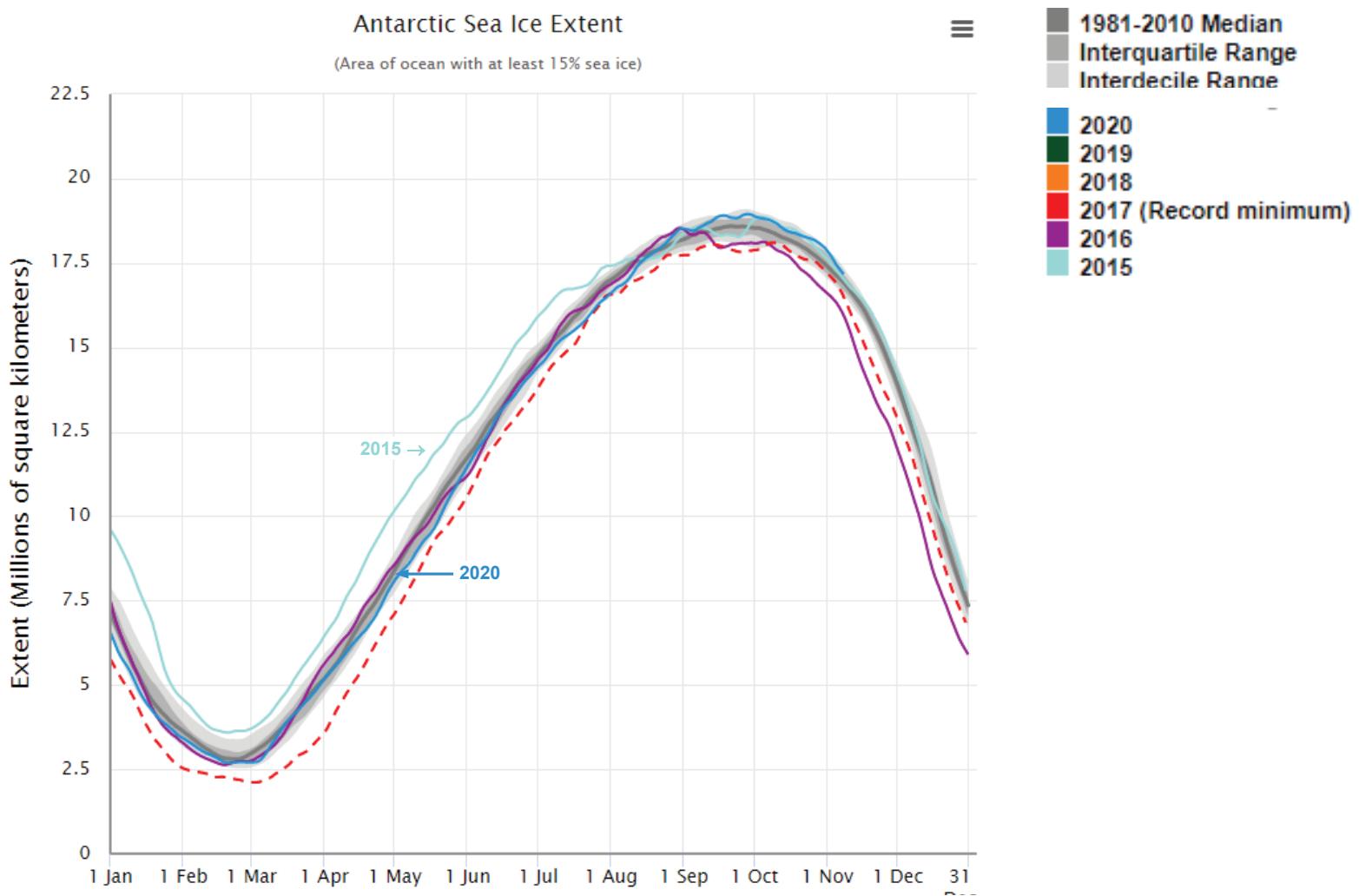
# Declining Arctic Sea Ice: Canary of Climate Change?



Don't need to use any heavy duty statistics to see the trend!

