

Review for First Exam

AOSC / CHEM 433 & AOSC / CHEM 633

Ross Salawitch

Class Web Sites:

<http://www2.atmos.umd.edu/~rjs/class/fall2020>

<https://myelms.umd.edu/courses/1291919>

Review A
6 October 2020

Announcements: Class

Original announcement:

Also, the first exam which will be on-line, open book, open web, that I hope you will take on 8 Oct (you'll have 75 mins time to take either 7 Oct, 8 Oct, or 9 Oct) will cover material up to Lecture 8 (today's lecture).

However, please note that:

- a) the exam will focus on a series of questions, about one per lecture, that you can only answer properly in the limited time IF you are already familiar with the contents of each lecture;
- b) there will be no calculations on the exam so, in this sense, the exam will be qualitative rather than quantitative
- c) if you have been doing all of the readings, answering the ATs based on a comprehensive understanding of the readings, and retaining knowledge from the readings and exams, as solidified by consistently high scores of the learning outcome quizzes, then you'll be in great shape for the first exam. On the other hand, if you have been skimming the readings, doing the bare minimum to answer the ATs, and not completing the learning outcome quizzes, you will need to impart greater effort to prepare for the exam, in order to do well. The exam will open more than a week from this moment in time ... so there is still time to prepare.

Announcements: Class

Modified announcement:

Also, the first exam which will be on-line, open book, open web, [*that everyone will take during normal class hour on Thurs, 8 Oct, from 2:00 to 3:15 pm.*](#)

However, please note that:

- a) the exam will focus on a series of questions, about one per lecture, that you can only answer properly in the limited time IF you are already familiar with the contents of each lecture;
- b) there will be no calculations on the exam so, in this sense, the exam will be qualitative rather than quantitative
- c) if you have been doing all of the readings, answering the ATs based on a comprehensive understanding of the readings, and retaining knowledge from the readings and exams, as solidified by consistently high scores of the learning outcome quizzes, then you'll be in great shape for the first exam. On the other hand, if you have been skimming the readings, doing the bare minimum to answer the ATs, and not completing the learning outcome quizzes, you will need to impart greater effort to prepare for the exam, in order to do well. The exam will open more than a week from this moment in time ... so there is still time to prepare.

Announcements: Class

Modified announcement:

Also, the first exam which will be on-line, open book, open web, that everyone will take during normal class hour on Thurs, 8 Oct, from 2:00 to 3:15 pm.

Details:

- Exam will open at 2:02 pm on Thurs, 8 Oct
- Exam will close at 3:17 pm on Thurs, 8 Oct
- Everyone is asked to connect to the class Zoom while conducting the exam, and have video running solely so that I can see you are “alone”
- Welcome to use any resource (readings, lectures, notes, etc) other than communicating with another person
- Exam will consist of 9 essay questions:
 - Students will answer 6 of the 9 questions
 - Questions will cover Lectures 1 to 8
 - Expected that each answer will require about 10 minutes of thought & writing
 - Questions should be sent to me through chat: if any question requires clarification, this will be announced via Zoom
 - One of the questions, **Question 9**, will be required to be answered by students in 633
 - When the exam is complete, you are asked to write the word “BLANK” in the text box of the 3 questions you do not answer
 - If you have internet connectivity issues during the exam, please text me at 626-487-5643 and include a cell phone image showing “the problem”

Announcements: Class

Modified announcement:

Also, the first exam which will be on-line, open book, open web, that everyone will take during normal class hour on Thurs, 8 Oct, from 2:00 to 3:15 pm.

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 - One of the questions, **Question 9**, will be required to be answered by students in 633
 - When the exam is complete, you are asked to write the word “BLANK” in the text box of the 3 questions you do not answer
 - **If you have an exceptional circumstance that requires taking this exam at a different time, please let me know ASAP**

Announcements: Class

Roll call to state you understand exam will be taken on ELMS during regular class time on Thurs, with your Zoom video running (fine to mute), and that you should text Ross at 626-487-5643 if you have internet connection issues, you should send a chat to Ross if you have questions during the exam, and that the exam is open book, open web, open everything except no contact with another person during the exam.



Student Name
Emma B AOSC433-0101
Nancy C CHEM433-0101
Ciara D AOSC433-0101
Ehiremen E CHEM633-0101
Endre F CHEM633-0101
Heather F AOSC433-0101
Nicholas F CHEM633-0101
HyunGee H CHEM433-0101
Michael H AOSC433-0101
Rachel K AOSC433-0101
Madeline L CHEM633-0101
Kishalay M CHEM633-0101



Student Name
Kotiba M AOSC633-0101
Jeffrey M CHEM433-0101
Malgorzata P CHEM433-0101
Michael R CHEM633-0101
Akanksha S AOSC633-0101
Stuart S CHEM633-0101
Nathaniel S AOSC433-0101
William Jay S AOSC433-0101
Pierce V CHEM633-0101
Alice W CHEM433-0101
Cindy X CHEM433-0101
Chencheng Z CHEM433-0101

Announcements: Outside of Class

1) Thurs, 8 Oct : AOSC Weekly Seminar (3:30 pm)

Dr. Jiwen Fan

Physical factors impacting convective storms and weather hazards

Deep convective clouds play a crucial role in atmospheric circulation, energy, and water cycle of our climate system. The extreme form of such storms produces weather hazards such as large hail, damaging winds and/or tornadoes, and torrential rainfall, causing significant property damages and economic losses. There is a large gap in our fundamental understanding of environmental factors impacting storm intensity, precipitation, and associated hazards. In particular, how aerosols from various sources modify these storms is still controversial. In this talk, I will present our efforts in understanding physical factors that impact storm intensity, extreme precipitation, and hailstones associated with:

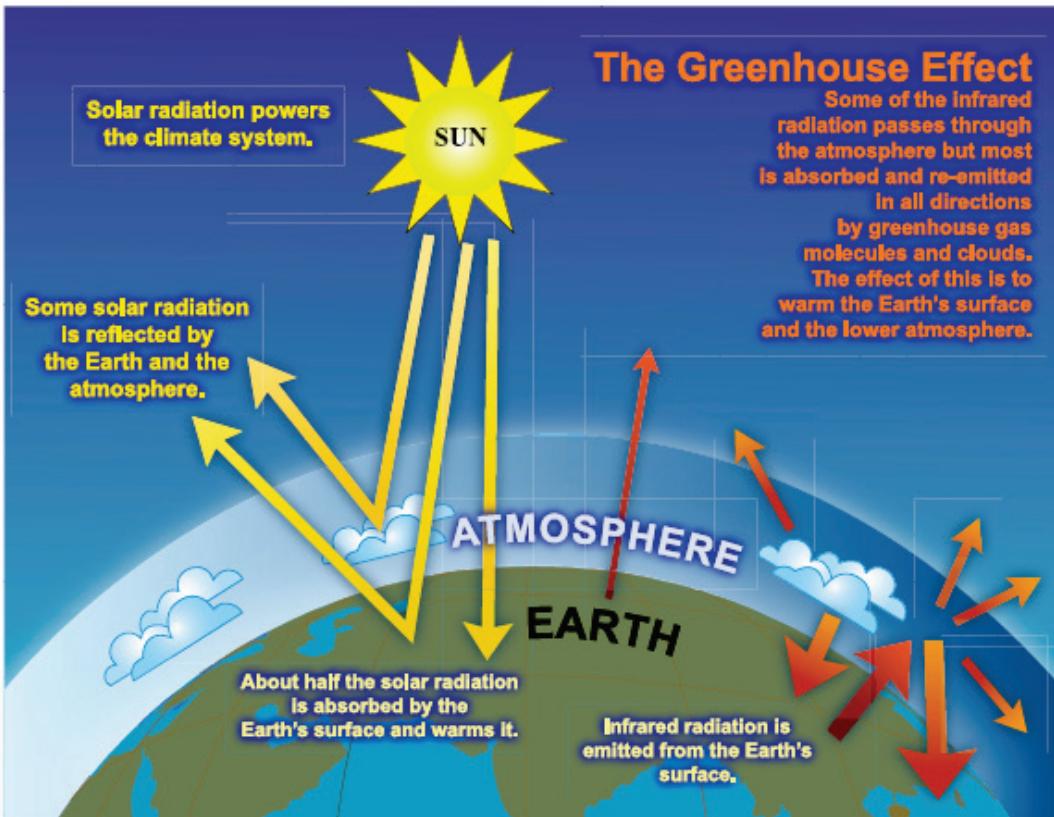
- (1) land cover and anthropogenic aerosol changes and
- (2) anthropogenic warming

I will focus on the understandings mainly gained from process-level studies with both advanced observations and high-resolution model simulations. The challenges in observing and modeling such deep convective clouds and their interactions with environmental factors will be discussed.

<https://aosc.umd.edu/seminars/department-seminar>

Email Joseph Knisely at jknisely@umd.edu for Zoom connection info

Greenhouse Effect

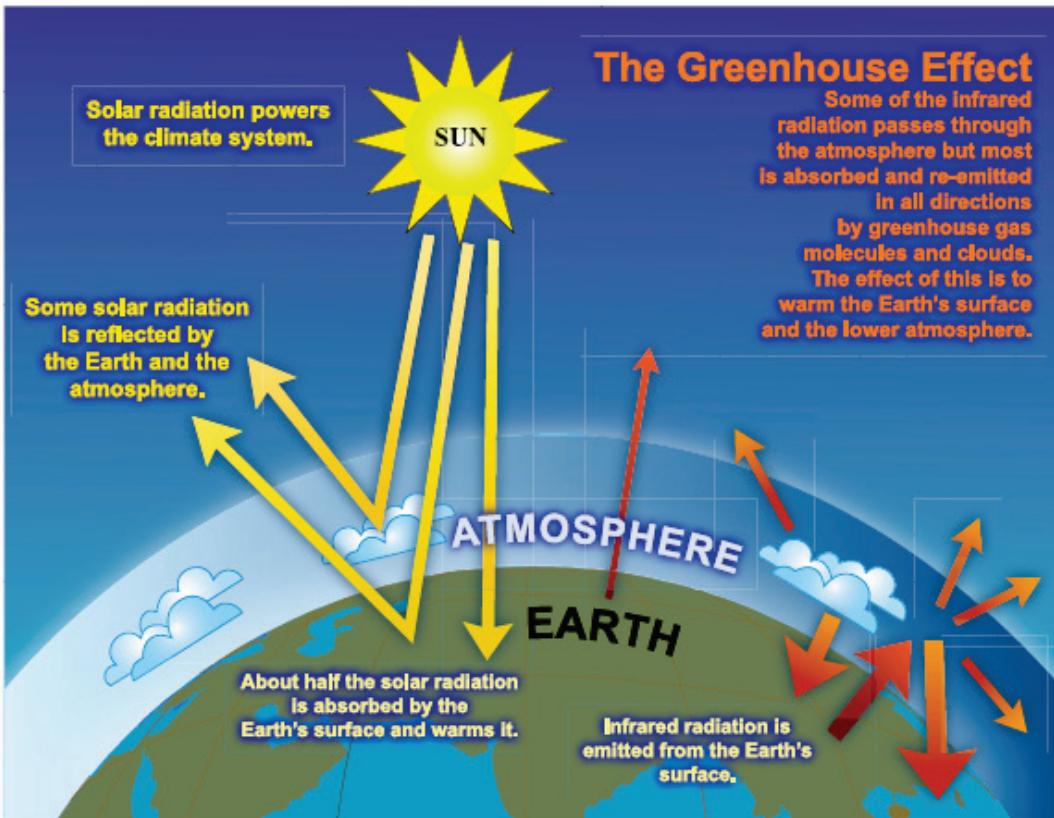


What is the most important greenhouse gas (GHG) ?

FAQ 1.3, Figure 1. An idealised model of the natural greenhouse effect. See text for explanation.

Question 1.3, IPCC, 2007

Greenhouse Effect

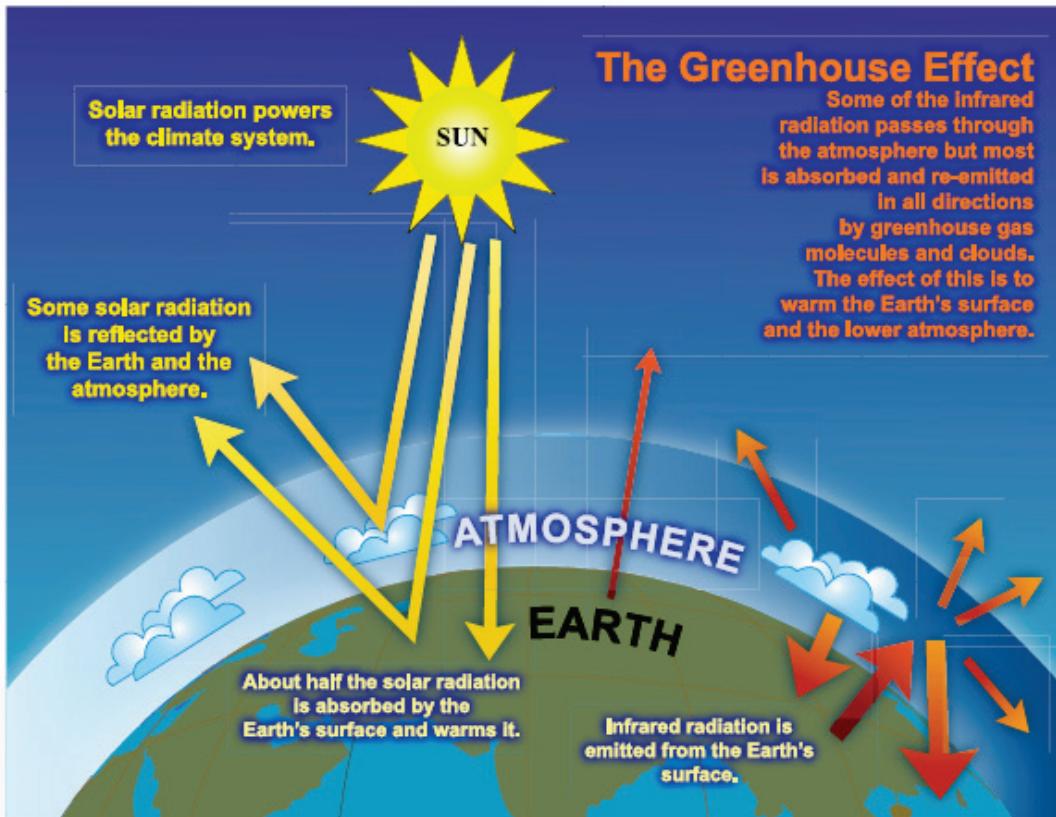


What is the most important anthropogenic greenhouse gas (GHG) ?

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Question 1.3, IPCC, 2007

Greenhouse Effect



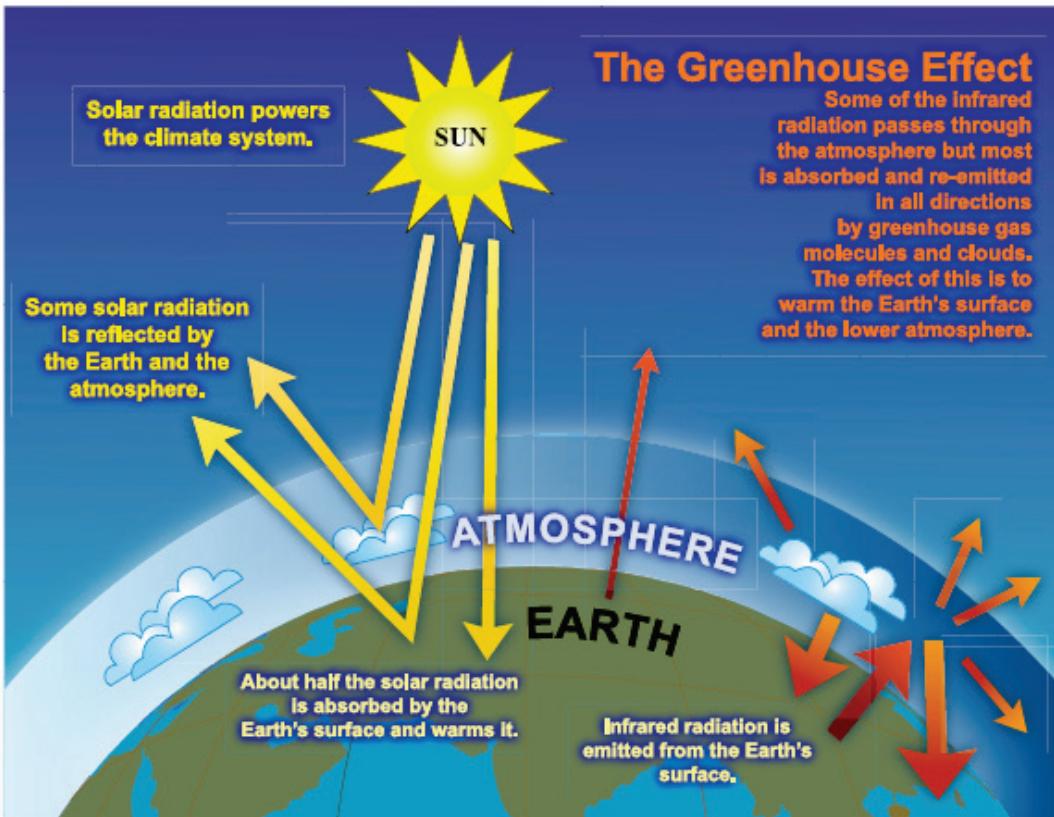
FAQ 1.3, Figure 1. An idealised model of the natural greenhouse effect. See text for explanation.

Question 1.3, IPCC, 2007

What is the most important anthropogenic greenhouse gas (GHG) ?

Second most important ?

Greenhouse Effect



FAQ 1.3, Figure 1. An idealised model of the natural greenhouse effect. See text for explanation.

Question 1.3, IPCC, 2007

What is the most important anthropogenic greenhouse gas (GHG) ?

Second most important ?

Third ?

