

# Fundamentals of Earth's Atmosphere

## AOSC / CHEM 433 & AOSC 633

Ross Salawitch & Walt Tribett

Class Web Site:

<http://www.atmos.umd.edu/~rjs/class/spr2019>

Email:

Ross: [rsalawit@umd.edu](mailto:rsalawit@umd.edu) or [rjs@atmos.umd.edu](mailto:rjs@atmos.umd.edu)

Walt: [wtribett@umd.edu](mailto:wtribett@umd.edu)

### Goals:

- 1) Tie up loose ends from last lecture
- 2) Barometric law (pressure vs height)
- 3) Thermal structure (temperature vs height)
- 4) Geostrophy (balance of pressure force & Coriolis Force  $\Rightarrow$  storms)
- 5) Ferrel Cell (mean circulation Earth's atmosphere  $\Rightarrow$  climate regimes)

## Lecture 3

7 February 2019

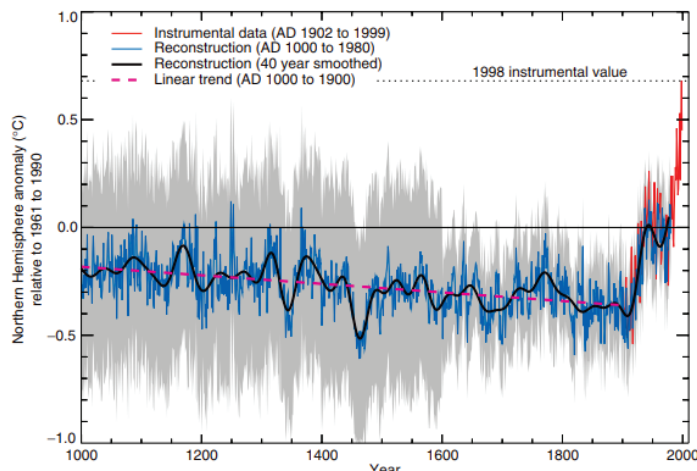
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1

In class, I had shown a version of the figure from a National Academies of Sciences report entitled *Surface Temperature Reconstructions for the Last 2000 Years*, published in 2006.

The original source of this figure is Chapter 2 of the 2001 IPCC report:



**Figure 2.20:** Millennial Northern Hemisphere (NH) temperature reconstruction (blue) and instrumental data (red) from AD 1000 to 1999, adapted from Mann *et al.* (1999). Smoother version of NH series (black), linear trend from AD 1000 to 1850 (purple-dashed) and two standard error limits (grey shaded) are shown.

<https://archive.ipcc.ch/ipccreports/tar/wg1/pdf/TAR-02.PDF>

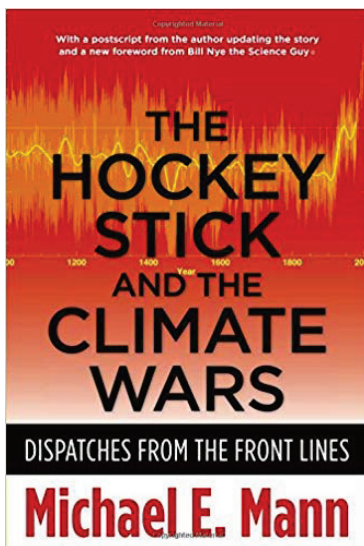
Mann *et al.*, *GRL*, 1999: <https://agupubs.onlinelibrary.wiley.com/doi/10.1029/1999GL900070>

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Lots more info in:



<https://www.amazon.com/Hockey-Stick-Climate-Wars-Dispatches/dp/0231152558>

This is a great book to get an understanding of the political war being waged over the climate change issue today, as well as a bit about how science works. It is not intended to explain climate science, although it does cover some of the techniques used in putting together an estimate of earth's surface temperature over the past 1000 years (the so-called "hockey stick"). But that is not the major part of either the book or climate science. The hockey stick is important not for scientific reasons but because it has become an icon, conveying visually that the stable climate that fostered the creation of human civilization supporting 7 billion people has undergone a radical change in the last 100 years. And there is no end in sight. First paragraph of top Amazon review by **Wendy Fehlauer**

Book published in 2012 so lacks salacious details of the recent political war being waged over climate change.

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## Tropospheric Pollutants (The Air We Breathe)

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Table 1.2 U.S. National Ambient Air Quality Standards		
Pollutant	Standard (ppm)	Approximate Equivalent Concentration ( $\mu\text{g}/\text{m}^3$ )
<b>Carbon monoxide</b>		
8-hr average	9	10,000
1-hr average	35	40,000
<b>Nitrogen dioxide</b>		
Annual average	0.053	100
<b>Ozone</b>		
8-hr average	0.075	147
1-hr average	0.12	235
<b>Particulates*</b>		
PM <sub>10</sub> , annual average	—	50
PM <sub>10</sub> , 24-hr average	—	150
PM <sub>2.5</sub> , annual average	—	15
PM <sub>2.5</sub> , 24-hr average <sup>†</sup>	—	35
<b>Sulfur dioxide</b>		
Annual average	0.03	80
24-hr average	0.14	365
3-hr average	0.50	1,300

\*PM<sub>10</sub> refers to all airborne particles 10  $\mu\text{m}$  in diameter or less. PM<sub>2.5</sub> refers to particles 2.5  $\mu\text{m}$  in diameter or less.