<https://nsidc.org/arcticseainews/charctic-interactive-sea-ice-graph>

# The Antarctic

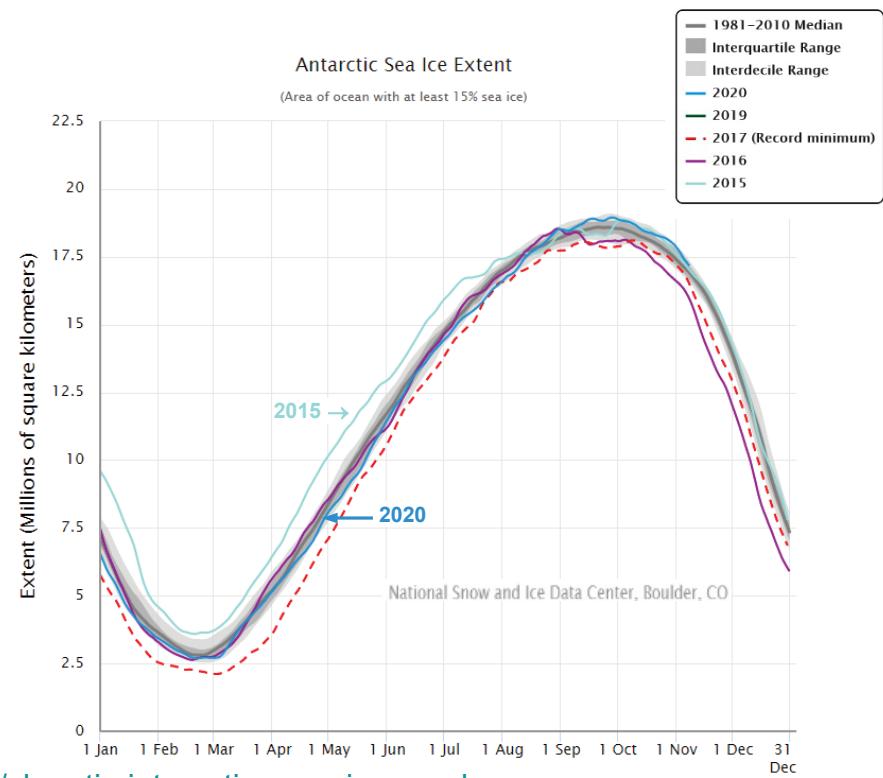
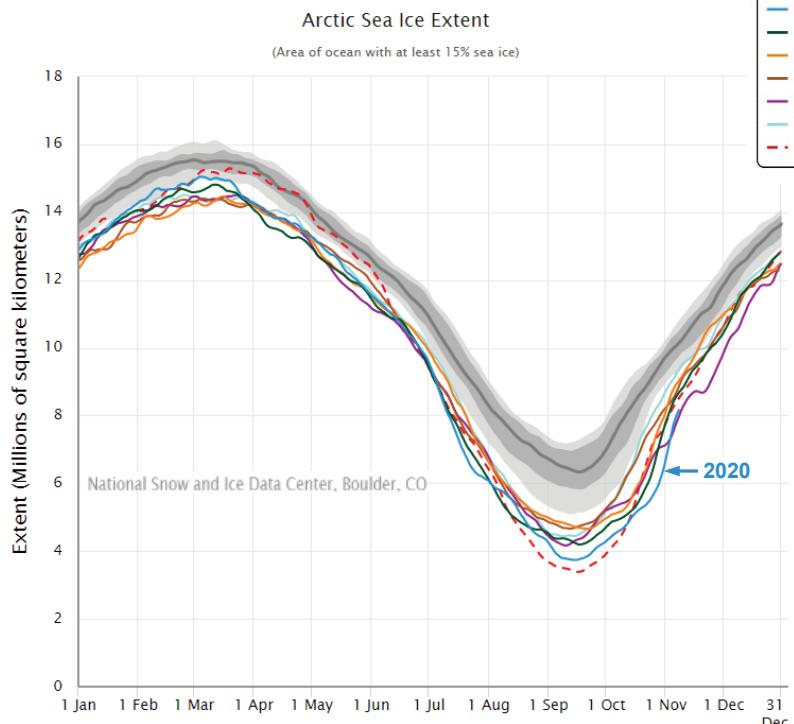