Radiative Forcing of Climate, 1750 to 2011

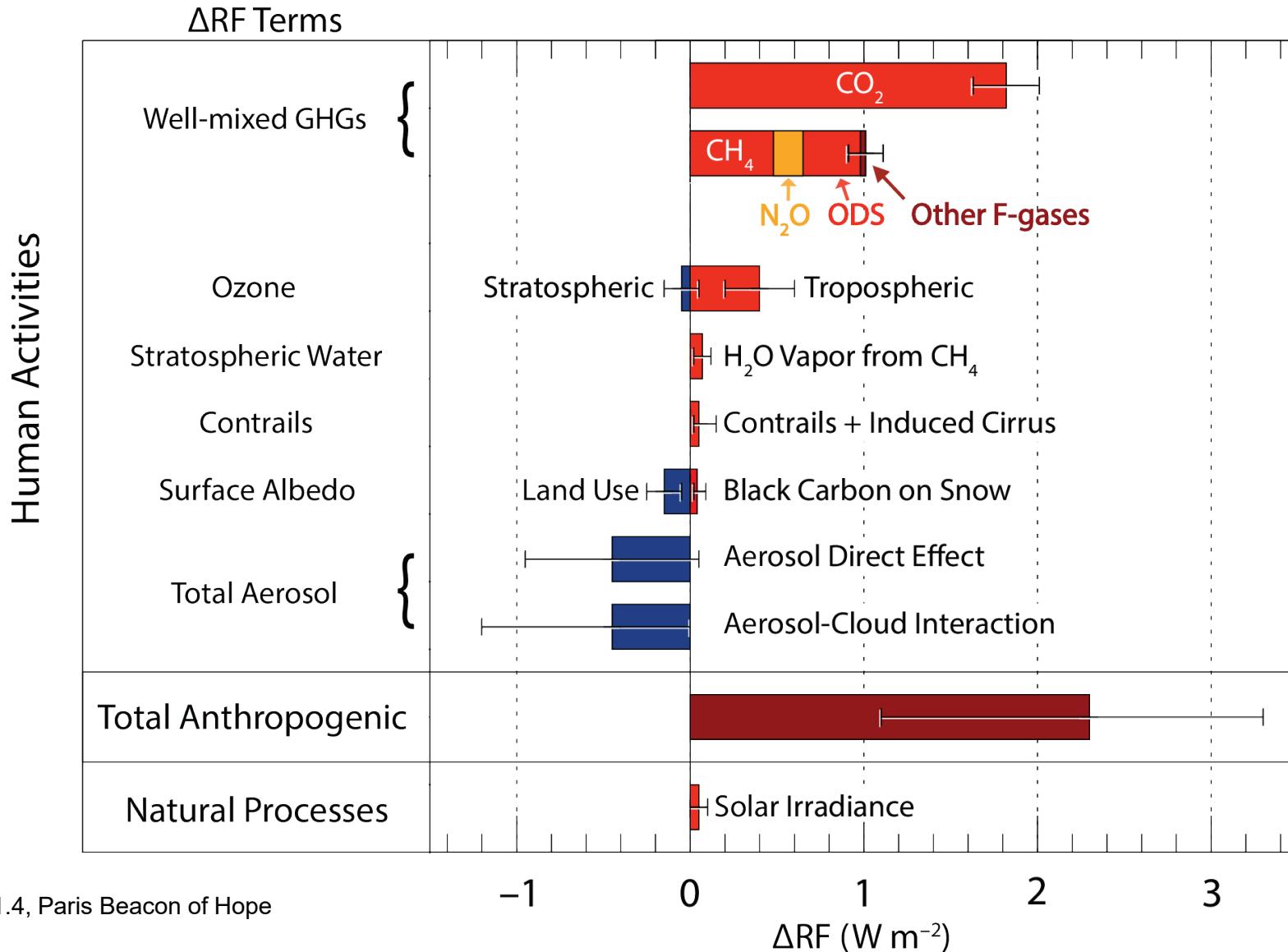
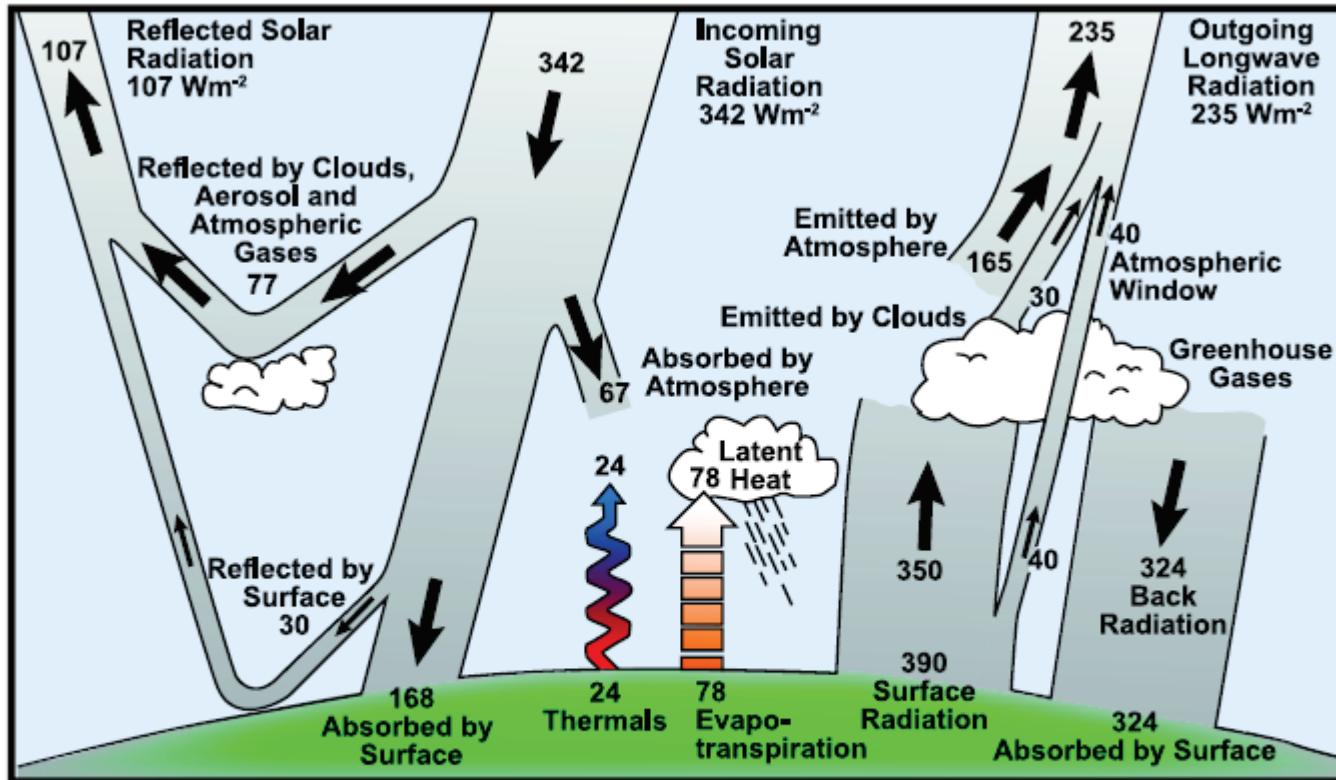


Figure 1.4, Paris Beacon of Hope

Radiative Forcing



FAQ 1.1, Figure 1. Estimate of the Earth's annual and global mean energy balance. Over the long term, the amount of incoming solar radiation absorbed by the Earth and atmosphere is balanced by the Earth and atmosphere releasing the same amount of outgoing longwave radiation. About half of the incoming solar radiation is absorbed by the Earth's surface. This energy is transferred to the atmosphere by warming the air in contact with the surface (thermals), by evapotranspiration and by longwave radiation that is absorbed by clouds and greenhouse gases. The atmosphere in turn radiates longwave energy back to Earth as well as out to space. Source: Kiehl and Trenberth (1997).

Question 1.1, IPCC, 2007

Radiative Forcing of Climate is Change in Energy
 reaching the lower atmosphere (surface to tropopause) as GHGs rise.
 “Back Radiation” is most important term.