—The unit of ppm is not applicable to particulates.

<sup>†</sup>PM<sub>2.5</sub> standards are likely to be revised after 2011.

Source: U.S. Environmental Protection Agency. Standards also exist for lead, but are not included here.

Chapter 1, Chemistry in Context

### AT3, Q1:

What was the U.S. NAAQS for exposure to ozone over an 8 hr period of time in 2012, the time the 7<sup>th</sup> edition of *Chemistry in Context* was published, in units of parts per million?

**0.075 ppm**

### AT3, Q2:

Express the answer to Qu 1 in units of parts per billion.

**$(0.075 \text{ parts} / 1 \times 10^6) \times (1 \times 10^9) = 75 \text{ ppb}$**

### AT3, Q3:

Based on your own internet research, what is the current U.S. NAAQS standard for ozone?

**2015 EPA NAAQS Standard for O<sub>3</sub> is 0.070 ppm or 70 ppb**

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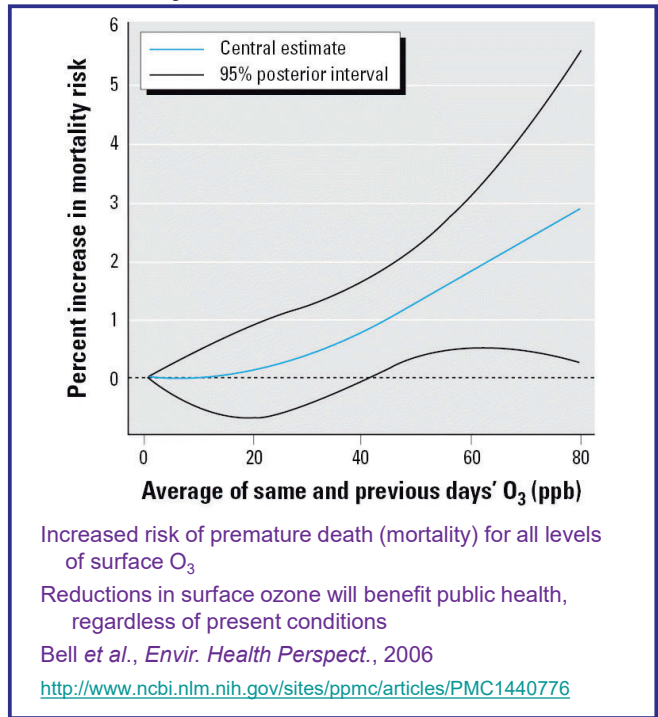
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# Air Quality Standards and Why We Care

Year	Averaging Period	EPA Surface Ozone Standard
1979	1 hr	125 ppb
1997	8 hr	85 ppb
2008	8 hr	75 ppb
2015 <sup>#</sup>	8 hr *	70 ppb

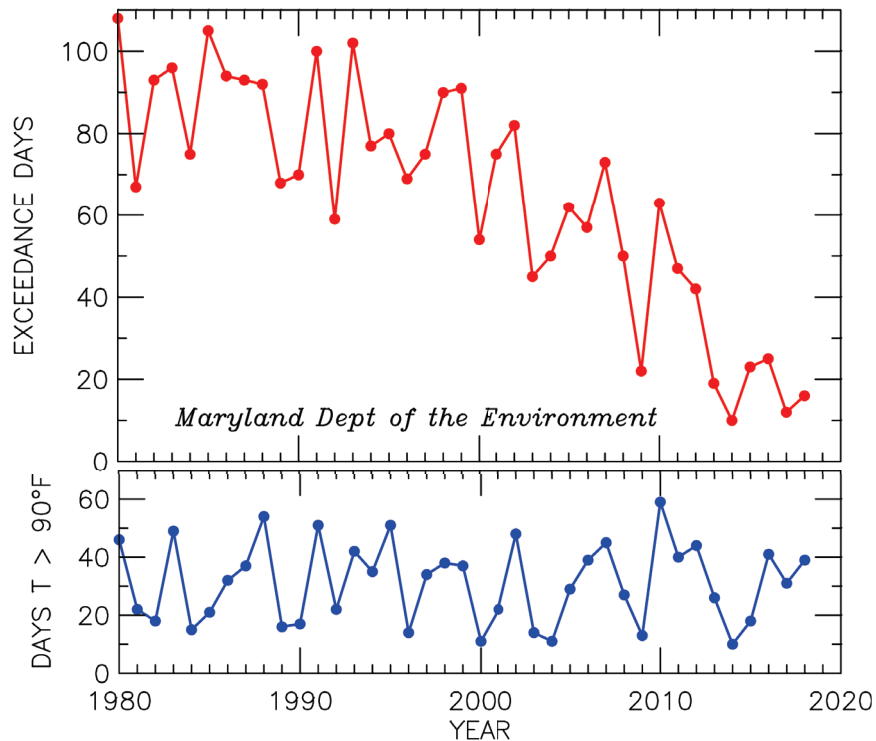
\* The 8 hr standard is met when the 3-yr average of the annual 4<sup>th</sup> highest daily maximum 8 hr O<sub>3</sub> is less than 70 ppb