<https://nsidc.org/arcticseaincnews/charctic-interactive-sea-ice-graph>

# The Arctic and the Antarctic

Shaded region: 1981 to 2010 mean  $\pm 2\sigma$

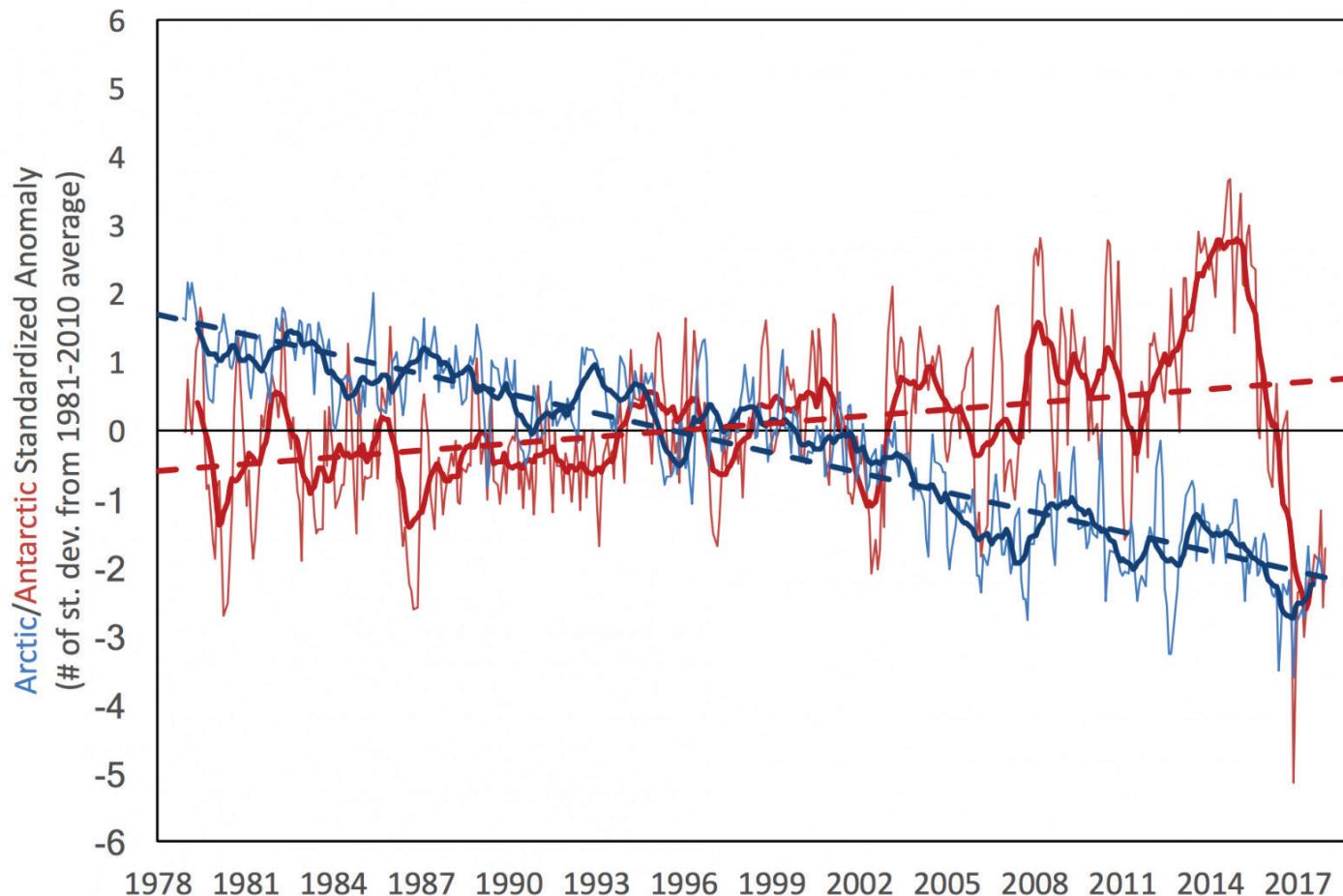
Colors: Individual years since 2017  
(plus 2012 for Arctic, since this  
was year of record minimum in NH)