Radiative Forcing of Climate, 1750 to 2011

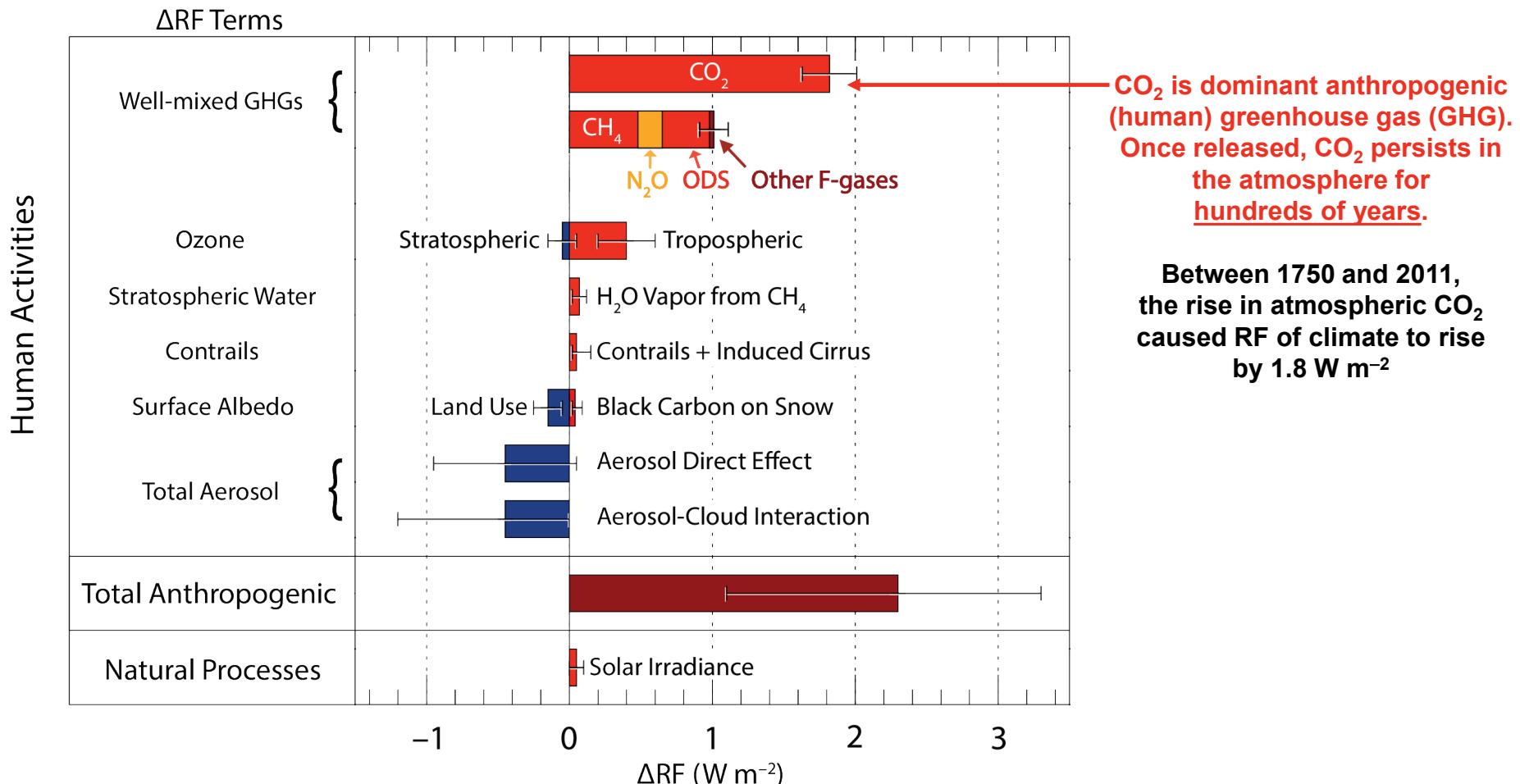


Figure 1.4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

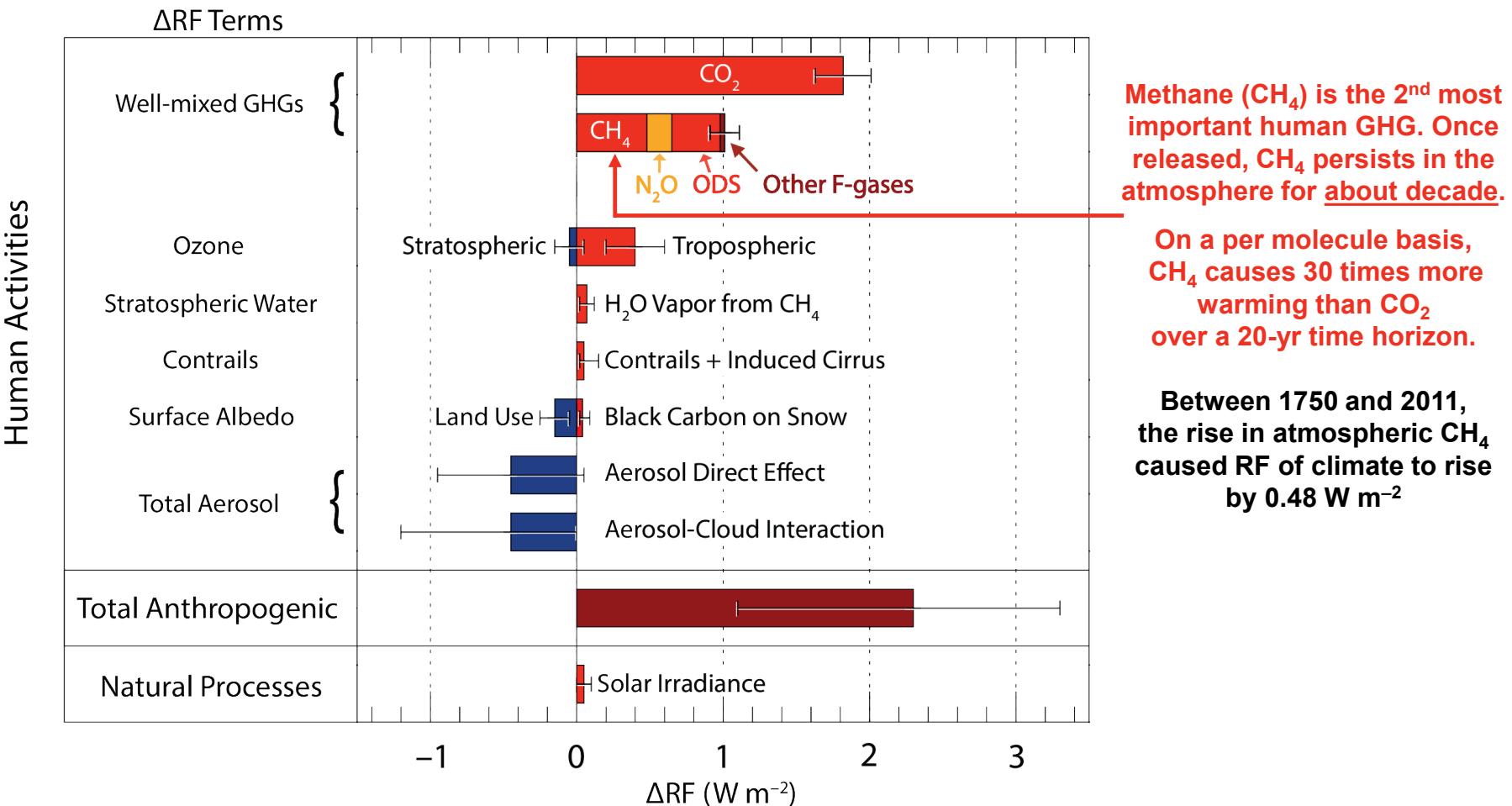


Figure 1.4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

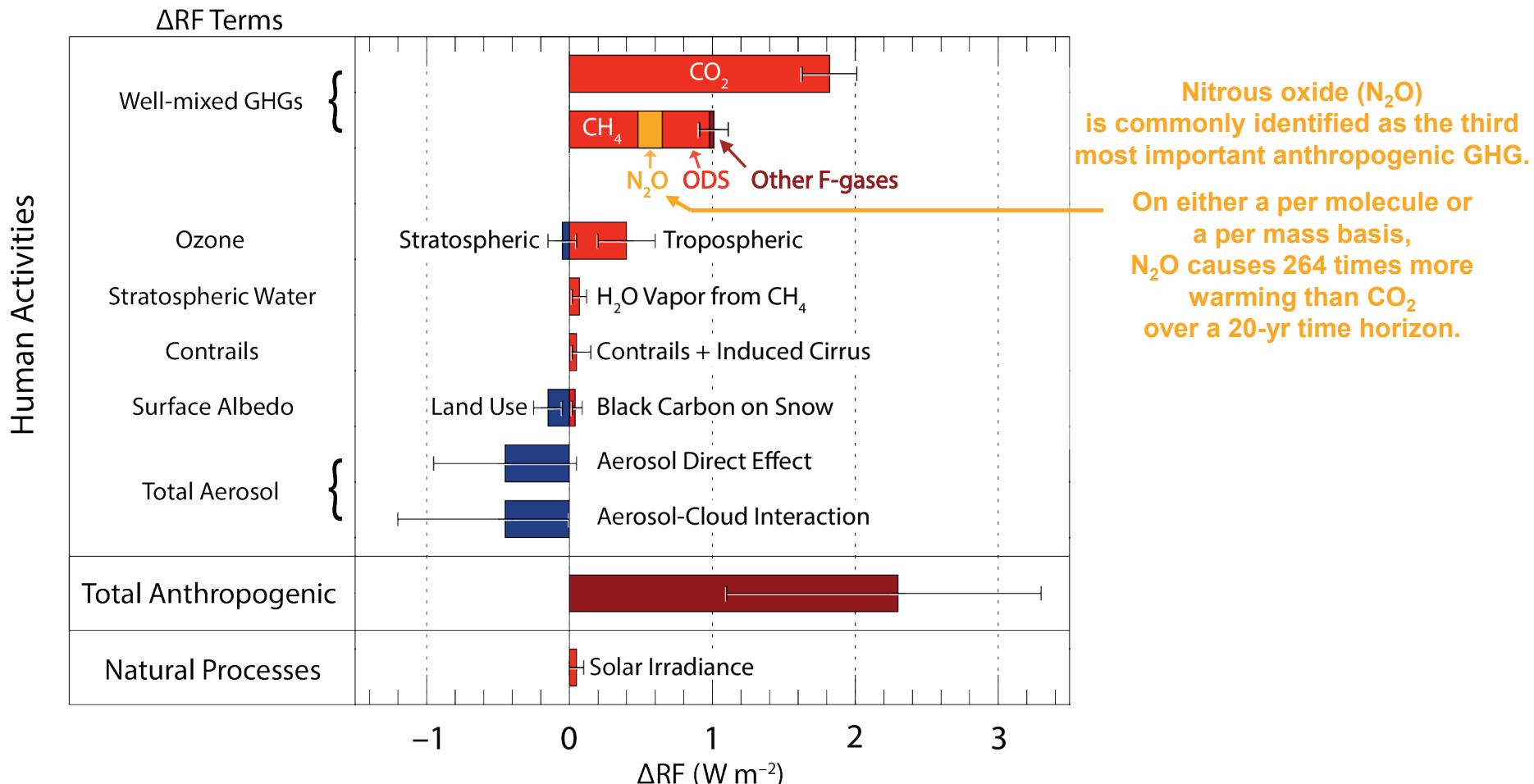


Figure 1-4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

Human Activities

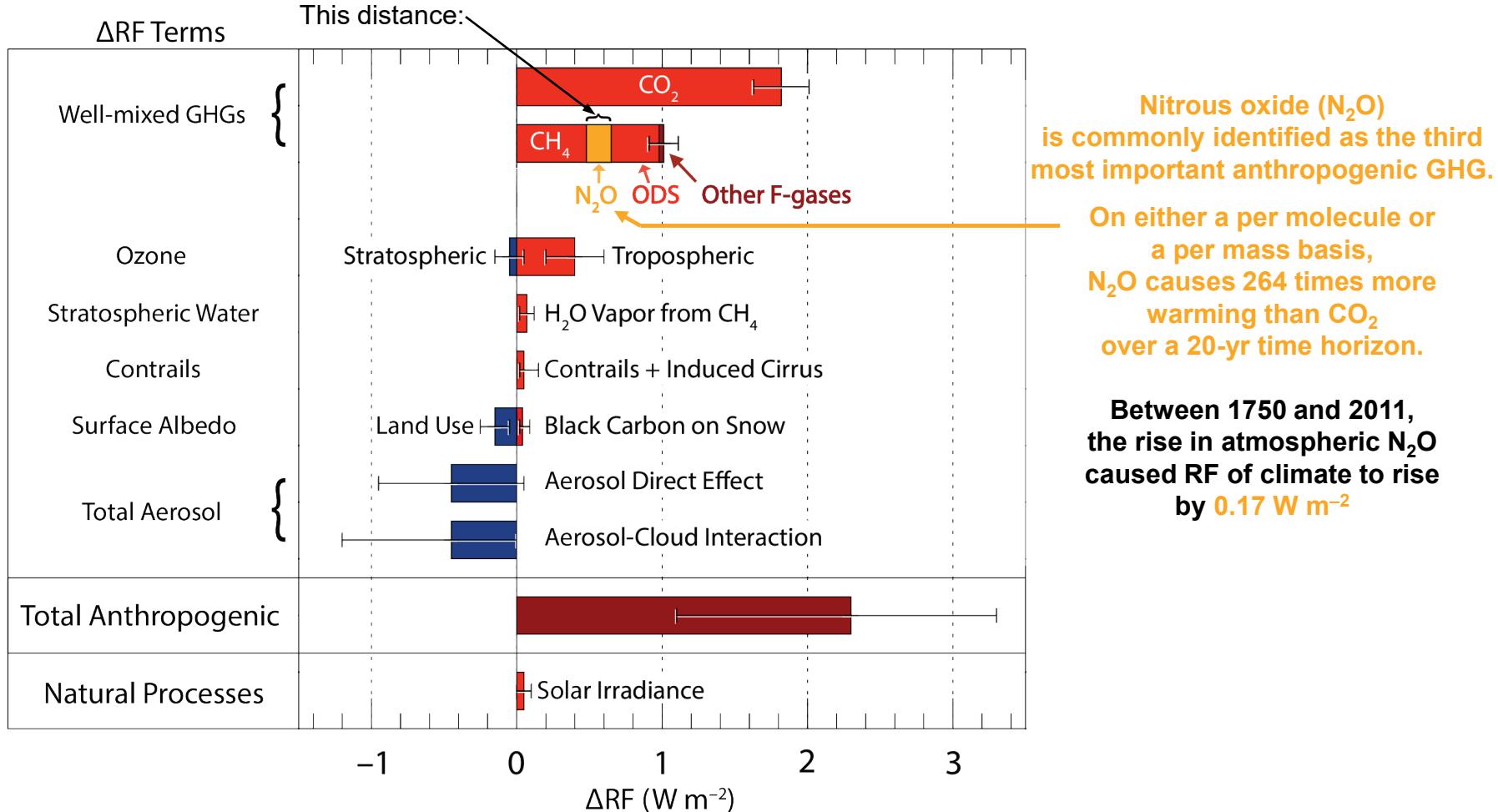


Figure 1-4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

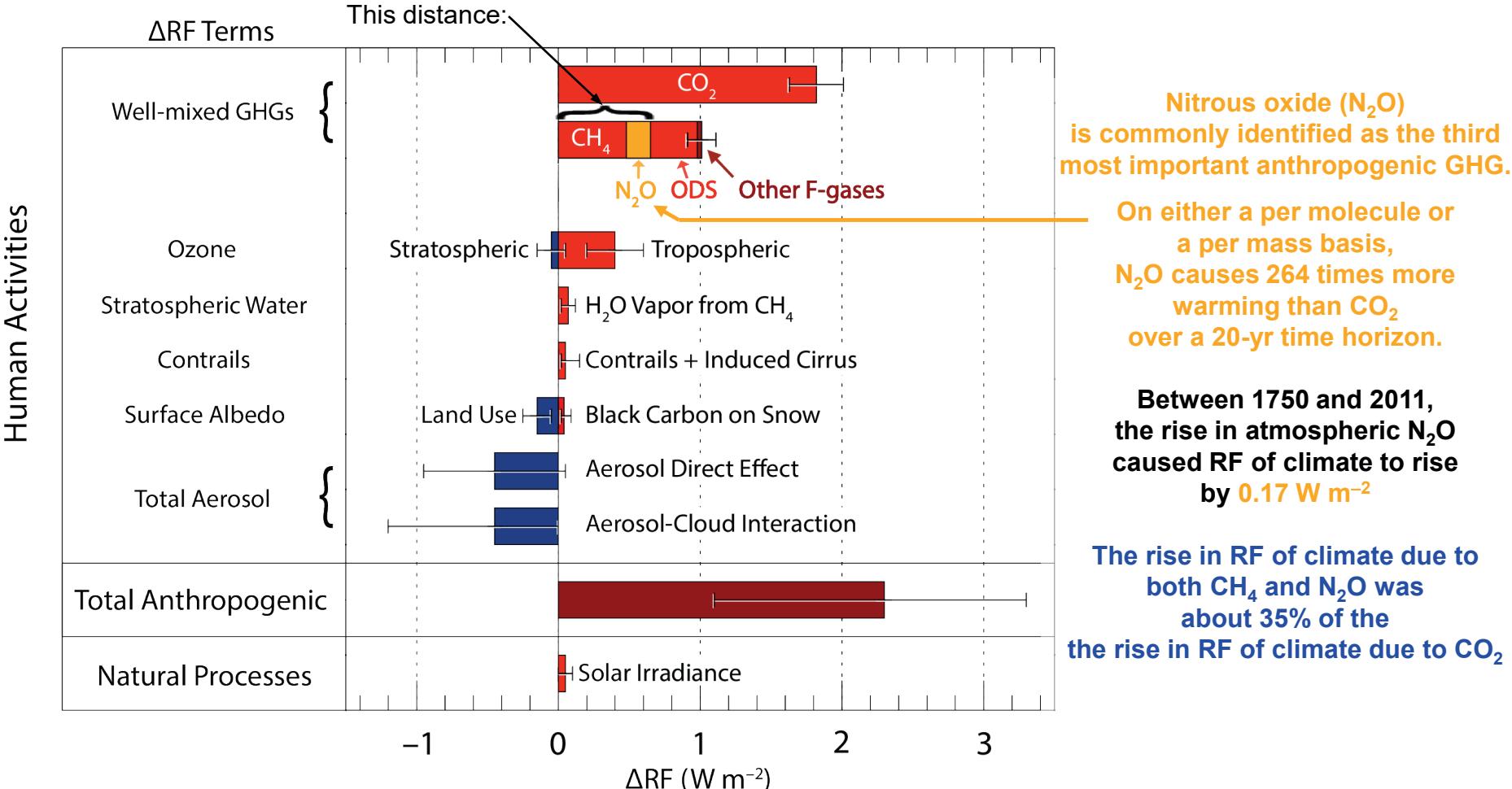


Figure 1-4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

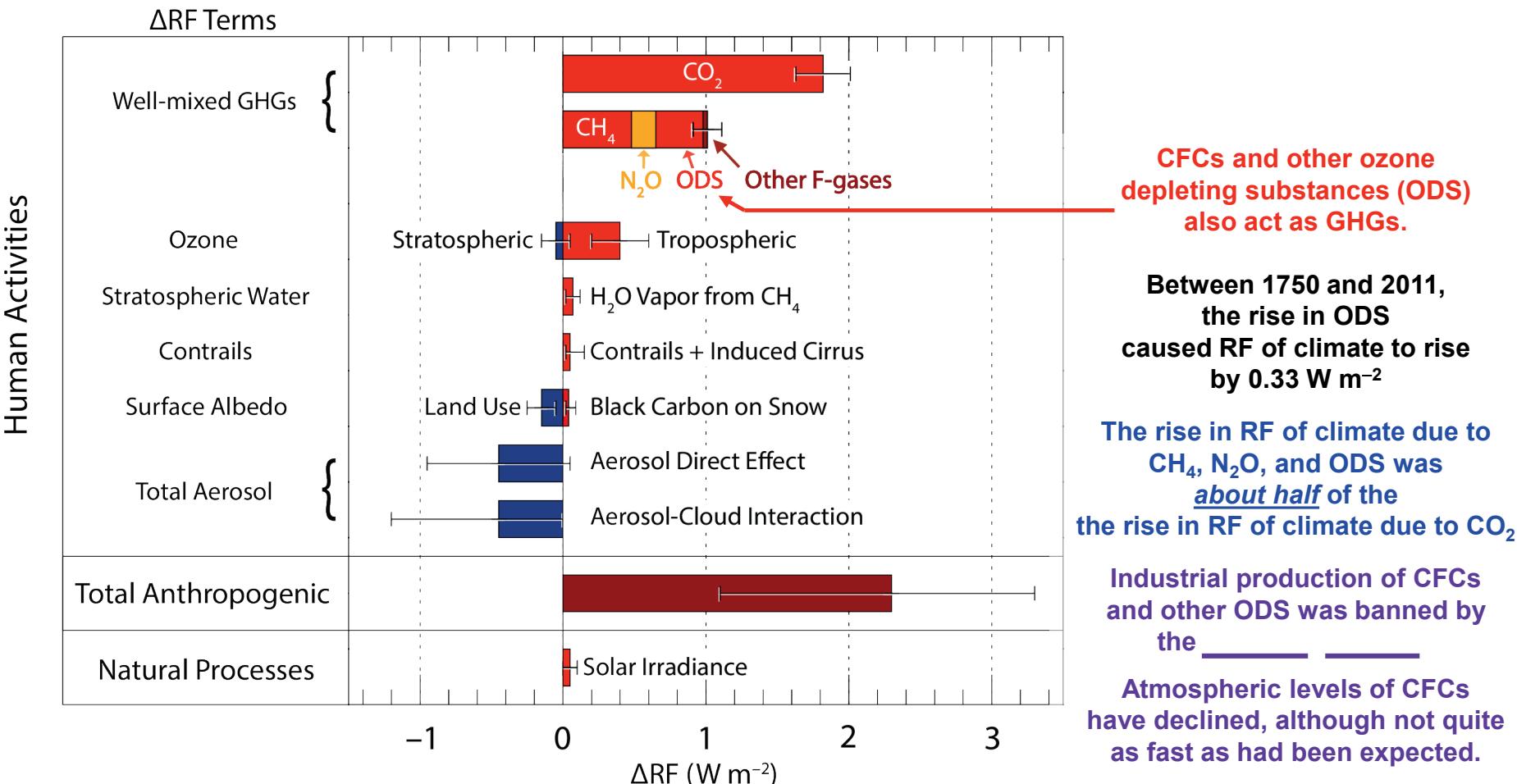


Figure 1-4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

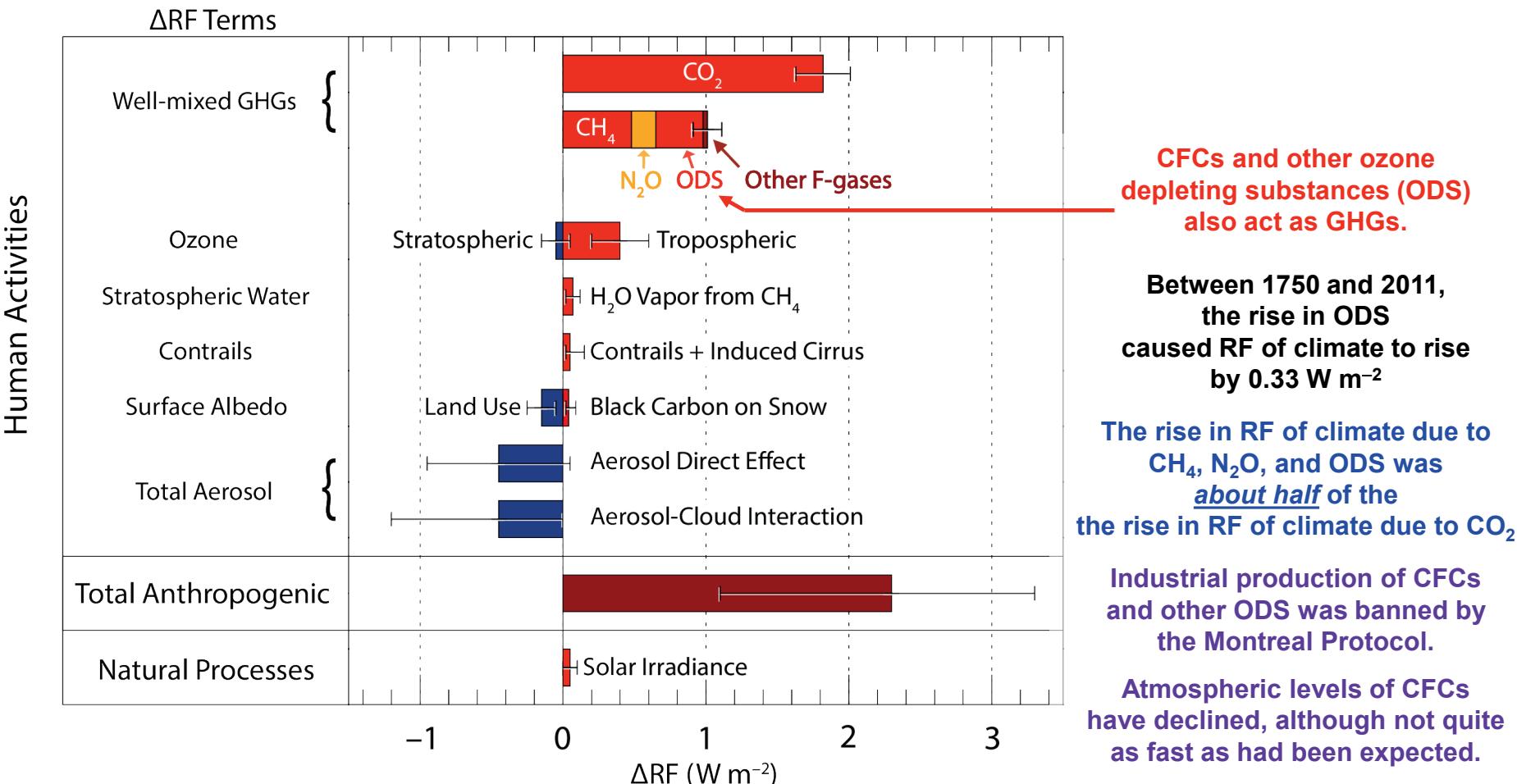


Figure 1-4, Paris Beacon of Hope

Ozone in the Atmosphere

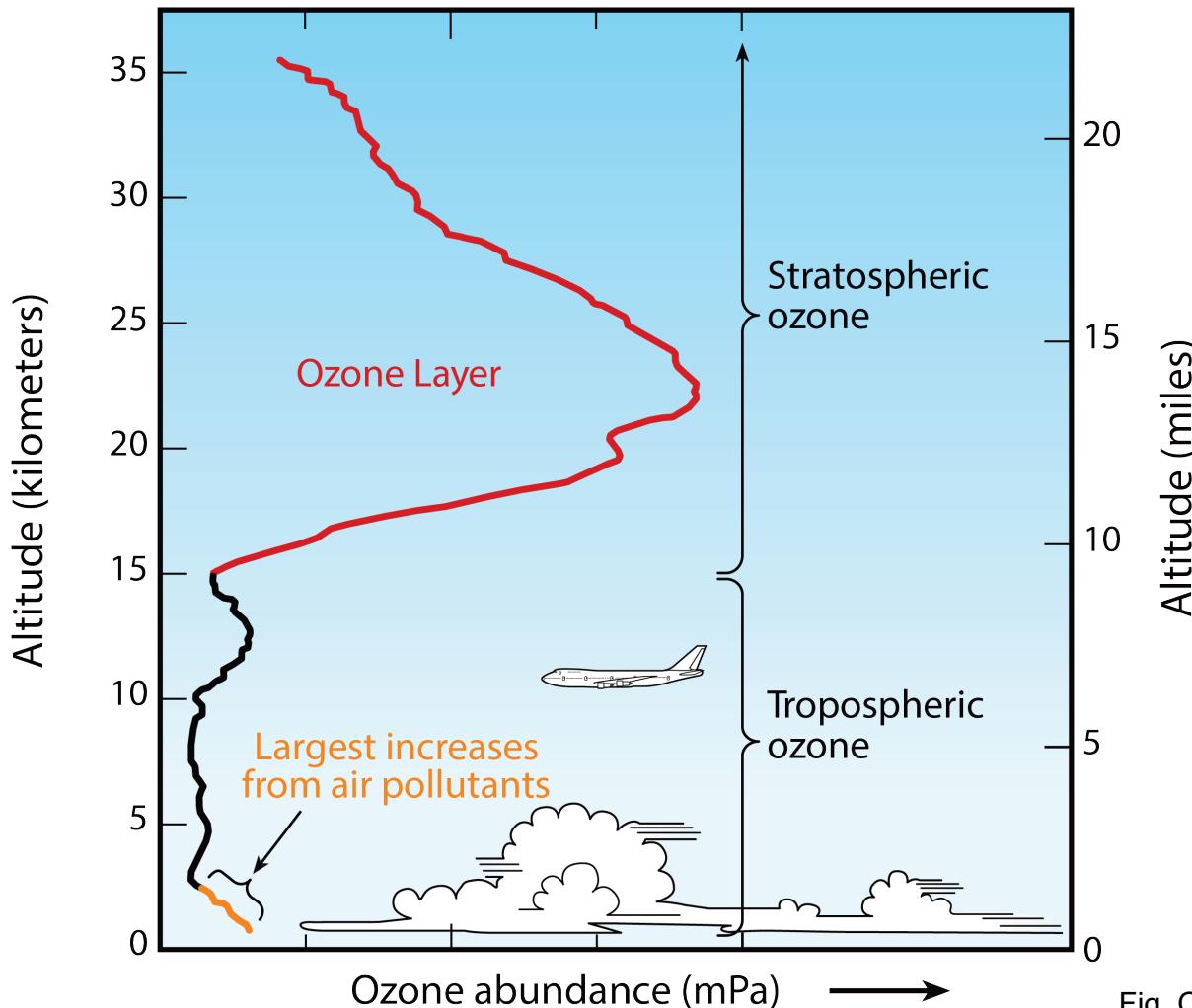
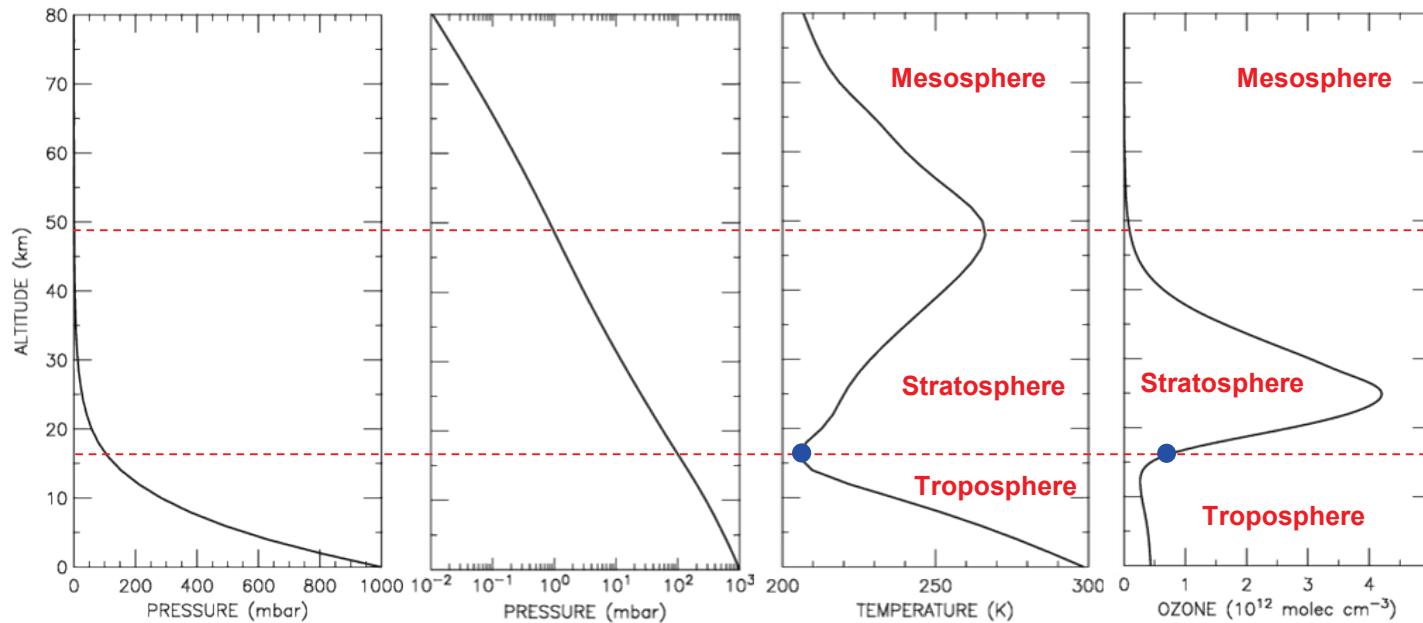


Fig. Q1-2
20 QAs about the Ozone Layer

It is incredible that human activity
both destroys stratospheric ozone (so-called good ozone)
and produces tropospheric ozone (so-called bad ozone)

Temperature versus Altitude

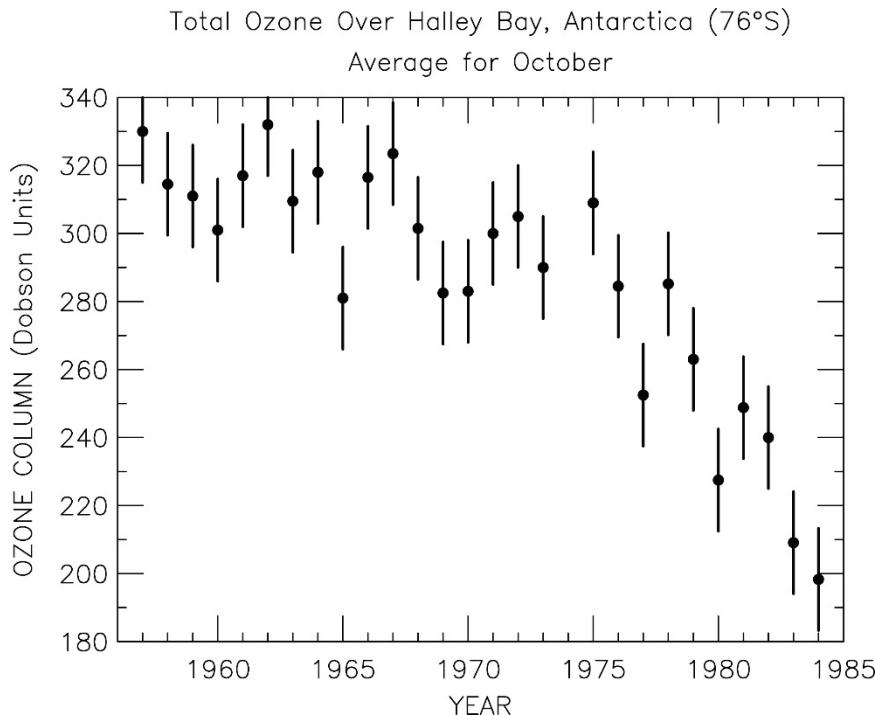


- T falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

Fourth chart expresses abundance of ozone concentration, or ozone density, or $[\text{O}_3]$, in units of molecules / cm^3