<sup>#</sup> On October 1, 2015 the EPA lowered the NAAQS for ground-level ozone 70 ppb, based on extensive scientific evidence about the harmful effects of tropospheric ozone

## Significant Improvements in Local Air Quality since early 1980s

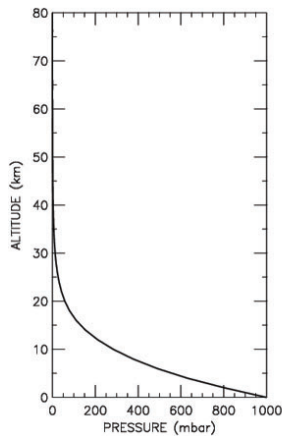
Days Exceeding New EPA Standard for 8 hr Ozone of 70 ppb (enacted in 2015) in Maryland



<http://www.mde.state.md.us/programs/Air/AirQualityMonitoring/Pages/SeasonalReports.aspx>



# Pressure versus Altitude



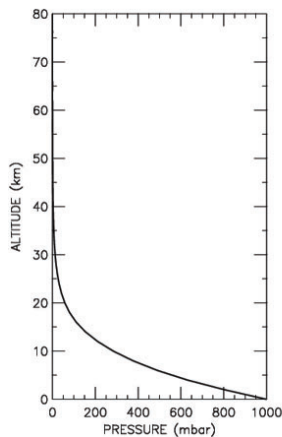
- **Pressure = Force per unit area**
- **Graph shows how “force” of atmosphere varies as a function of altitude**
- **Pressure shown in units of mbar : 1 mbar =  $10^3$  dynes/cm<sup>2</sup>**
- **1 dyne = gm cm / sec<sup>2</sup>; therefore 1 mbar =  $10^3$  gm / cm sec<sup>2</sup>**
- **Also:**
  - **European community prefers to write hPa; 1 hPa is exactly equal to 1 mbar**
  - **1 atmosphere =  $p/p_{\text{STANDARD}}$ , where  $p_{\text{STANDARD}} = 1013.25$  mbar (or 1013.25 hPa)**

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# Pressure versus Altitude



Derivation of the Barometric Law involves use of the Ideal Gas Law:

$$p \text{ Vol} = n R T$$

where  $p$  is pressure,  $\text{Vol}$  is volume,  $n$  is the number of moles of a gas,

$R$  is the gas constant ( $8.3143 \times 10^7 \frac{\text{ergs}}{\text{K mole}}$ ), and  $T$  is temperature

as shown on Extra Slide 1 at end of Class Notes (you will not be tested on this!)

- **Barometric law describes the variation of Earth’s pressure with respect to altitude:**

$$\text{Pressure (z)} = \text{Pressure (surface)} \times e^{-z/H}$$

where  $H$  is called the “scale height”

$$\text{Can show } H = R_{\text{EARTH}} T(z) / \text{grav},$$

$$\text{Since } R_{\text{EARTH}} = 2.88 \times 10^6 \text{ ergs} / \text{K gm}$$

$$\text{grav} = 981 \text{ cm sec}^{-2} \quad \text{and } T(\text{lower trop}) \approx 272 \text{ K}$$

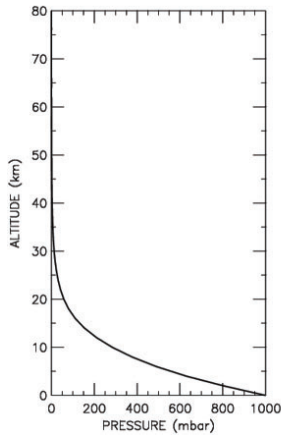
$$\text{then } H(z=0) = 8.0 \times 10^5 \text{ cm} = 8 \text{ km}$$

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# Pressure versus Altitude



In modern atmospheric sciences, the most handy version of the Ideal Gas Law is:

$$p = M k T$$

where  $p$  is pressure (force per unit area),  $M$  is number density (molecules/volume),  $k$  is Boltzmann's constant ( $1.38 \times 10^{-16}$  ergs/K), and  $T$  is temperature.