<http://nsidc.org/arcticseainews/charctic-interactive-sea-ice-graph>

## Arctic and Antarctic Standardized Anomaly and Trend

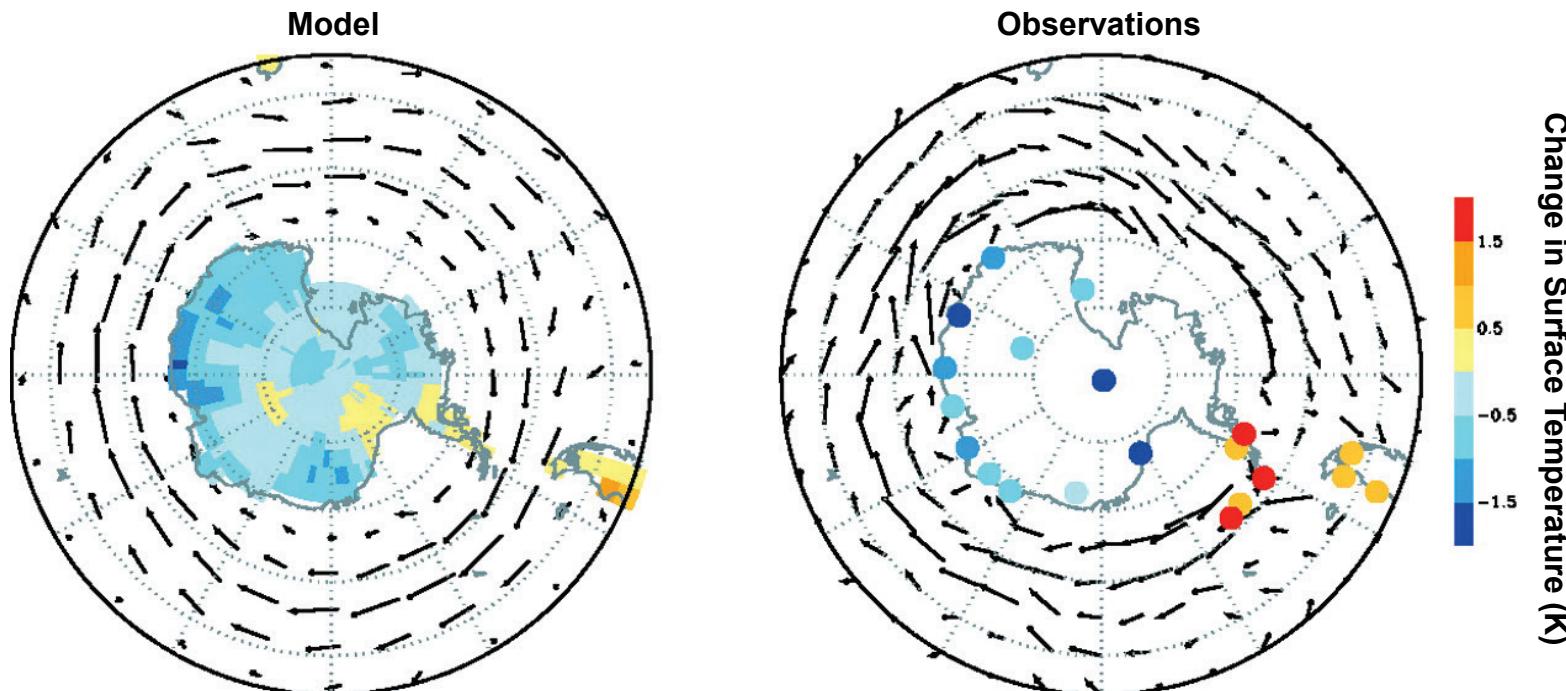
Nov. 1978 - Dec. 2017



Changes in the extent of Arctic (blue) and Antarctic sea ice (red) from November 1978 to December 2017, relative to a 1981-2010 baseline. Thick lines show changes to the yearly average and thin lines show changes to the monthly anomalies. Source: [National Snow and Ice Data Center, University of Colorado, Boulder](#)