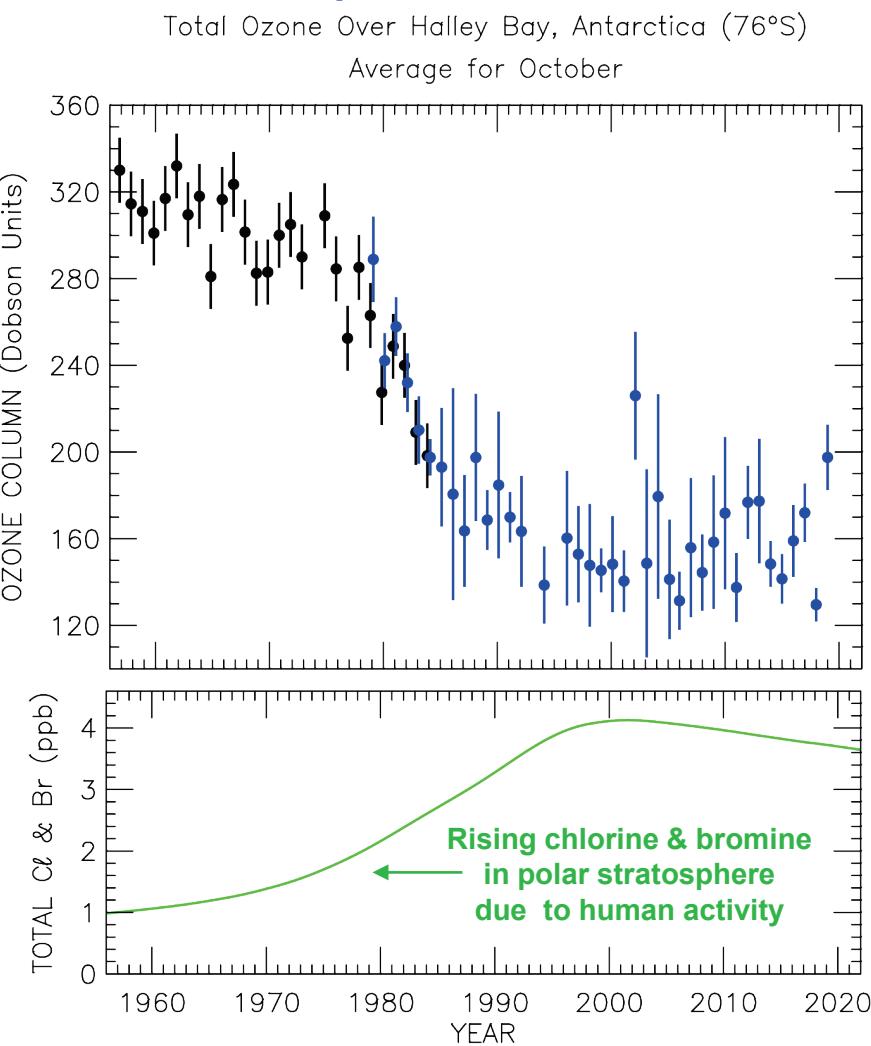
Earth's Atmosphere – Effect of Humans

Stratospheric Ozone – shields surface from solar UV radiation

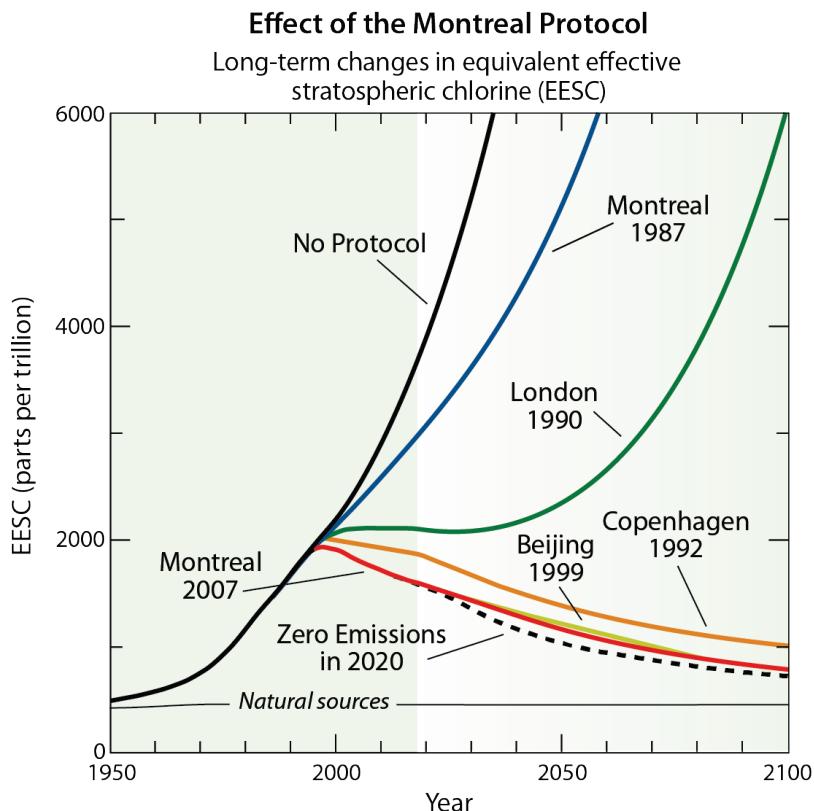


After Farman *et al.*, Large losses of total ozone in Antarctica reveal Seasonal ClO_x/NO_x interaction, Nature, 315, 207, 1985.

Update



And Atmospheric Levels of these Pollutants are Declining



CFCs: Chlorofluorocarbons

Contain some combination of chlorine, fluorine, and at least one carbon. Freons are a trade name for CFCs.

Bromocarbons:

Contain bromine, perhaps chlorine, and at least one carbon. Halons are a trade name for bromocarbons.

HCFCs: Hydro-chlorofluorocarbons

Same as CFCs, except one or more hydrogen has replaced a chlorine.

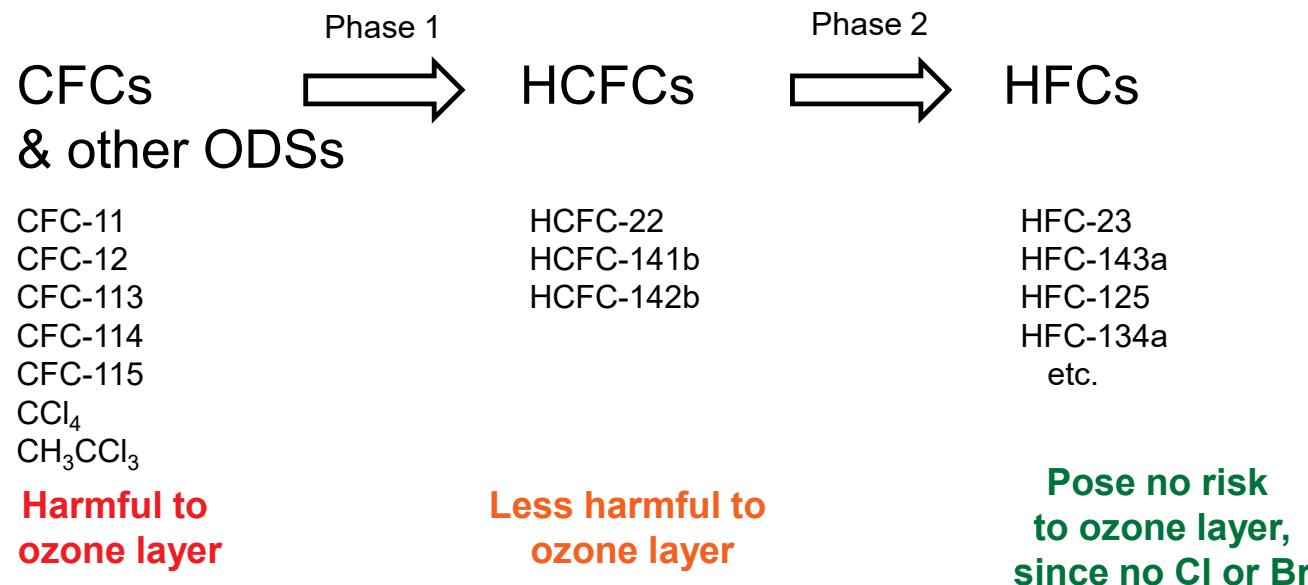
HFCs: Hydrofluorocarbons

Contain some combination of hydrogen, fluorine, and carbon. **These gases do not contain any bromine or chlorine, and hence pose no damage to the ozone layer.** Some HFCs are potent GHGs.

EESC: Equivalent, effective stratospheric chlorine. Reflects combined influence of chlorine and bromine on ozone, via a simple formula: $[\text{Chlorine}] + 60 \times [\text{Bromine}]$

Figure Q14-1, 20 QAs about the Ozone Layer

Phase out of CFCs and other Ozone Depleting Substances (ODSs)



See http://www.atmos.umd.edu/~rjs/class/spr2020/supplemental_readings/Naming_Convention_for_CFCs_Halons.pdf
for a guide to CFC naming convention

Phase out of CFCs and other Ozone Depleting Substances (ODSs)

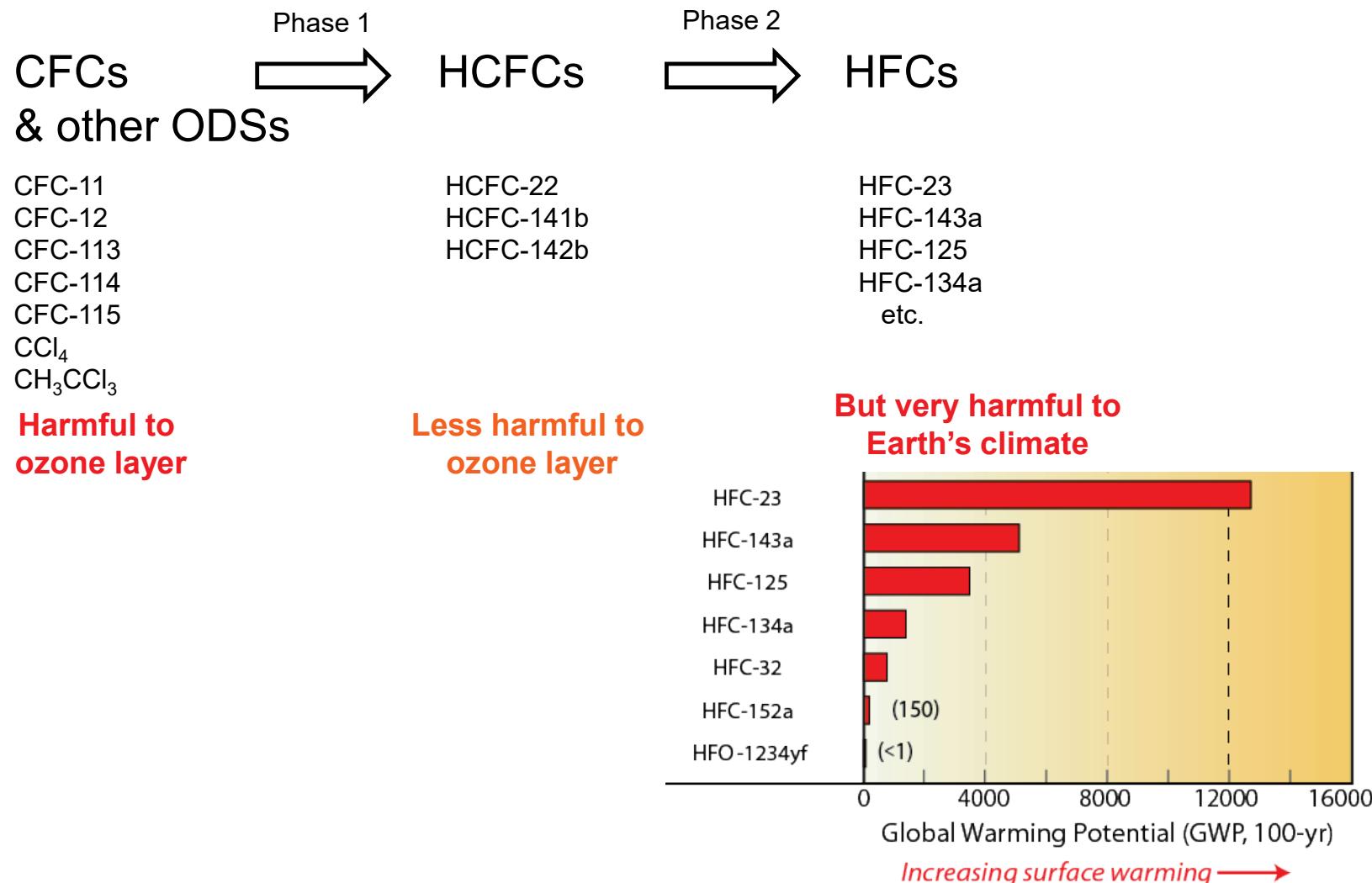


Figure Q17-3, 20 QAs about the Ozone Layer

Kigali Amendment



Tina Birnpili, Ozone Secretariat

PhD in Environmental Management and Economics from *Imperial College of Science, Technology and Medicine*, London.

As of 1 January 2019, the future production of HFCs is controlled by the Montreal Protocol, based on amendment passed in Kigali, Rwanda in Fall 2017, hosted by Rwanda's Minister of Environment Vincent Biruta and Canada's Minister of Environment and Climate Change Catherine McKenna.

"Ozone Secretariat Executive Secretary Tina Birnpili added: "2017 marks the 30th anniversary of the Protocol's life and there is no better way to celebrate this anniversary than by seeking country support to ratify the Kigali Amendment and build on the next 30 years."