If  $p$  is given in units of mbar (or hPa),  $M$  is in units of  $\frac{\text{molecules}}{\text{cm}^3}$ , and  $T$  is in K, then can show  $k$  must be  $1.38 \times 10^{-19} \frac{\text{mbar cm}^3}{\text{K molecules}}$

- Barometric law describes the variation of Earth's pressure with respect to altitude:

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where  $H$  is called the "scale height"

$$\text{Can show } H = R_{\text{EARTH}} T(z) / \text{grav},$$

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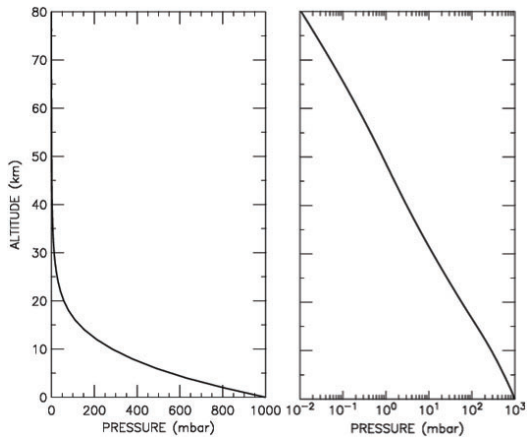
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# Pressure versus Altitude



- Barometric law describes the variation of Earth's pressure with respect to altitude:

$$\text{Pressure (z)} = \text{Pressure (surface)} \times e^{-z/H}$$

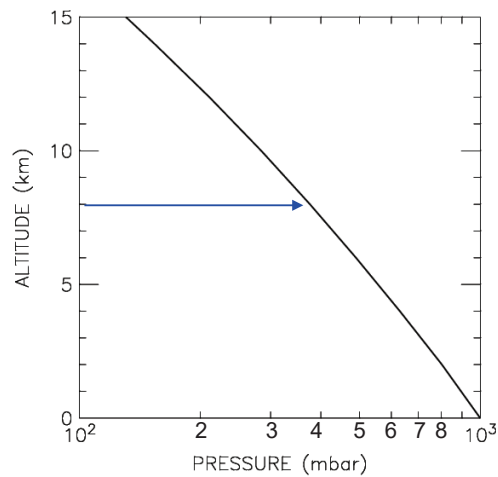
Two plots convey the same information !

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# Pressure versus Altitude



$$p(z=0) = 1013.25 \text{ mbar}$$

$$p(z=8 \text{ km}) =$$

- Barometric law describes the variation of Earth's pressure with respect to altitude:

$$\text{Pressure}(z) = \text{Pressure}(\text{surface}) \times e^{-z/H}$$

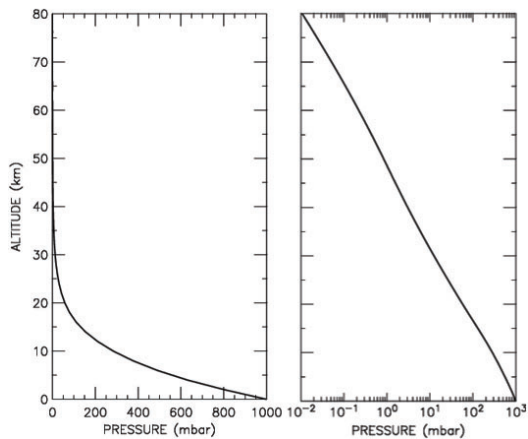
Let's take a closer look at log pressure versus altitude, in the troposphere

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# Pressure versus Altitude



- Barometric law describes the variation of Earth's pressure with respect to altitude:

$$\text{Pressure}(z) = \text{Pressure}(\text{surface}) \times e^{-z/H}$$

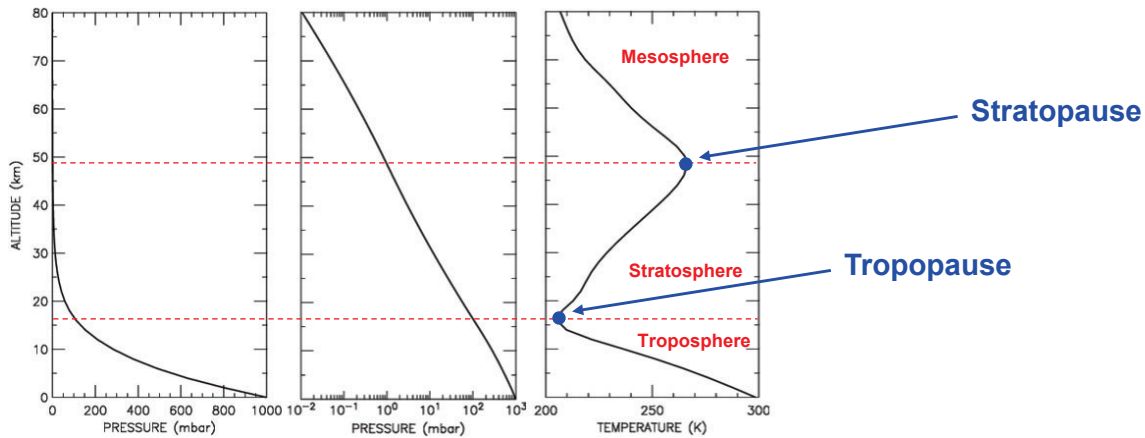
How does pressure vary as a function of depth, in the ocean ?