[https://nsidc.org/cryosphere/sotc/sea\\_ice.html](https://nsidc.org/cryosphere/sotc/sea_ice.html)

# The Ozone Hole may have shielded the Antarctic surface from warming!



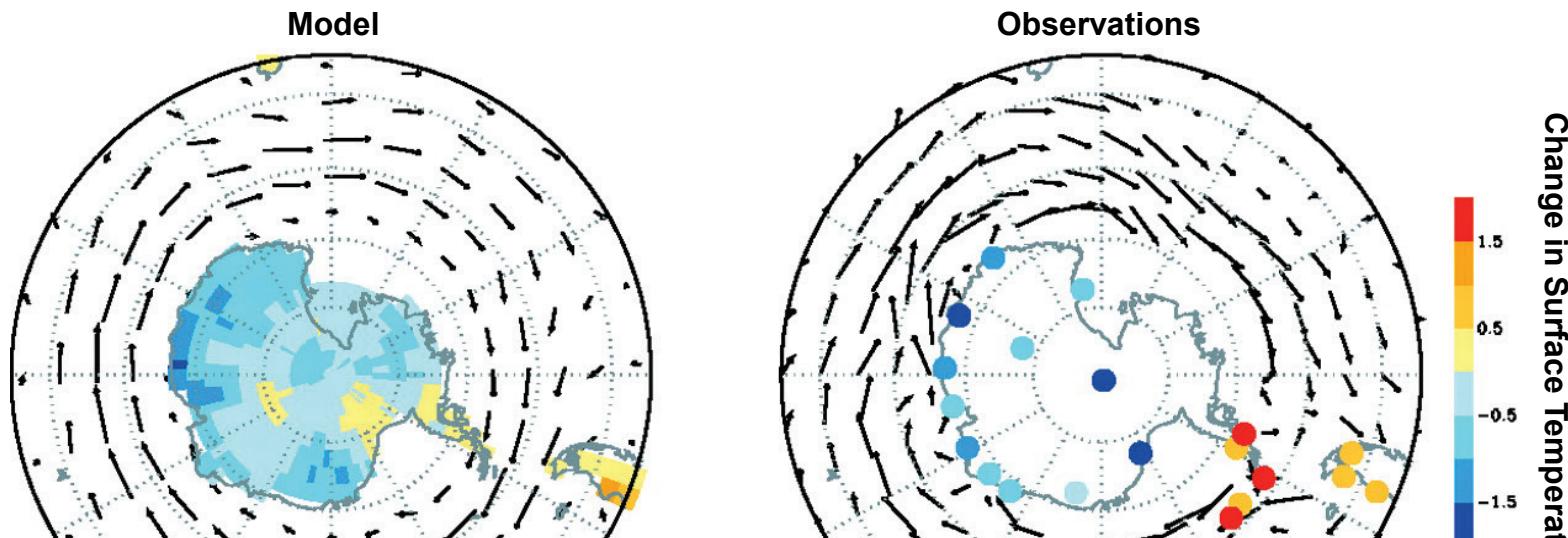
Simulated and observed changes in surface temperature (K) and wind speed, 1969 to 2000, averaged over December to May. The longest wind vector corresponds to 4 m/s.