<https://ozone.unep.org/unga-high-level-event-ratification-kigali-amendment>

Climate Benefit of the Kigali Amendment

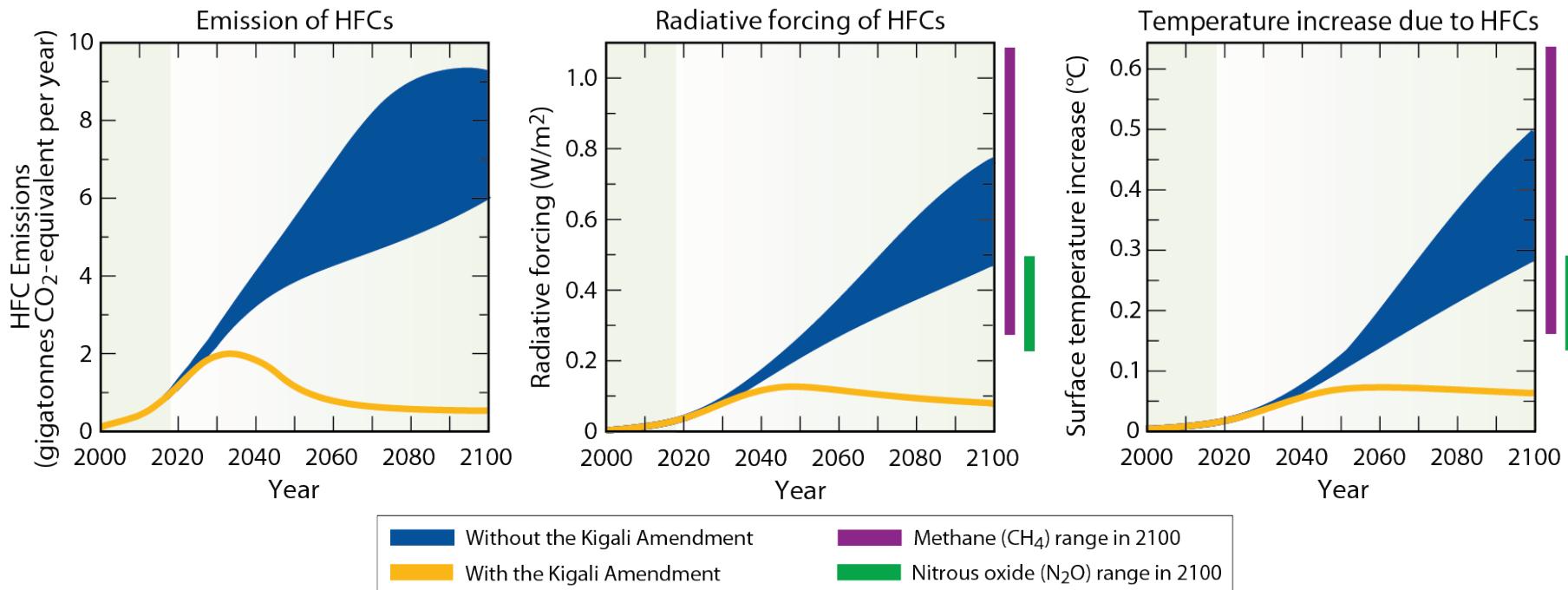


Figure Q19-2, 20 QAs about the Ozone Layer

Radiative Forcing of Climate, 1750 to 2011

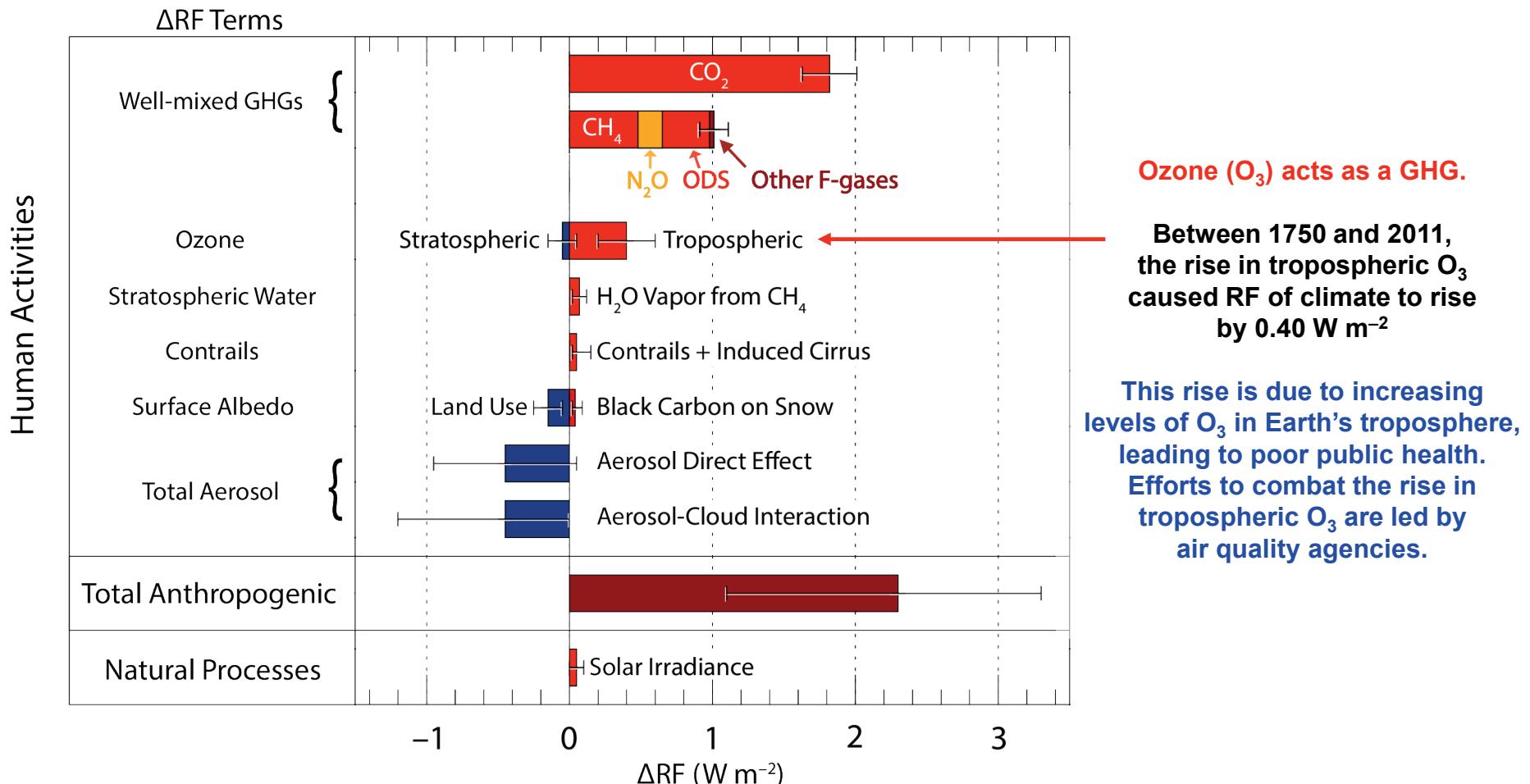


Figure 1-4, Paris Beacon of Hope

Radiative Forcing of Climate, 1750 to 2011

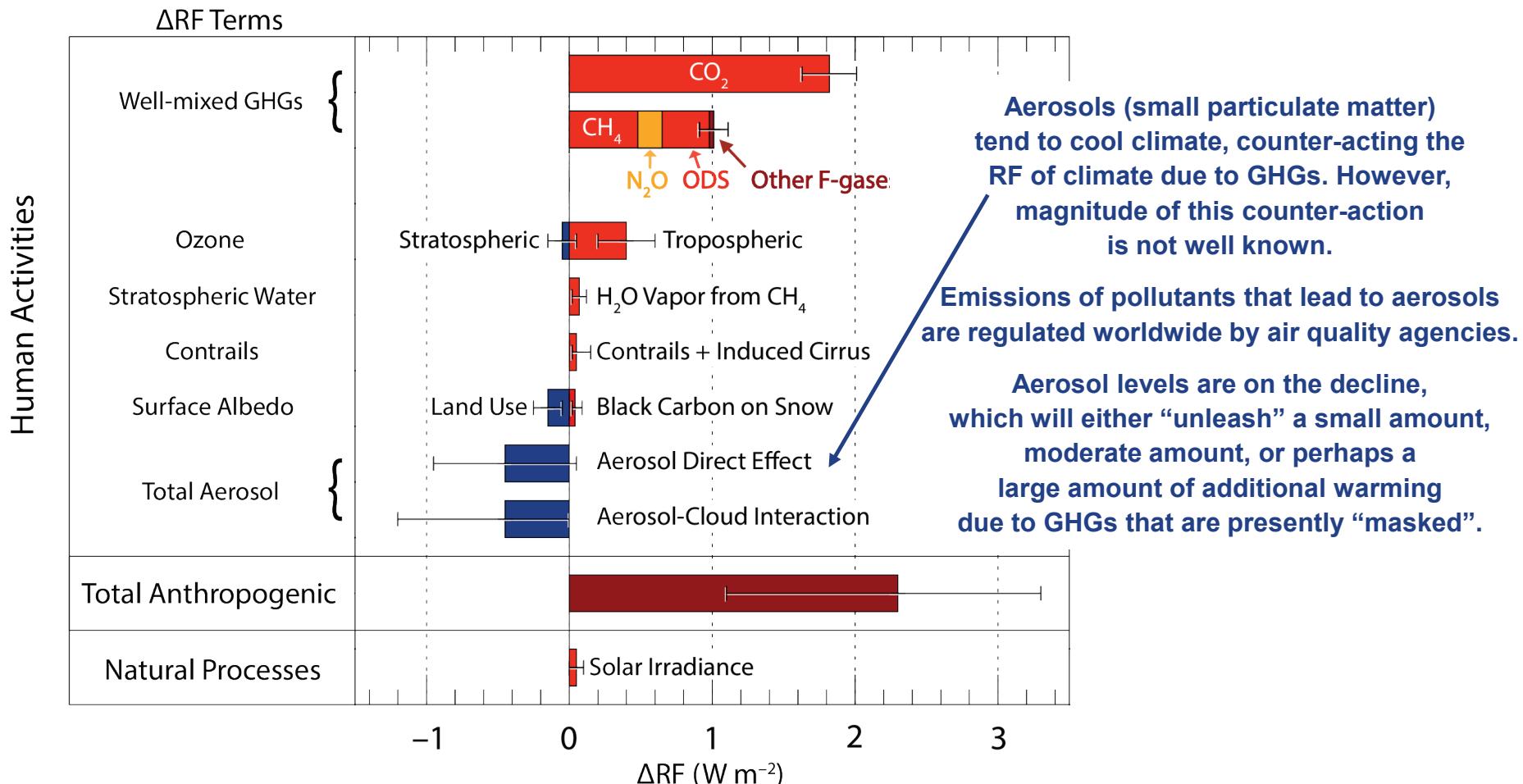


Figure 1-4, Paris Beacon of Hope

Earth's Climate History

Accordion-like unraveling of Earth's climate and CO₂

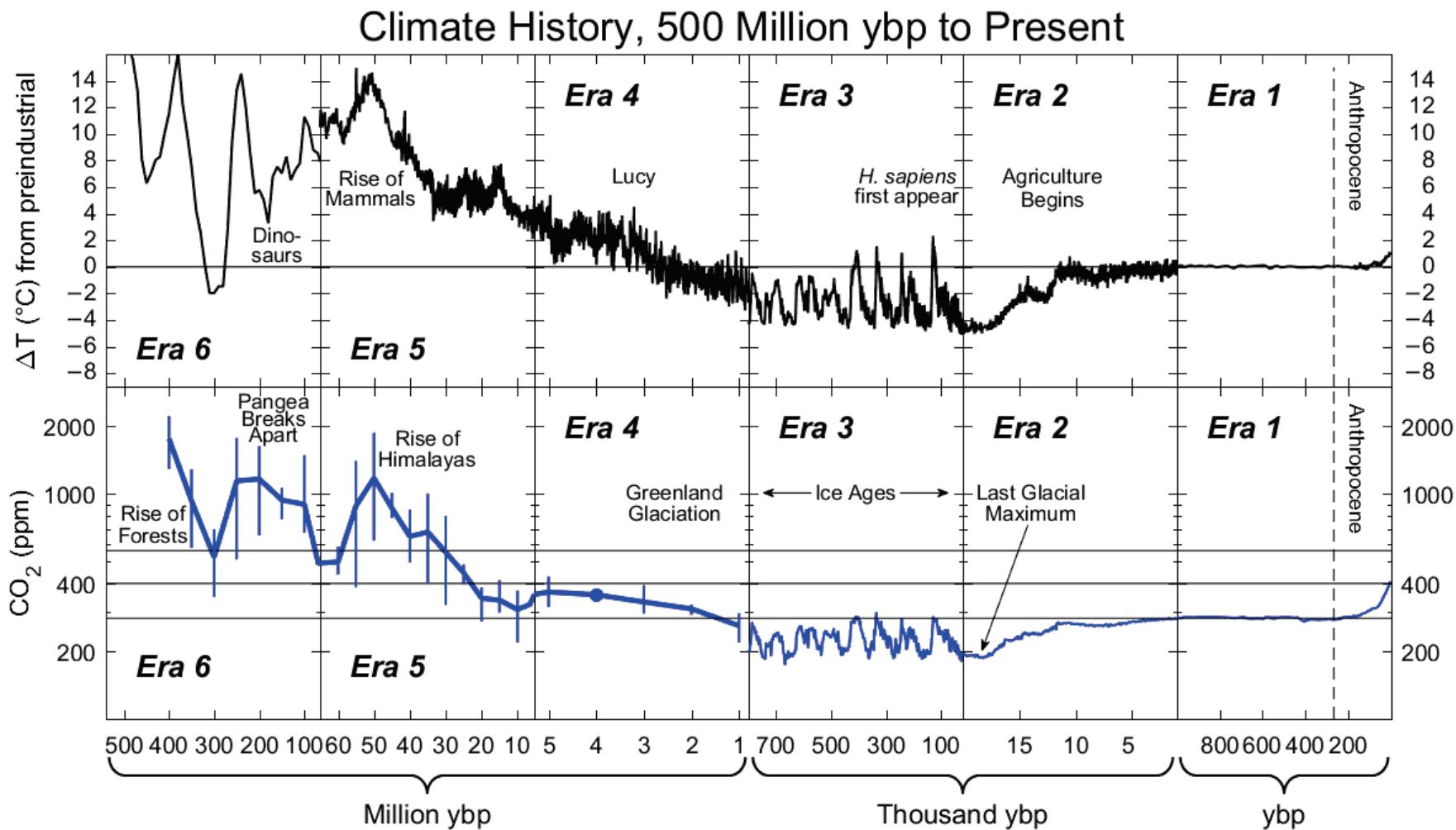


Fig 1.1, Paris Beacon of Hope

Going Back 600,000 years

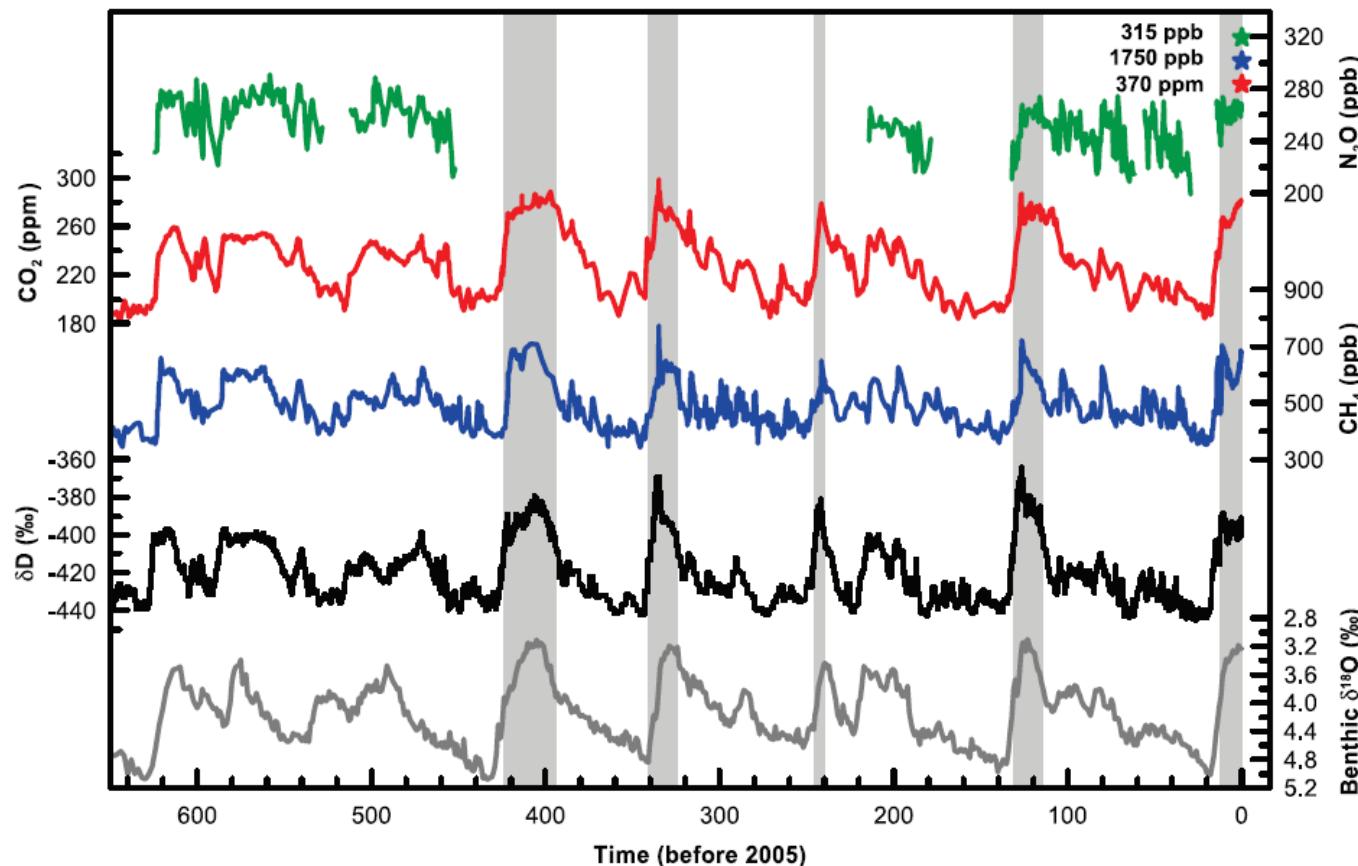


Figure 6.3. Variations of deuterium (δD ; black), a proxy for local temperature, and the atmospheric concentrations of the greenhouse gases CO_2 (red), CH_4 (blue), and nitrous oxide (N_2O ; green) derived from air trapped within ice cores from Antarctica and from recent atmospheric measurements (Petit et al., 1999; Lüthi et al., 2000; EPICA community members, 2004; Spahni et al., 2005; Siegenthaler et al., 2005a,b). The shading indicates the last Interglacial warm periods. Interglacial periods also existed prior to 450 ka, but these were apparently colder than the typical Interglacials of the latest Quaternary. The length of the current Interglacial is not unusual in the context of the last 650 kyr. The stack of 57 globally distributed benthic $\delta^{18}O$ marine records (dark grey), a proxy for global ice volume fluctuations (Lisicki and Raymo, 2005), is displayed for comparison with the ice core data. Downward trends in the benthic $\delta^{18}O$ curve reflect increasing ice volumes on land. Note that the shaded vertical bars are based on the ice core age model (EPICA community members, 2004), and that the marine record is plotted on its original time scale based on tuning to the orbital parameters (Lisicki and Raymo, 2005). The stars and labels indicate atmospheric concentrations at year 2000.

Figure 6.3, IPCC 2007

See <https://epic.awi.de/id/eprint/18400/1/Oer2008a.pdf> for description of EPICA , European Project for Ice Coring in Antarctica

GHG Record Over Last Several Millennia

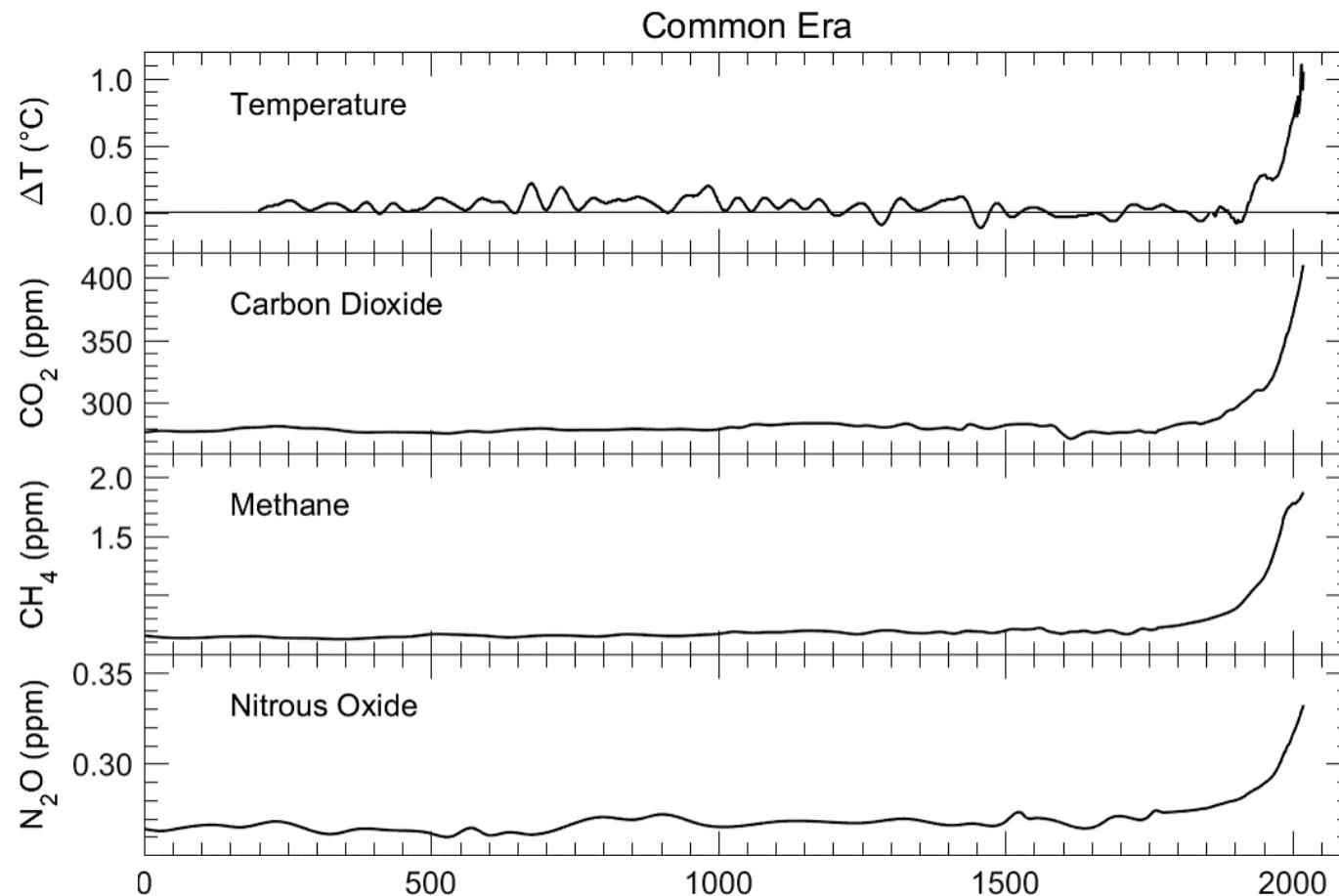


Figure 1.2, Paris Beacon of Hope (updated)

GHG Record Over Last Several Millennia

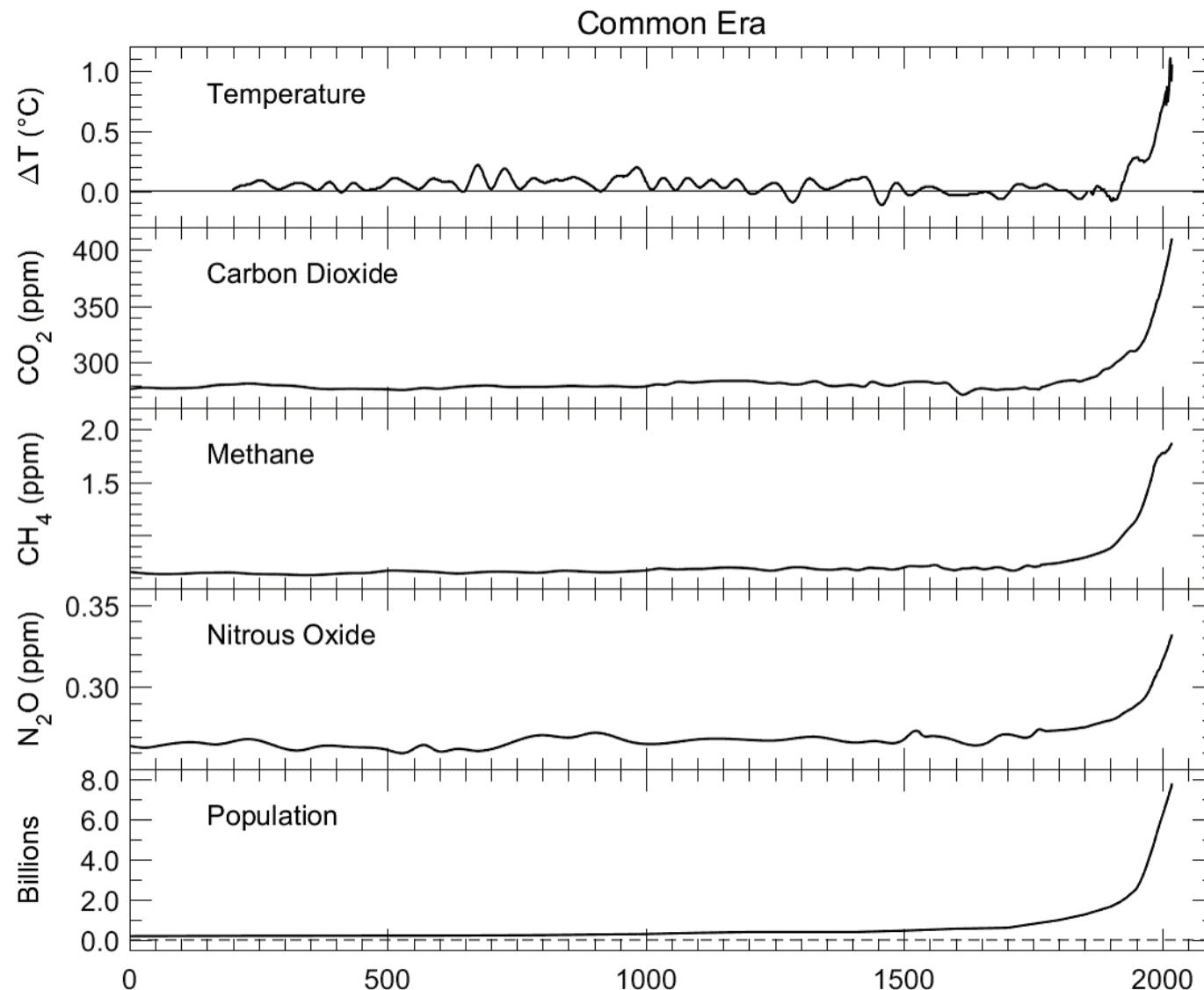


Figure 1.2, Paris Beacon of Hope (updated)

GWP – Global Warming Potential

$$\text{GWP}(\text{CH}_4) = \frac{\int_{\text{time initial}}^{\text{time final}} a_{\text{CH}_4} \times [\text{CH}_4(t)] dt}{\int_{\text{time initial}}^{\text{time final}} a_{\text{CO}_2} \times [\text{CO}_2(t)] dt}$$

where:

a_{CH_4} = Radiative Efficiency ($\text{W m}^{-2} \text{ kg}^{-1}$) due to an increase in CH_4

a_{CO_2} = Radiative Efficiency ($\text{W m}^{-2} \text{ kg}^{-1}$) due to an increase in CO_2

$\text{CH}_4(t)$ = time-dependent response to an instantaneous release of a pulse of certain mass of CH_4

$\text{CO}_2(t)$ = time-dependent response to an instantaneous release of a pulse of the same mass of CO_2

$$\text{GWP}(\text{N}_2\text{O}) = \frac{\int_{\text{time initial}}^{\text{time final}} a_{\text{N}_2\text{O}} \times [\text{N}_2\text{O}(t)] dt}{\int_{\text{time initial}}^{\text{time final}} a_{\text{CO}_2} \times [\text{CO}_2(t)] dt}$$