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# Temperature versus Altitude



- T falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

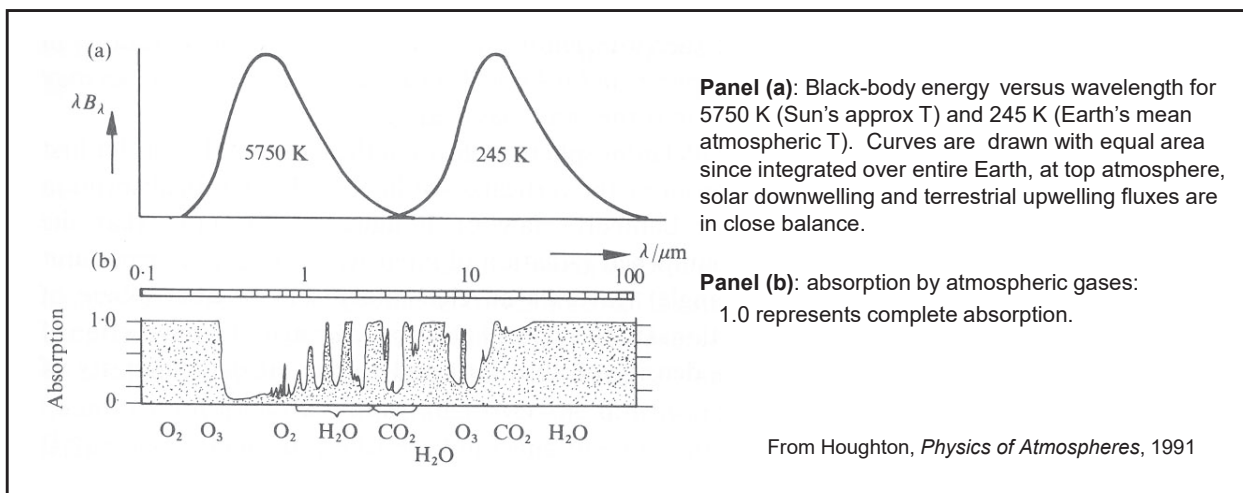
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# Atmospheric Radiation

- Solar irradiance (downwelling) at top of atmosphere occurs at wavelengths between ~200 and 2000 nm (~5750 K “black body” temperature)
- Thermal irradiance (upwelling) at top of the atmosphere occurs at wavelengths between ~5 and 50  $\mu\text{m}$  (~245 K “black body” temperature)