Gillett and Thompson, *Science*, 2003

As ozone depletion occurs:

The positive phase of the southern annular mode (SAM) increases, causing Antarctic surface westerlies to intensify, resulting in cooling of Antarctic continent

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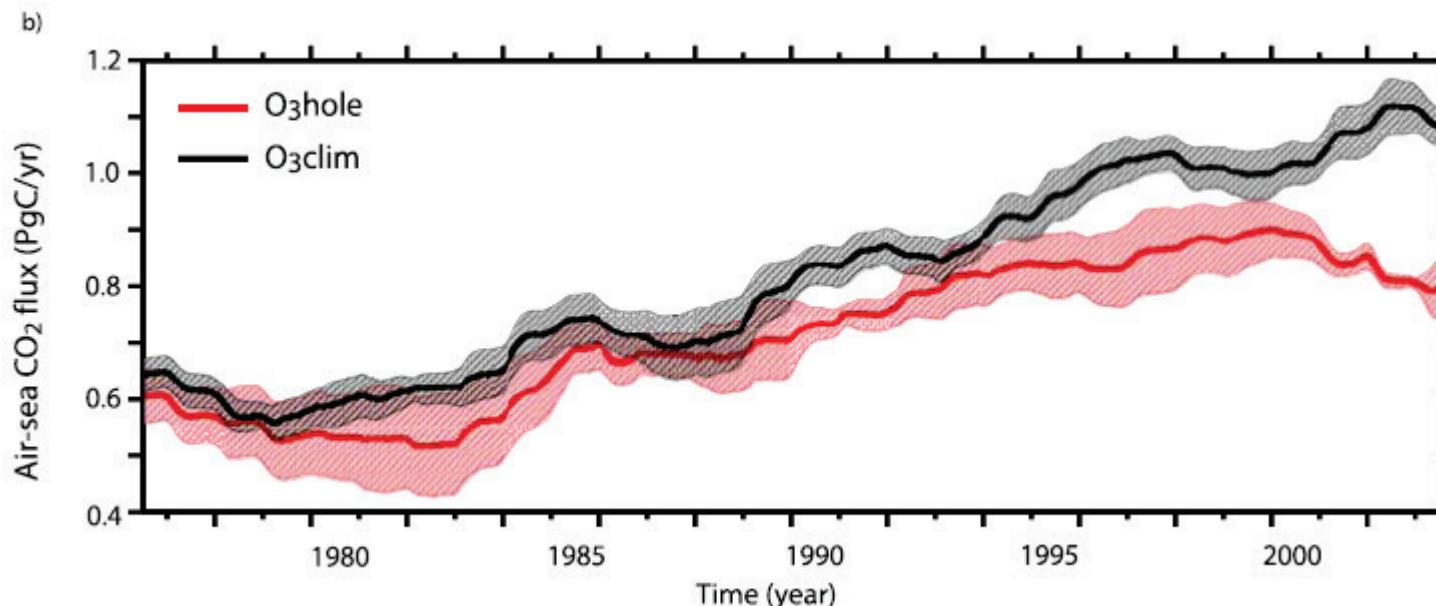
**SAM: difference in zonal mean sea-level pressure between 40°S and 65°S.**  
The pattern associated with SAM is a nearly annular pattern with a large low pressure anomaly centered on the South Pole and a ring of high pressure anomalies at mid-latitudes. The SAM effects storm tracks, precipitation patterns, etc.

[http://www.climate.be/textbook/chapter5\\_node6.html](http://www.climate.be/textbook/chapter5_node6.html)

As ozone depletion occurs:

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# The Ozone Hole may have lead to increased ventilation of CO<sub>2</sub> from southern ocean



(b) Integrated air to sea CO<sub>2</sub> flux (south of 40°S) showing stratospheric ozone depletion (O<sub>3</sub>hole) significantly reduces CO<sub>2</sub> uptake (relative to O<sub>3</sub>clim), and is strongly correlated with changes in ΔpCO<sub>2</sub>.