GWP – Global Warming Potential

Table 1-1, Paris Beacon of Hope

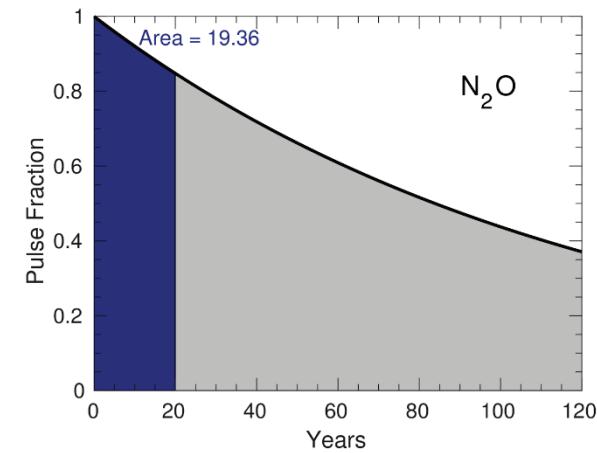
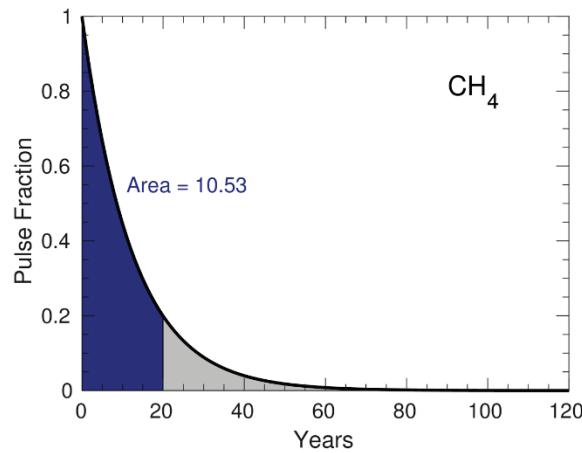
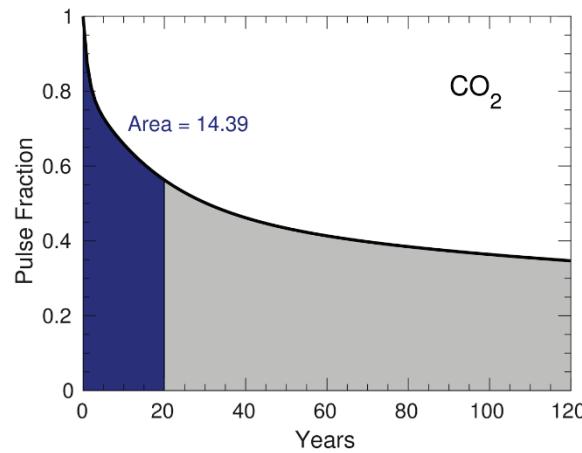
$$GWP(CH_4) = \frac{\int_{time\ initial}^{time\ final} \alpha_{CH_4} \times [CH_4(t)] dt}{\int_{time\ initial}^{time\ final} \alpha_{CO_2} \times [CO_2(t)] dt}$$

$$GWP(N_2O) = \frac{\int_{time\ initial}^{time\ final} \alpha_{N_2O} \times [N_2O(t)] dt}{\int_{time\ initial}^{time\ final} \alpha_{CO_2} \times [CO_2(t)] dt}$$

GHG	IPCC (1995)	IPCC (2001)	IPCC (2007)	IPCC (2013)
<i>100 Year Time Horizon</i>				
CH ₄	21	23	25	28, 34*
N ₂ O	310	296	298	265, 298*
<i>20 Year Time Horizon</i>				
CH ₄	56	62	72	84, 86*
N ₂ O	280	275	289	264, 268*

*Allowing for carbon cycle feedback

20 Year Time Horizon means time final = 20 years in these integrals



$$CO_2(t) = 0.217 + 0.186 \times CO_2(t=0) e^{-t/1.286} + 0.338 \times CO_2(t=0) e^{-t/18.59} + 0.249 \times CO_2(t=0) e^{-t/172.9}$$

$$CH_4(t) = CH_4(t=0) e^{-t/12.4}$$

$$N_2O(t) = N_2O(t=0) e^{-t/121.0}$$

where all times are given in units of year

GWP – Global Warming Potential

Table 1-1, Paris Beacon of Hope

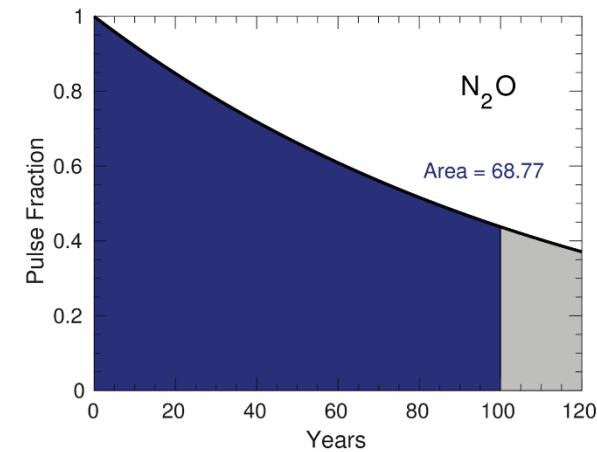
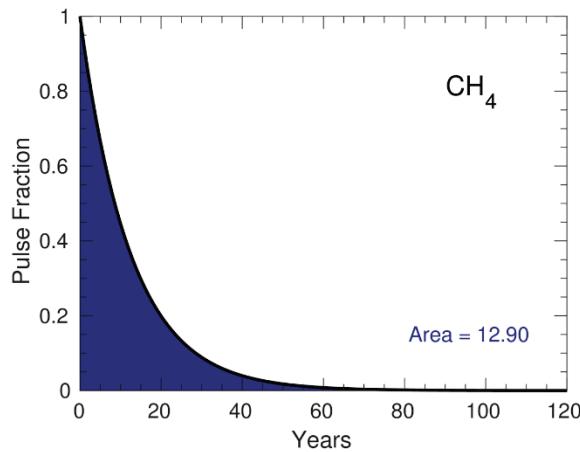
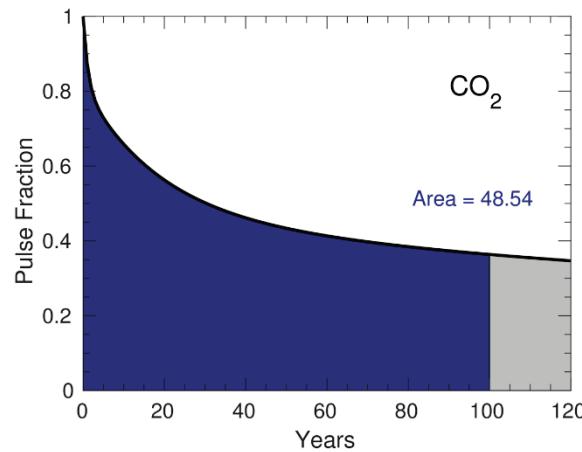
$$GWP(CH_4) = \frac{\int_{time\ initial}^{time\ final} \alpha_{CH_4} \times [CH_4(t)] dt}{\int_{time\ initial}^{time\ final} \alpha_{CO_2} \times [CO_2(t)] dt}$$

$$GWP(N_2O) = \frac{\int_{time\ initial}^{time\ final} \alpha_{N_2O} \times [N_2O(t)] dt}{\int_{time\ initial}^{time\ final} \alpha_{CO_2} \times [CO_2(t)] dt}$$

GHG	IPCC (1995)	IPCC (2001)	IPCC (2007)	IPCC (2013)
100 Year Time Horizon				
CH ₄	21	23	25	28 34*
N ₂ O	310	296	298	265 298*
20 Year Time Horizon				
CH ₄	56	62	72	84, 86*
N ₂ O	280	275	289	264, 268*

*Allowing for carbon cycle feedback

100 Year Time Horizon means time final = 100 years in these integrals



$$CO_2(t) = 0.217 + 0.186 \times CO_2(t=0) e^{-t/1.286} + 0.338 \times CO_2(t=0) e^{-t/18.59} + 0.249 \times CO_2(t=0) e^{-t/172.9}$$

$$CH_4(t) = CH_4(t=0) e^{-t/12.4}$$

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where all times are given in units of year

GWP – Global Warming Potential

Table 1-1, Paris Beacon of Hope

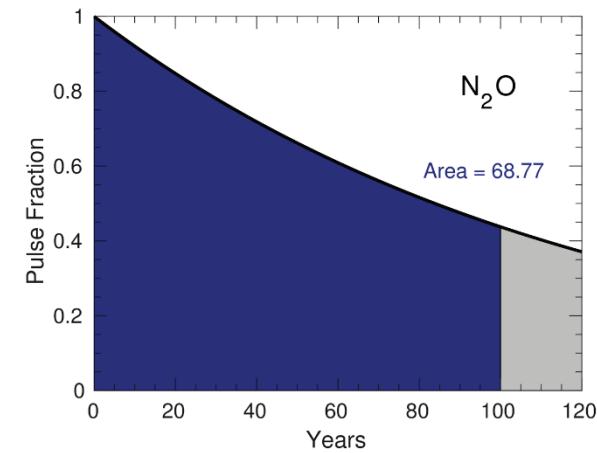
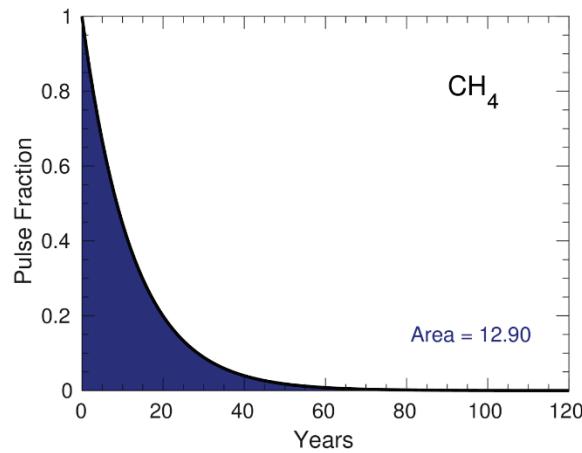
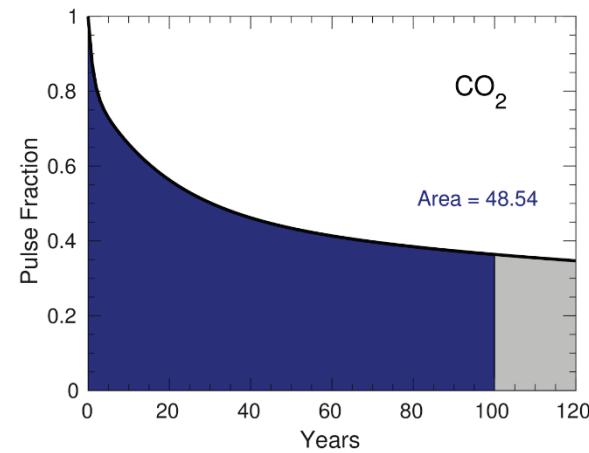
$$GWP(CH_4) = \frac{\int_{time\ initial}^{time\ final} a_{CH_4} \times [CH_4(t)] dt}{\int_{time\ initial}^{time\ final} a_{CO_2} \times [CO_2(t)] dt}$$

$$GWP(N_2O) = \frac{\int_{time\ initial}^{time\ final} a_{N_2O} \times [N_2O(t)] dt}{\int_{time\ initial}^{time\ final} a_{CO_2} \times [CO_2(t)] dt}$$

GHG	IPCC (1995)	IPCC (2001)	IPCC (2007)	IPCC (2013)
<i>100 Year Time Horizon</i>				
CH ₄	21	23	25	28, 34*
N ₂ O	310	296	298	265, 298*
<i>20 Year Time Horizon</i>				
CH ₄	56	62	72	84, 86*
N ₂ O	280	275	289	264, 268*

*Allowing for carbon cycle feedback

These numbers (i.e., 1995, 2001, 2007, & 2013) are publication years



$$CO_2(t) = 0.217 + 0.186 \times CO_2(t=0) e^{-t/1.286} + 0.338 \times CO_2(t=0) e^{-t/18.59} + 0.249 \times CO_2(t=0) e^{-t/172.9}$$

$$CH_4(t) = CH_4(t=0) e^{-t/12.4}$$

$$N_2O(t) = N_2O(t=0) e^{-t/121.0}$$

where all times are given in units of year

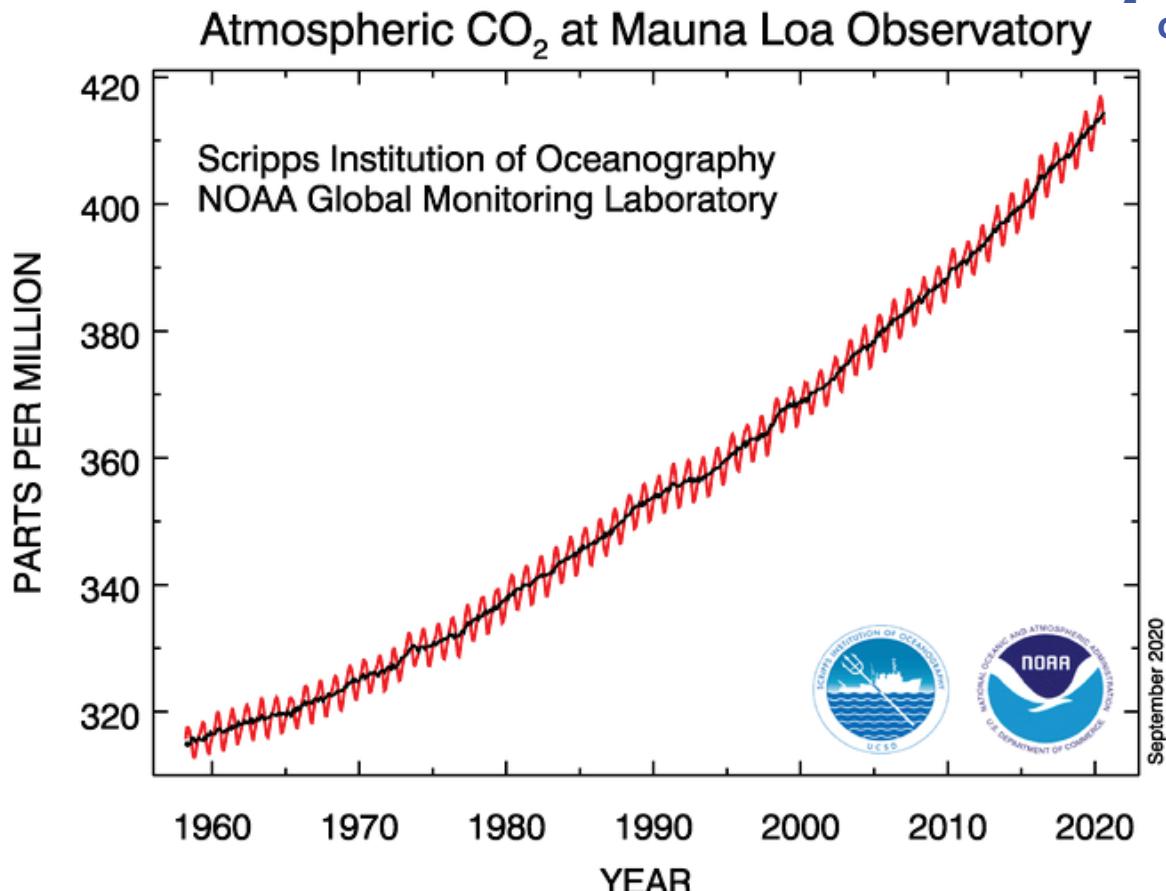
Modern CO₂ Record

CO₂ at MLO on 15 Sept 2020: 411.84 parts per million (ppm)

CO₂ at MLO on 14 Sept 2019: 408.12 parts per million (ppm)

$\Delta\text{CO}_2 = 3.71 \text{ ppm per year}$

or 0.91 % per year



Legacy of Charles Keeling, Scripps Institution of Oceanography, La Jolla, CA

https://www.esrl.noaa.gov/gmd/webdata/ccgg/trends/co2_data_mlo.png

See also <https://www.co2.earth/daily-co2>

Atmospheric CH₄

AT6, Q1:

According to Table 3.2 of Chemistry in Context, what was pre-industrial atmospheric abundance of CH₄ **and** is this consistent with Figure 3.7 of the Houghton reading?

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700 ppb

Table 3.2

Examples of Greenhouse Gases

Broadly

Name and Chemical Formula	Preindustrial Concentration (1750)	Concentration in 2008	Atmospheric Lifetime (years)	Anthropogenic Sources	Global Warming Potential
carbon dioxide CO ₂	270 ppm	388 ppm	50-200*	Fossil fuel combustion, deforestation, cement production	1
methane CH ₄	700 ppb	1760 ppb	12	Rice paddies, waste dumps, livestock	21
nitrous oxide N ₂ O					310
CFC-12 CCl ₂ F ₂					8100

*A single value for based on several measurements

The given is an estimate

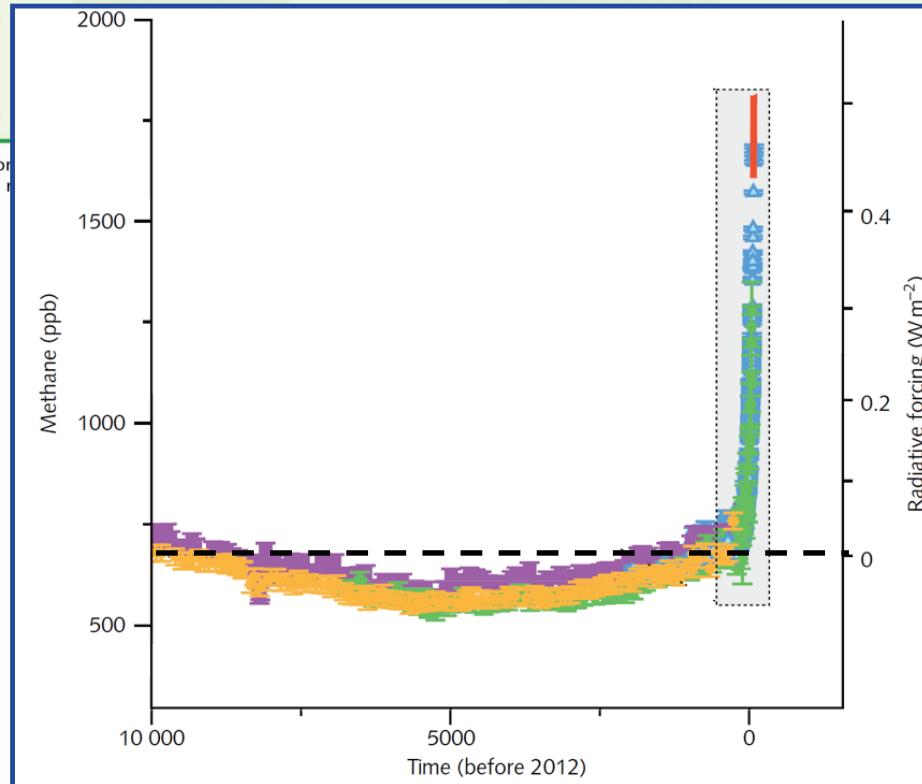


Figure 3.7, Houghton

Atmospheric CH₄

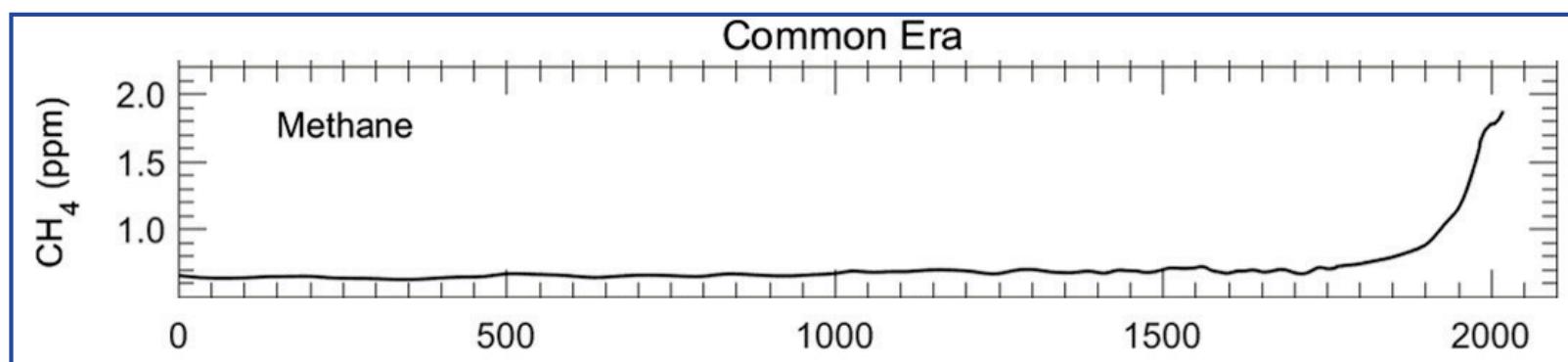
AT6, Q2:

What is the approximate current atmospheric abundance of CH₄?