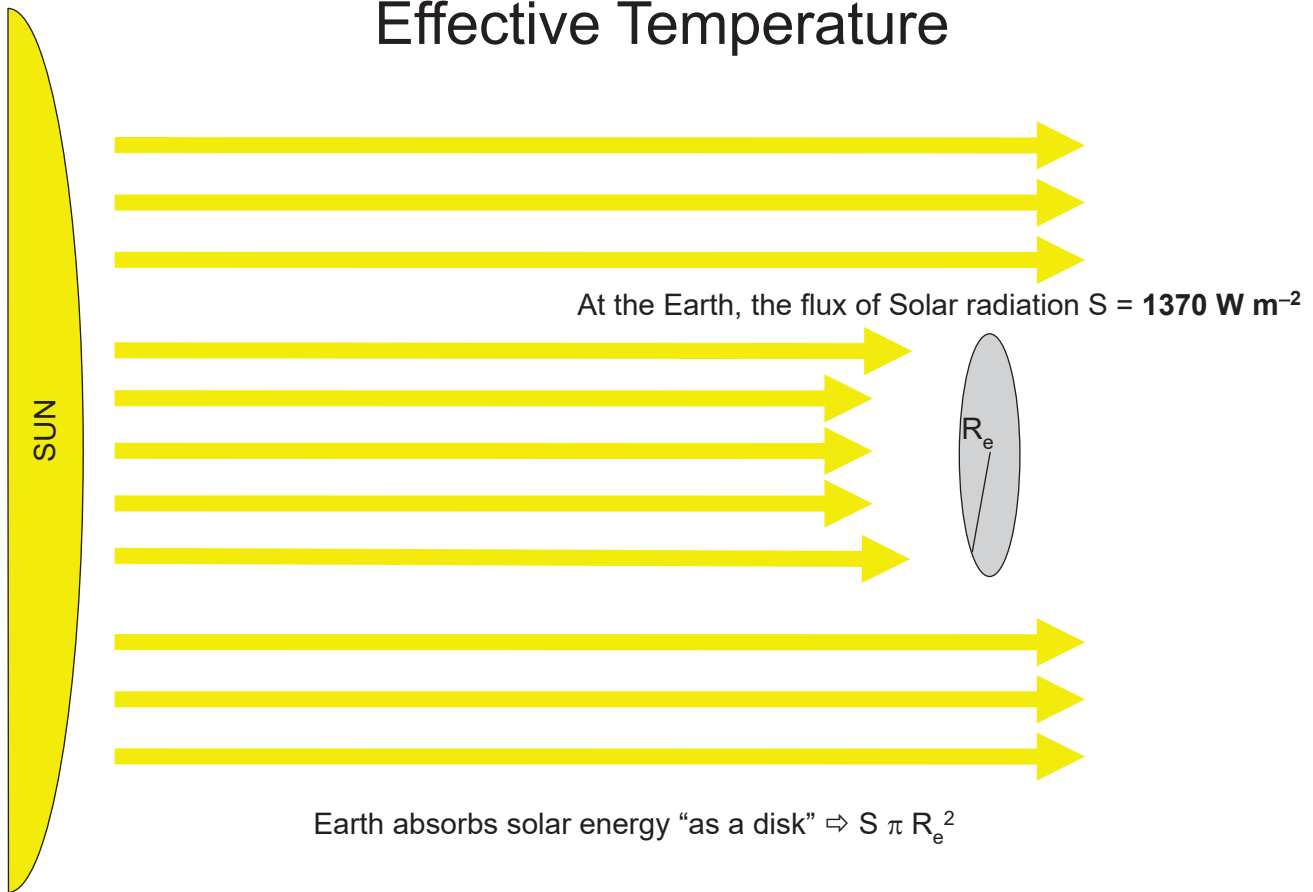
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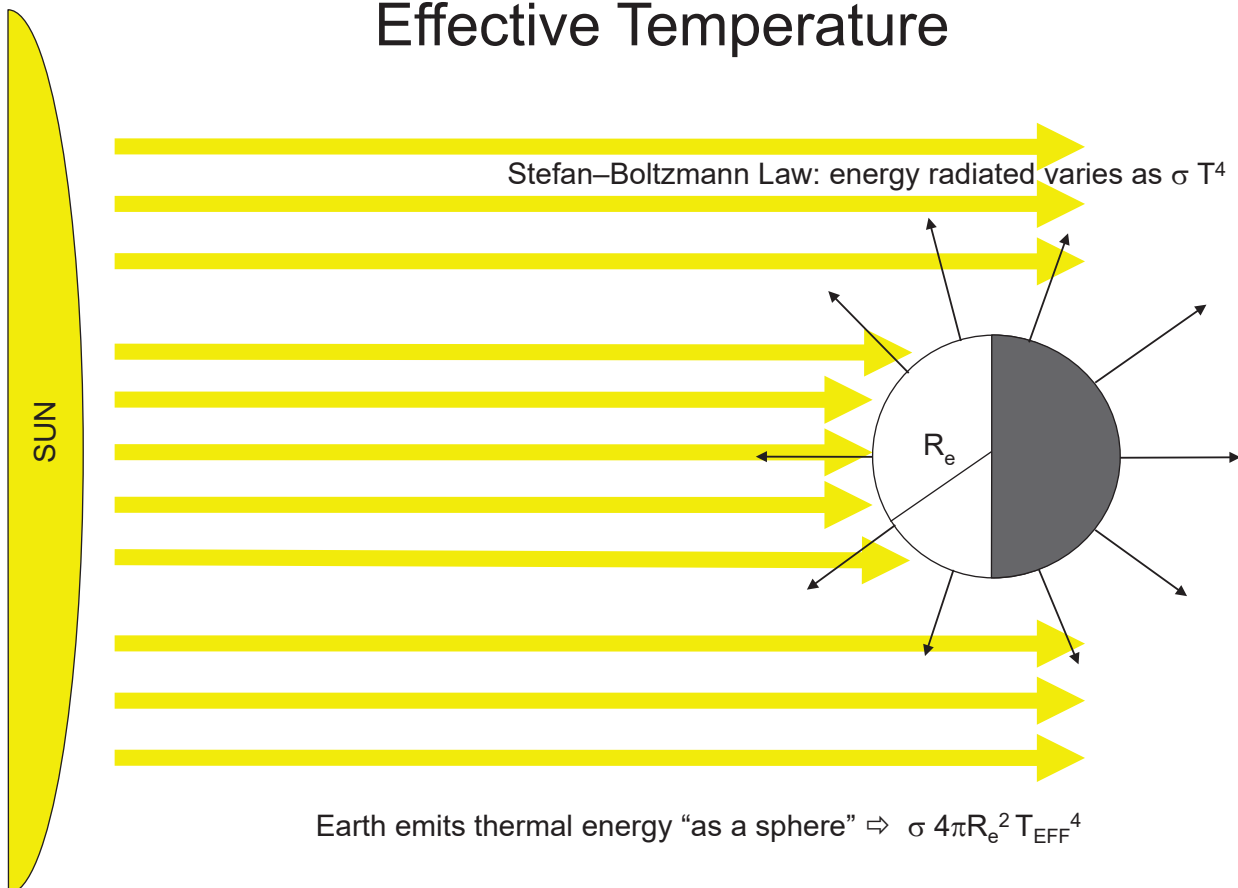