Lenton *et al.*, GRL, 2009

As ozone depletion occurs:

The positive phase of the southern annular mode (SAM) increases, causing Antarctic surface westerlies to intensify, resulting in increased ventilation of CO<sub>2</sub> from southern ocean

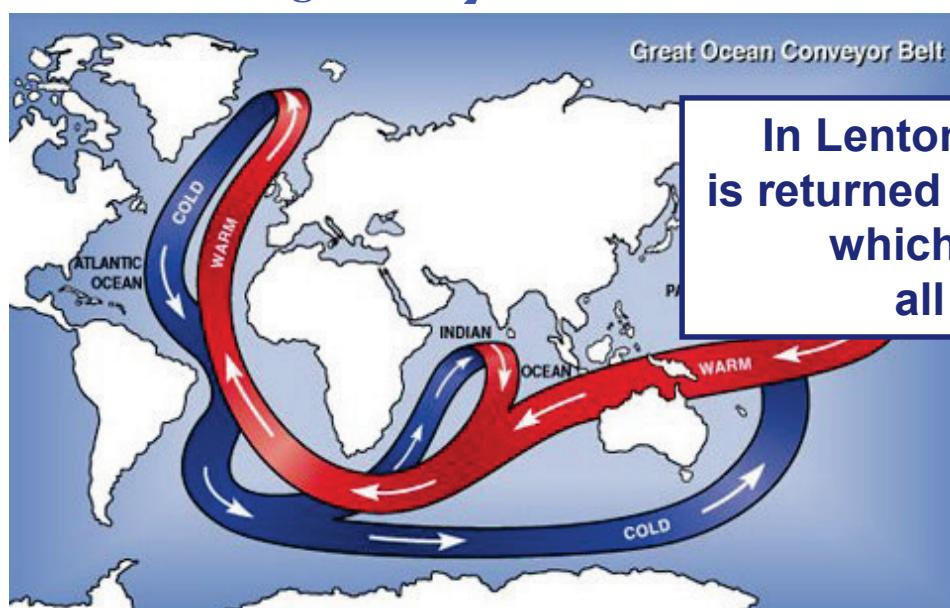
# Uptake of Atmospheric CO<sub>2</sub> by Oceans

## – Solubility Pump:

- a) More CO<sub>2</sub> can dissolve in cold polar waters than in warm equatorial waters. As major ocean currents (e.g. the Gulf Stream) move waters from tropics to the poles, they are cooled and take up atmospheric CO<sub>2</sub>
- b) Deep water forms at high latitude. *As deep water sinks, ocean carbon ( $\Sigma CO_2$ ) accumulated at the surface is moved to the deep ocean interior.*

## – Biological Pump:

- a) Ocean biology limited by availability of nutrients such as NO<sub>3</sub><sup>-</sup>, PO<sub>4</sub><sup>-</sup>, and Fe<sup>2+</sup> & Fe<sup>3+</sup>. Ocean biology is never carbon limited.
- b) Detrital material “rains” from surface to deep waters, *contributing to higher CO<sub>2</sub> in intermediate and deep waters*



[http://science.nasa.gov/headlines/y2004/05mar\\_arctic.htm](http://science.nasa.gov/headlines/y2004/05mar_arctic.htm)

# Chemistry Climate Coupling

CCMs (chemistry climate models): developed to quantify impacts of climate change on stratospheric ozone and impacts of ozone depletion/recovery on climate:

As GHGs rise:

1. Brewer-Dobson circulation predicted to accelerate leading to:
  - a) less ozone in tropical lower stratosphere (“permanent depletion”)
  - b) more ozone in mid-latitude lower stratosphere (“super recovery”)
2. Upper stratosphere cools, slowing down rate limiting steps for ozone loss and therefore leading to “super recovery”
3. Eventually, CH<sub>4</sub> and N<sub>2</sub>O will drive future levels of ozone

Data analysis suggests “coldest Arctic winters getting colder”:

1. Possibly due to rising GHGs
2. Not represented well by CCMs

As Antarctic ozone depletion had occurred:

The positive phase of the southern annular mode (SAM) increases, causing Antarctic surface westerlies to intensify, resulting in:

1. Cooling of Antarctic continent (good for sea-level)
2. Increased ventilation of CO<sub>2</sub> from southern ocean (bad for climate)

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As Antarctic ozone recovery occurs:

The positive phase of the southern annular mode (SAM) may decline, causing Antarctic surface westerlies to weaken, resulting in:

1. Warming of Antarctic continent (bad for sea-level)
2. Decreased ventilation of CO<sub>2</sub> from southern ocean (good for climate)