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Examples of Greenhouse Gases					
Name and Chemical Formula	Preindustrial Concentration (1750)	Concentration in 2008	Atmospheric Lifetime (years)	Anthropogenic Sources	Global Warming Potential
carbon dioxide CO ₂	270 ppm	388 ppm	50-200*	Fossil fuel combustion, deforestation, cement production	1
methane CH ₄	700 ppb	1760 ppb	12	Rice paddies, waste dumps, livestock	21
nitrous oxide N ₂ O	275 ppb	322 ppb	120	Fertilizers, industrial production, combustion	310
CFC-12 CCl ₂ F ₂	0	0.56 ppb	102	Liquid coolants, foams	8100

*A single value for the atmospheric lifetime of CO₂ is not possible. Removal mechanisms take place at different rates. The range given is an estimate based on several removal mechanisms.



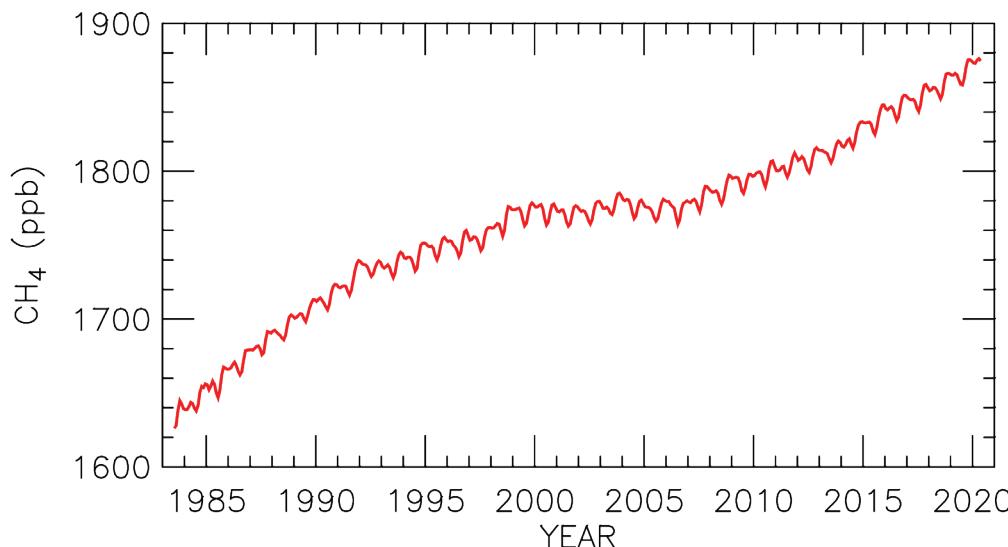
Atmospheric CH₄

AT6, Q2:

What is the approximate current atmospheric abundance of CH₄?

NOAA Earth System Research Laboratory (Boulder, Co) is “go to” place for information regarding GHGs

Latest data indicate CH₄ is at 1876.2 ppb and rising, and also that CH₄ exceeded 1760 ppb in late-1990s and exceeded 1.84 ppm in mid-2017.



https://www.esrl.noaa.gov/gmd/ccgg/trends_ch4

Simple Climate Model

$$\Delta T = \lambda_p (1 + f_{H_2O}) (\Delta F_{CO_2} + \Delta F_{CH_4+N_2O} + \Delta F_{OTHER\ GHGs} + \Delta F_{AEROSOLS}) - OHE$$

where

$$\lambda_p = 0.3 \text{ K} / \text{W m}^{-2}$$

OHE = Ocean Heat Export

Climate models that consider water vapor feedback find:

$$\lambda \approx 0.63 \text{ K} / \text{W m}^{-2}, \text{ from which we deduce } f_{H_2O} = 1.08$$

Simple Climate Model

$$\Delta T = \lambda_p (1 + f_{H2O \text{ & LR}}) (\Delta F_{CO_2} + \Delta F_{CH4+N2O} + \Delta F_{\text{OTHER GHGs}} + \Delta F_{\text{AEROSOLS}}) - OHE$$

where

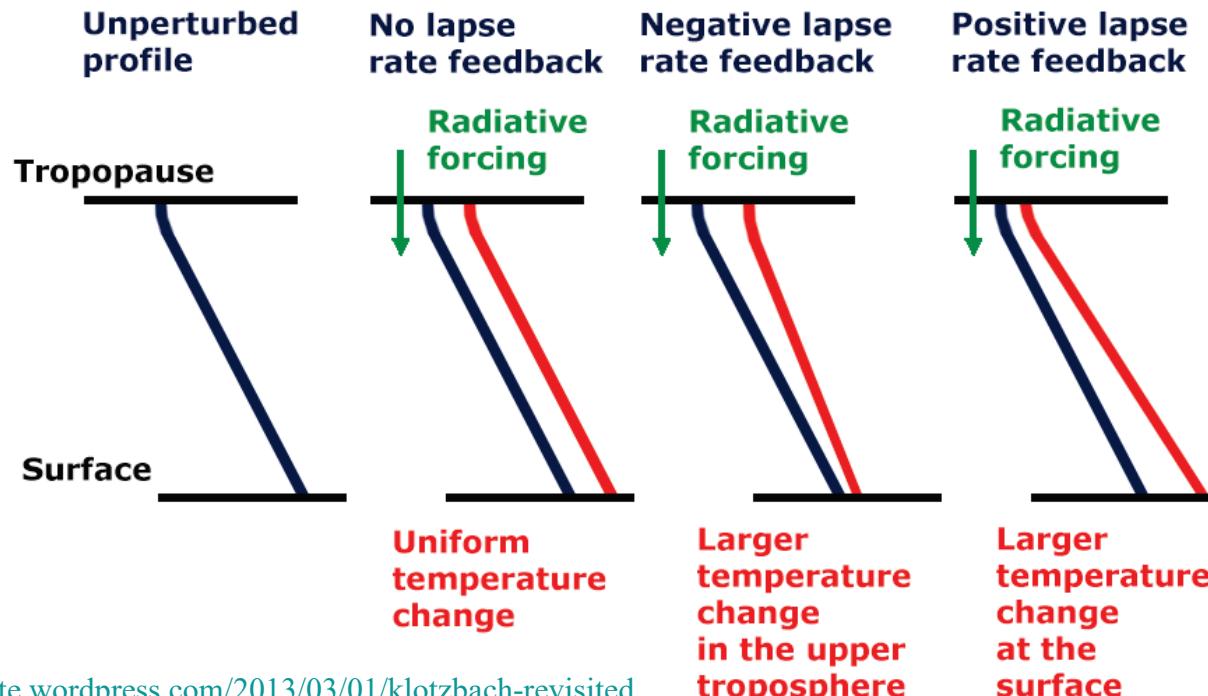
$$\lambda_p = 0.3 \text{ K} / \text{W m}^{-2}$$

OHE = Ocean Heat Export

Climate models that consider water vapor feedback & lapse rate feedback find:

$$\lambda \approx 0.44 \text{ K} / \text{W m}^{-2}, \text{ from which we deduce } f_{H2O \text{ & Lapse Rate}} = 0.45$$

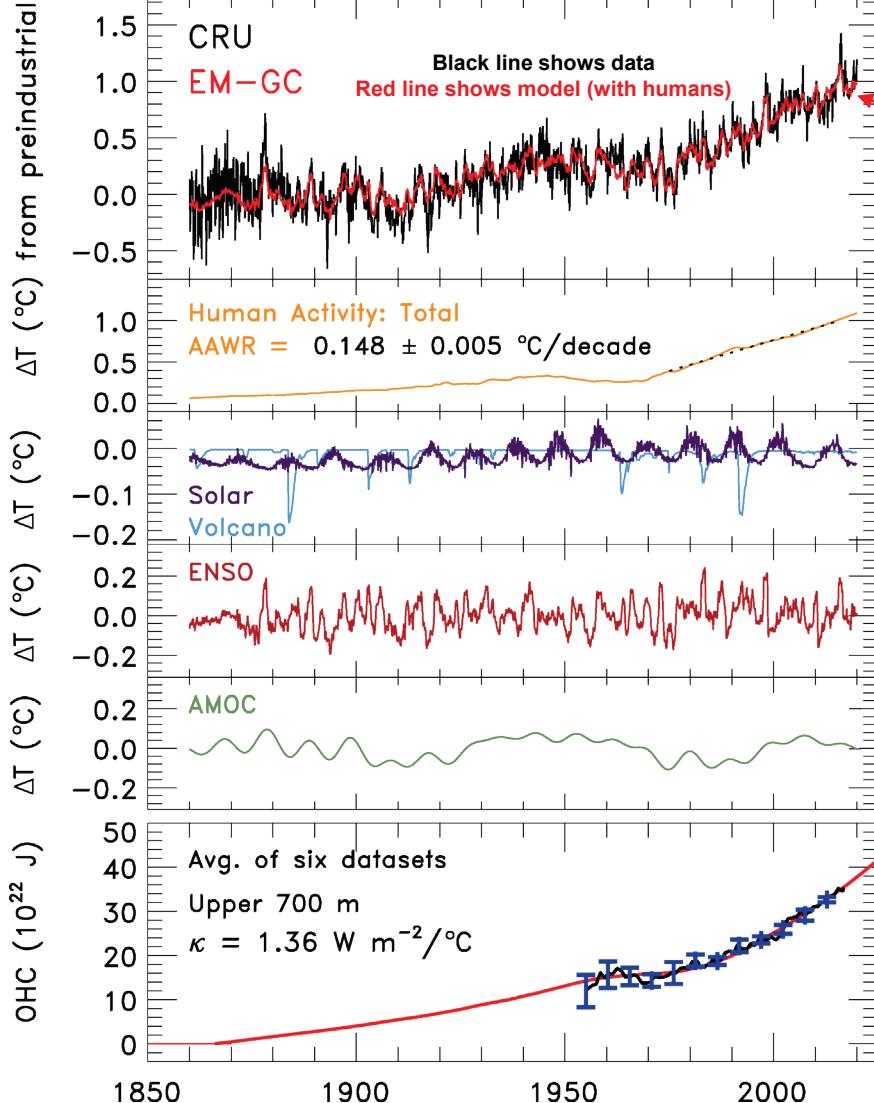
Lapse Rate Feedback



<https://ourchangingclimate.wordpress.com/2013/03/01/klotzbach-revisited>

- Photons emitted in UT can escape to space more easily than photons emitted near surface
- If UT warms more than surface, bulk atmospheric emissivity increases
 - UT :upper troposphere Emissivity: efficiency in which thermal energy is radiated
- GCMs indicate water vapor & lapse rate feedbacks are intricately linked, with the former almost certainly being positive (in response to rising GHGs), the latter almost certainly being negative, and the sum probably being positive
- Definition of empirical lapse rate feedback has some controversy due to differences in opinion regarding proper interpretation of temperature data in UT from various Microwave Sounding Unit (MSU) instruments

Are humans responsible?



CRU: Climate Research Unit of East Anglia, United Kingdom

EM-GC: Empirical Model of Global Climate, Univ of Maryland

$$\Delta T = \lambda_p (1 + f_{\text{TOTAL}}) (\Delta F_{\text{CO}_2} + \Delta F_{\text{CH}_4+\text{N}_2\text{O}} + \Delta F_{\text{OTHER GHGs}} + \Delta F_{\text{AEROSOLS}}) - \text{OHE}$$

Here, $f_{\text{TOTAL}} = 0.99$

This model run assumes $\Delta F_{\text{AEROSOLS}}$ in 2011 is -0.9 W m^{-2}

where f_{TOTAL} is dimensionless climate sensitivity parameter that represents feedbacks, and is related to IPCC definition of feedbacks (Bony et al., *J. Climate*, 2006) via:

$$1 + f_{\text{TOTAL}} = \frac{1}{1 - \text{FB}_{\text{TOTAL}} \lambda_p}$$

and $\text{FB}_{\text{TOTAL}} = \text{FB}_{\text{WATER VAPOR}} + \text{FB}_{\text{LAPSE RATE}} + \text{FB}_{\text{CLOUDS}} + \text{FB}_{\text{SURFACE ALBEDO}} + \text{etc}$

Each FB term has units of $\text{W m}^{-2} \text{ K}^{-1}$, the reciprocal of the units of λ_p . The utility of this approach is that feedbacks can be summed to get FB_{TOTAL} .

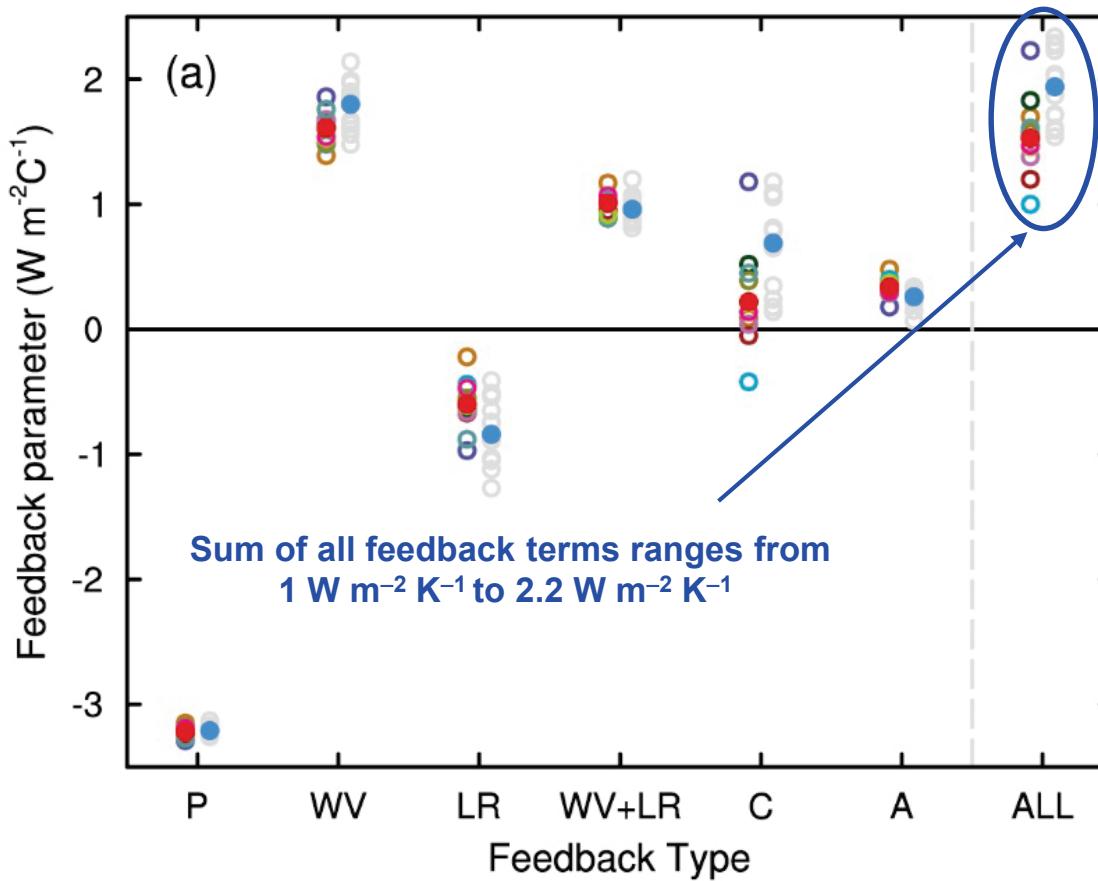
Lecture 2, Slide 37

Canty et al., 2013 <https://www.atmos-chem-phys.net/13/3997/2013/acp-13-3997-2013.html>

Hope et al., 2017 https://link.springer.com/chapter/10.1007/978-3-319-46939-3_2

as well as Hope et al. (2020, submitted) & McBride et al. (2020, submitted).

Figure provided by Laura McBride.



Low end of IPCC climate models (i.e., $\text{FB}_{\text{ALL}} = 1.0 \text{ W m}^{-2} \text{ K}^{-1}$) implies

$f_{\text{TOTAL}} = 0.45$; i.e., all feedbacks besides water vapor & lapse rate cancel

High end of IPCC climate models (i.e., $\text{FB}_{\text{ALL}} = 2.2 \text{ W m}^{-2} \text{ K}^{-1}$) implies

$f_{\text{TOTAL}} = 2.98$; i.e., clouds must exert a very strong positive feedback

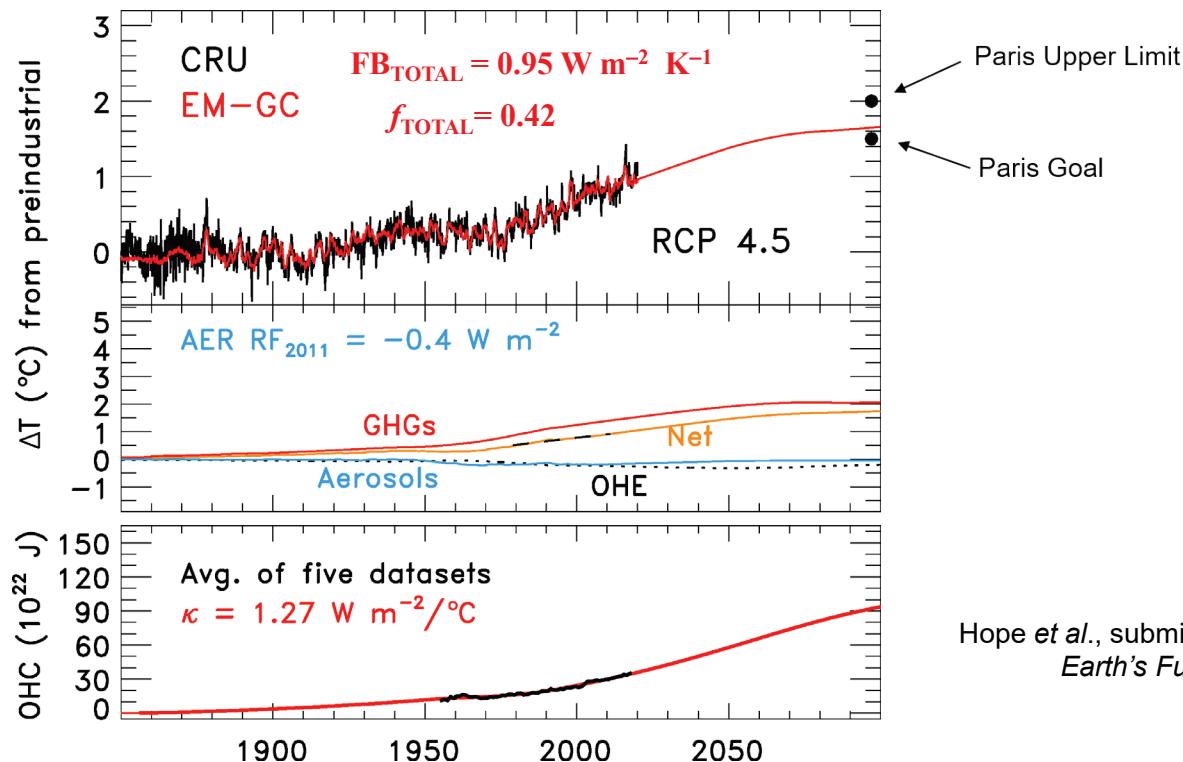
Uncertainty in RF due to aerosols is a huge complication that places a fundamental uncertainty on how well future global warming can be forecast

$$\Delta T \approx \lambda_p (1 + f_{\text{TOTAL}}) \Delta RF - OHE ; \quad 1 + f_{\text{TOTAL}} = \frac{1}{1 - FB_{\text{TOTAL}} \lambda_p}$$

f_{TOTAL} : total feedback term (dimensionless) due to water vapor, lapse rate, clouds, etc. that directly amplifies RF

FB_{TOTAL} : sum of individual feedback terms (units $\text{W m}^{-2} \text{ K}^{-1}$) computed using IPCC formalism

OHE : export of heat from atmosphere to world's ocean



Hope et al., submitted to AGU journal
Earth's Future, 2020.

We assume that whatever value of climate feedback is inferred from the climate record will persist into the future. For Aerosol RF in 2011 of -0.4 W m^{-2} & assuming best estimate for H_2O and Lapse Rate feedback is correct, this simulation implies sum of other feedbacks (clouds, surface albedo) must be **close to zero**.