# Effective Temperature



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# Effective Temperature



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# Effective Temperature

Earth absorbs solar energy “as a disk”  $\Rightarrow (1 - \text{Albedo}) \times S \pi R_e^2$

Earth emits thermal energy “as a sphere”  $\Rightarrow \sigma 4\pi R_e^2 T_{\text{EFF}}^4$

$$(1 - \text{Albedo}) \times S = 4 \sigma T_{\text{EFF}}^4$$

or

$$T_{\text{EFF}} = \{ (1 - \text{Albedo}) \times S / 4 \sigma \}^{1/4}$$

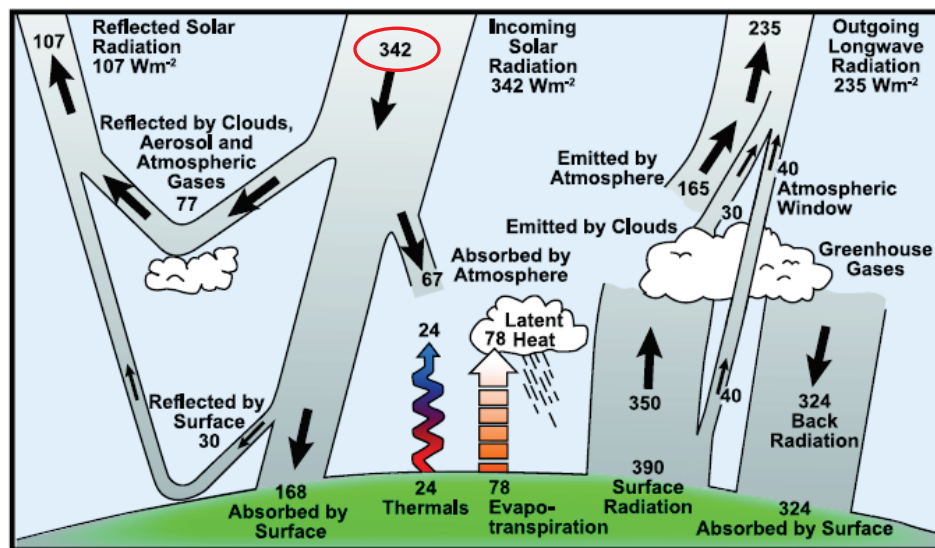
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# Effective Temperature

Let's take a closer look at  $S = 1370 \text{ W m}^{-2}$



FAQ 1.1, Figure 1. Estimate of the Earth's annual and global mean energy balance. Over the long term, the amount of incoming solar radiation absorbed by the Earth and atmosphere is balanced by the Earth and atmosphere releasing the same amount of outgoing longwave radiation. About half of the incoming solar radiation is absorbed by the Earth's surface. This energy is transferred to the atmosphere by warming the air in contact with the surface (thermals), by evapotranspiration and by longwave radiation that is absorbed by clouds and greenhouse gases. The atmosphere in turn radiates longwave energy back to Earth as well as out to space. Source: Kiehl and Trenberth (1997).

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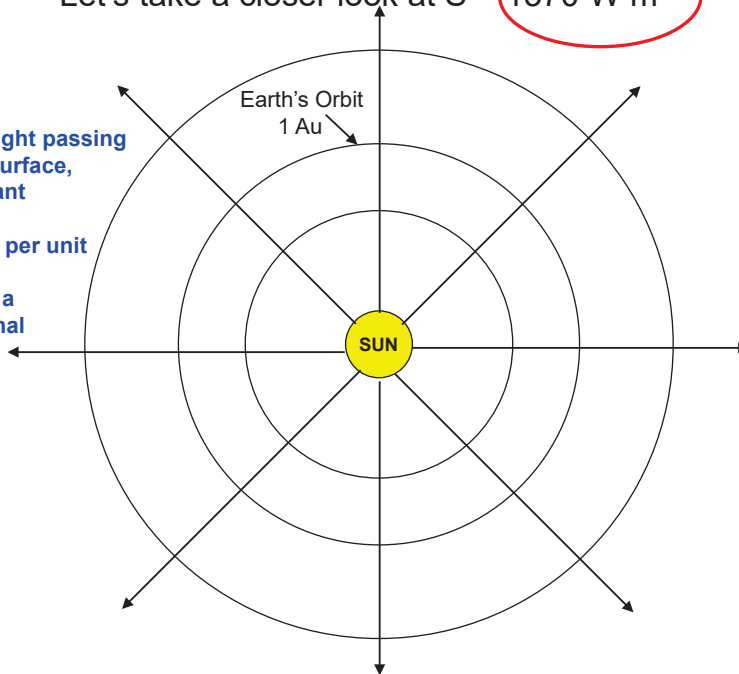
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# Effective Temperature

Let's take a closer look at  $S = 1370 \text{ W m}^{-2}$

The total amount of sunlight passing through each spherical surface, of various radii, is constant

Therefore the energy (W) per unit area ( $\text{m}^{-2}$ ) decreases wrt distance from the Sun in a manner that is proportional to: \_\_\_\_\_



Notes: 1) Au, or Astronomical Unit, is a measure of the distance of a planet from the Sun, normalized by the mean distance of Earth from the Sun. So by definition, **Earth's orbit is 1 Au from the Sun**

2) The diagram above represents orbits as perfect spheres, which is suitable for our study of effective temperatures. In reality, of course, planets orbit the Sun in an elliptical manner.