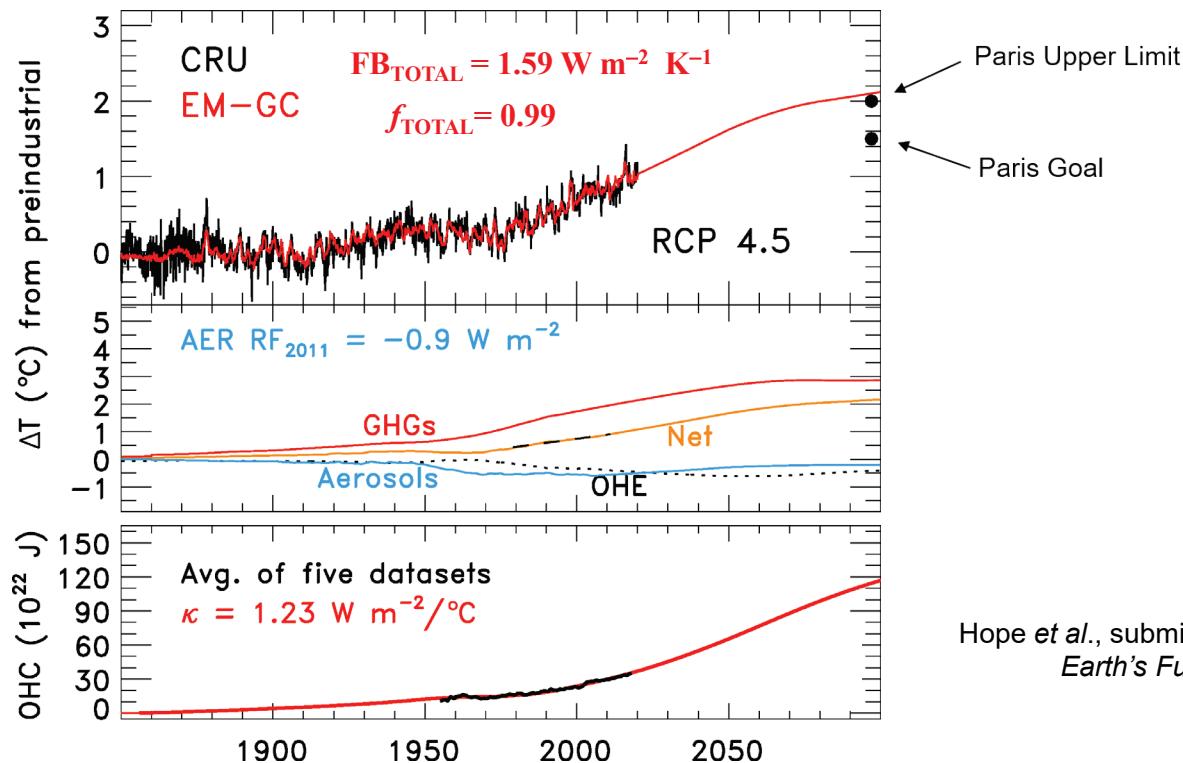
Uncertainty in RF due to aerosols is a huge complication that places a fundamental uncertainty on how well future global warming can be forecast

$$\Delta T \approx \lambda_p (1 + f_{\text{TOTAL}}) \Delta RF - OHE ; \quad 1 + f_{\text{TOTAL}} = \frac{1}{1 - FB_{\text{TOTAL}} \lambda_p}$$

f_{TOTAL} : total feedback term (dimensionless) due to water vapor, lapse rate, clouds, etc. that directly amplifies RF

FB_{TOTAL} : sum of individual feedback terms (units $\text{W m}^{-2} \text{ K}^{-1}$) computed using IPCC formalism

OHE : export of heat from atmosphere to world's ocean



We assume that whatever value of climate feedback is inferred from the climate record will persist into the future. For Aerosol RF in 2011 of **-0.9 W m⁻²** & assuming best estimate for H₂O and Lapse Rate feedback is correct, this simulation implies sum of other feedbacks (clouds, surface albedo) must be **moderately positive**.

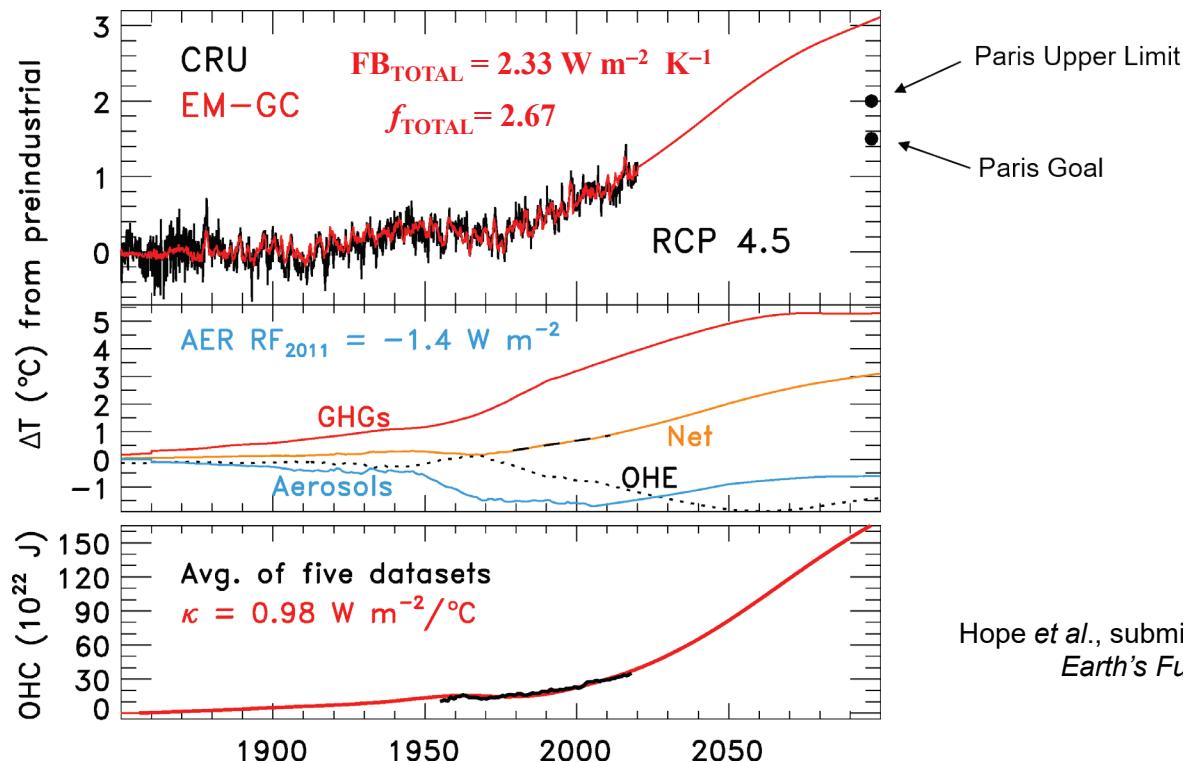
Uncertainty in RF due to aerosols is a huge complication that places a fundamental uncertainty on how well future global warming can be forecast

$$\Delta T \approx \lambda_p (1 + f_{\text{TOTAL}}) \Delta RF - OHE ; \quad 1 + f_{\text{TOTAL}} = \frac{1}{1 - FB_{\text{TOTAL}} \lambda_p}$$

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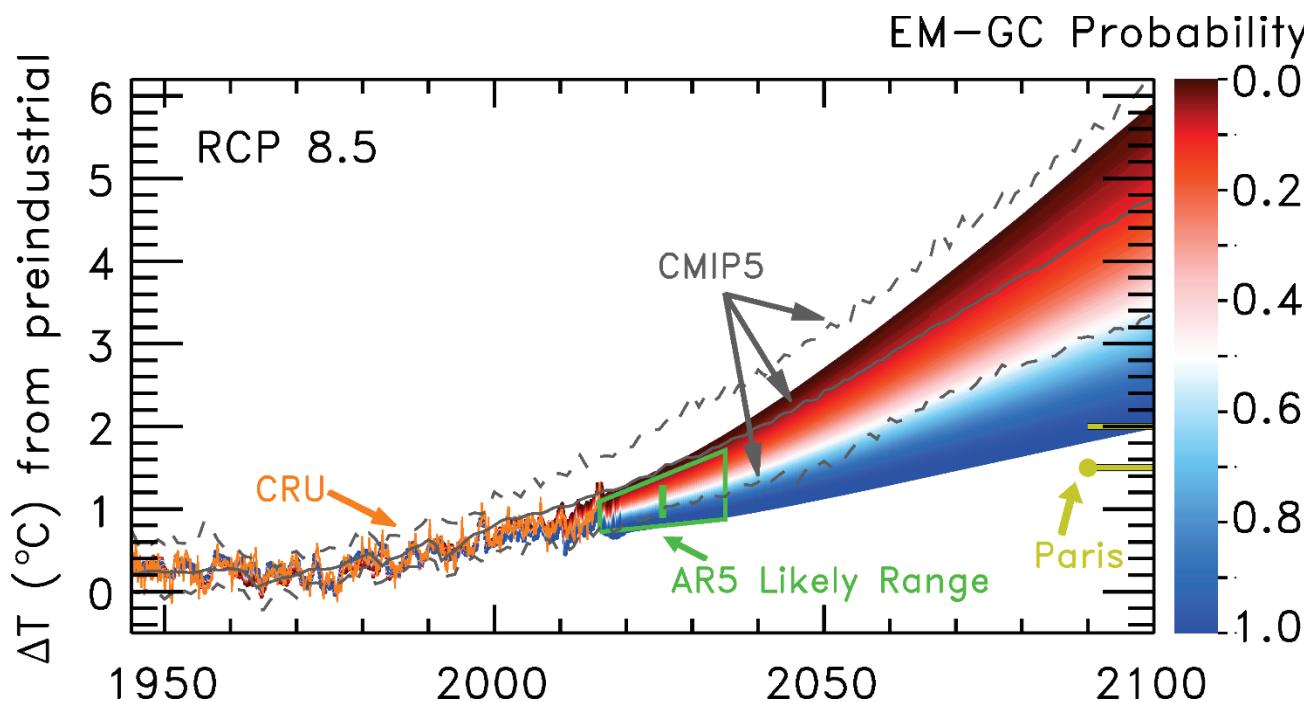
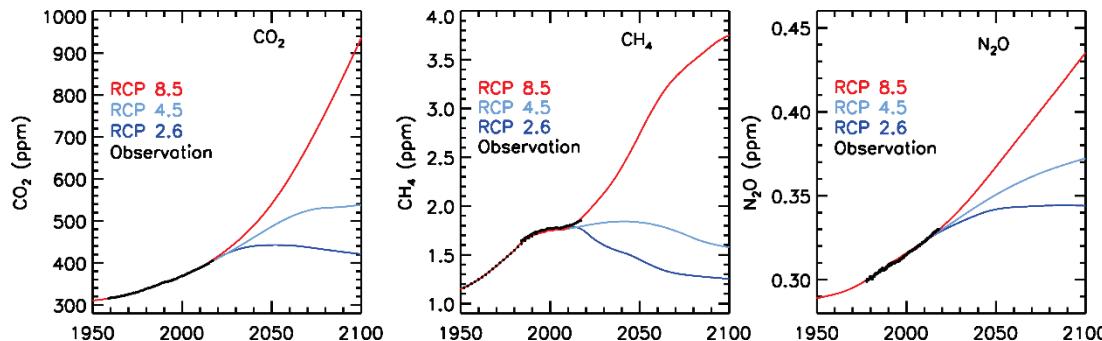
OHE : export of heat from atmosphere to world's ocean



Hope et al., submitted to AGU journal
Earth's Future, 2020.

We assume that whatever value of climate feedback is inferred from the climate record will persist into the future. For Aerosol RF in 2011 of -1.5 W m^{-2} & assuming best estimate for H₂O and Lapse Rate feedback is correct, this simulation implies sum of other feedbacks (clouds, surface albedo) must be **strongly positive**.

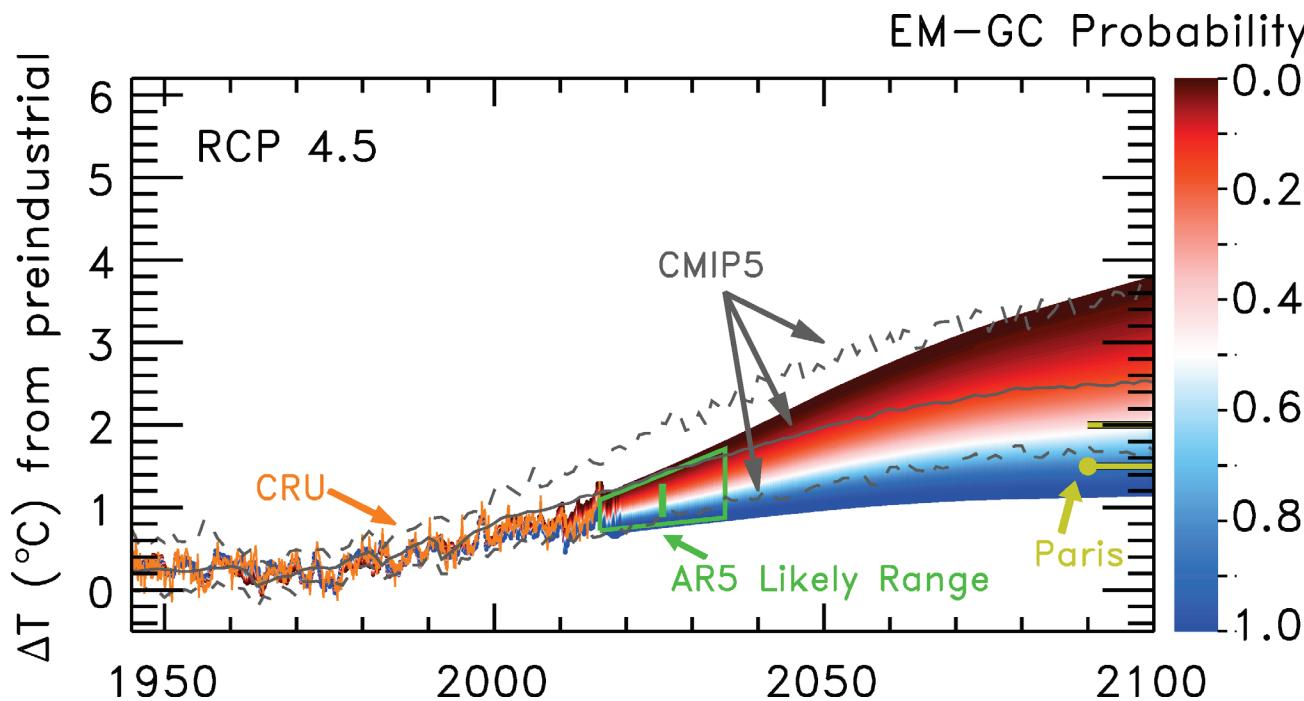
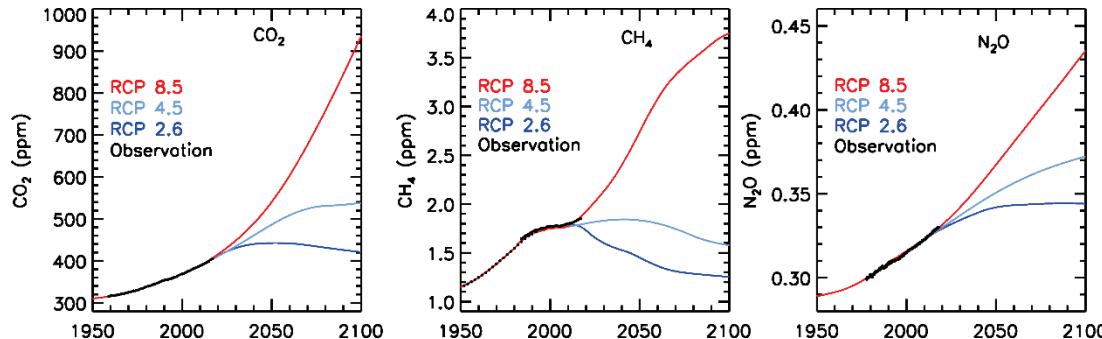
EM-GC Forecast vs CMIP5



If GHGs follow RCP 8.5, **0% chance** rise GMST stays below **1.5°C** and **0% chance** stays below **2.0°C**

Hope et al., *Earth's Future*, submitted, 2020.

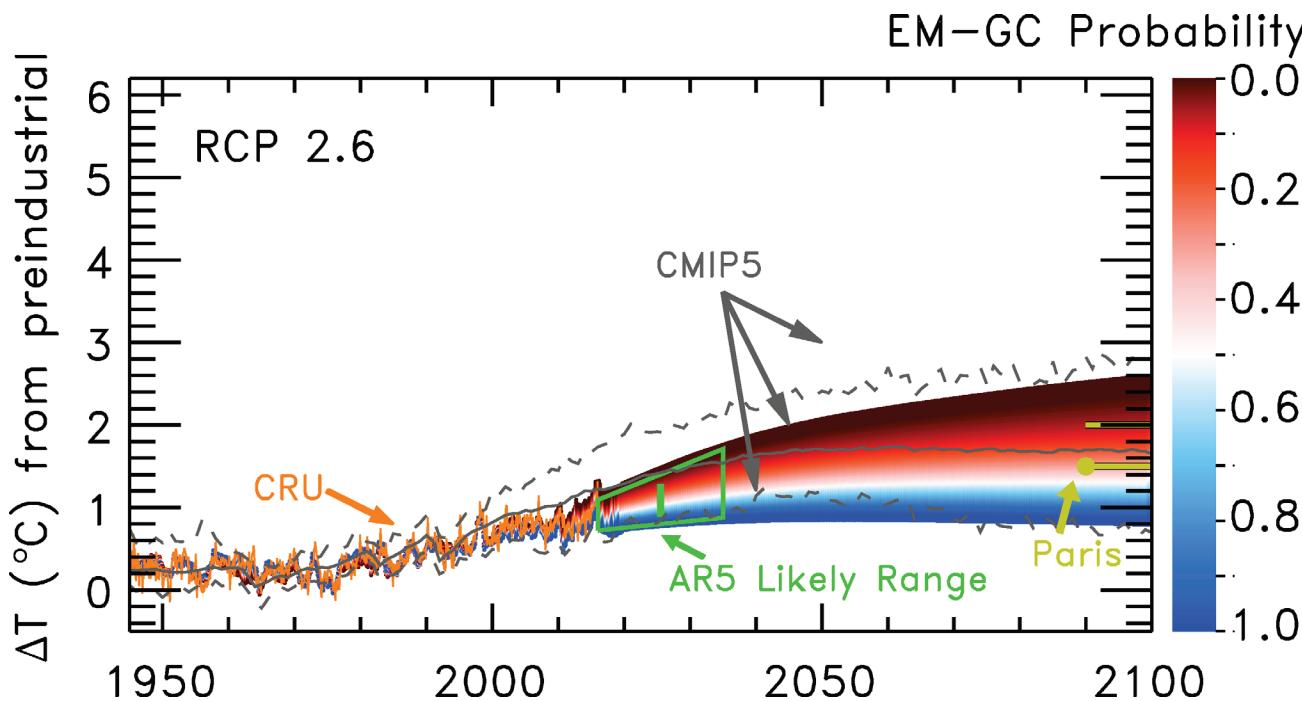
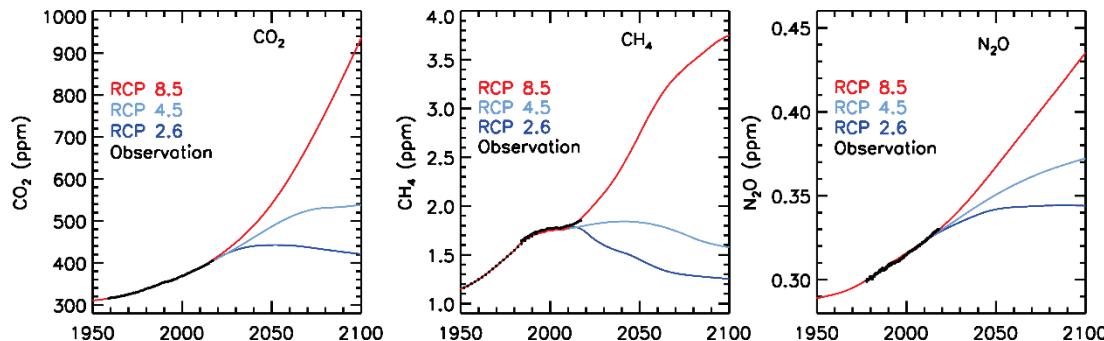
EM-GC Forecast vs CMIP5



If GHGs follow RCP 4.5, 10% chance rise GMST stays below 1.5°C and 50% chance stays below 2.0°C

Hope et al., *Earth's Future*, submitted, 2020.

EM-GC Forecast vs CMIP5



If GHGs follow RCP 2.6, **67%** chance rise GMST stays below **1.5°C** and **92%** chance stays below **2.0°C**

Hope et al., *Earth's Future*, submitted, 2020.