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$$T_{\text{EFF}} = \left\{ (1 - \text{Albedo}) \times S / 4 \sigma \right\}^{1/4}$$

633 students: find  $T_{\text{EFF}}$  for Earth, using:

$$\begin{aligned} \sigma &= 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4} \\ S &= 1370 \text{ W m}^{-2} \\ \text{Albedo} &= 0.3 \end{aligned}$$

433 student whose last name begins with letters A-M:

Find  $T_{\text{EFF}}$  for **Mars** using:

$$\begin{aligned} \sigma &= 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4} \\ S &= 1370 \text{ W m}^{-2} \\ \text{Albedo} &= 0.17 \end{aligned}$$

**Distance from Sun = 1.5 AU**

433 student whose last name begins with letters N-Z:

Find  $T_{\text{EFF}}$  for **Venus** using:

$$\begin{aligned} \sigma &= 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4} \\ S &= 1370 \text{ W m}^{-2} \\ \text{Albedo} &= 0.75 \end{aligned}$$

**Distance from Sun = 0.72 AU**

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# Effective Temperature

## My Favorite Planets



Venus:

$$T_{\text{SURFACE}} \approx 753 \text{ K}$$

$$T_{\text{EFFECTIVE}} \approx ???$$

Earth:

$$T_{\text{SURFACE}} \approx 288 \text{ K}$$

$$T_{\text{EFFECTIVE}} \approx ???$$

Mars

$$T_{\text{SURFACE}} \approx 217 \text{ K}$$

$$T_{\text{EFFECTIVE}} \approx ???$$

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## Geological Evolution of Earth's Atmosphere: *Earth, Mars, and Venus*

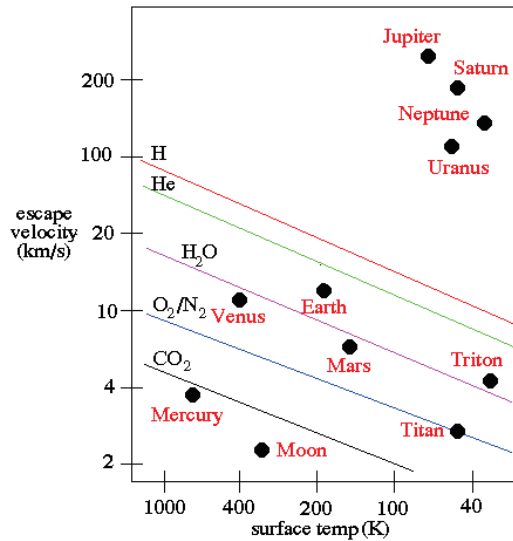
	Earth	Venus	Mars
Radius (km)	6400	6100	3400
Mass ( $10^{24}$ kg)	6.0	4.9	0.6
Albedo	0.3	0.8	0.22
Distance from Sun (A.U.)	1	0.72	1.52
Surface Pressure (atm)	<b>1</b>	<b>91</b>	<b>0.007</b>
Surface Temperature (K)	$\sim 15$ °C	$\sim 460$ °C	$-140$ °C to $20$ °C
N <sub>2</sub> (mol/mol)	0.78	$3.4 \times 10^{-2}$	$2.7 \times 10^{-2}$
O <sub>2</sub> (mol/mol)	0.21	$6.9 \times 10^{-5}$	$1.3 \times 10^{-3}$
CO <sub>2</sub> (mol/mol)	$3.7 \times 10^{-4}$	<b>0.96</b>	0.95
H <sub>2</sub> O (mol/mol)	$1 \times 10^{-2}$	$3 \times 10^{-3}$	$3 \times 10^{-4}$
SO <sub>2</sub> (mol/mol)	$1 \times 10^{-9}$	$1.5 \times 10^{-4}$	Nil
Cloud Composition	H <sub>2</sub> O	H <sub>2</sub> SO <sub>4</sub>	Mineral Dust

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# Geological Evolution of Earth's Atmosphere: *Earth. Mars. and Venus*



<http://abyss.uoregon.edu/~js/ast121/lectures/lec14.html>

Billions of years ago Venus had water. Most of this would have been in the form of vapor, since Venus is closer to the Sun. Water vapor is a greenhouse gas so heat from the Sun would have been trapped, temperature would have risen, and CO<sub>2</sub> would have been “baked” out of carbonate rocks. These circumstances likely led to increasing temperature in a strongly positive, upward spiraling fashion, leading to a “hot house” planet that persists today.

Scientists debate whether a runaway greenhouse effect could occur on Earth

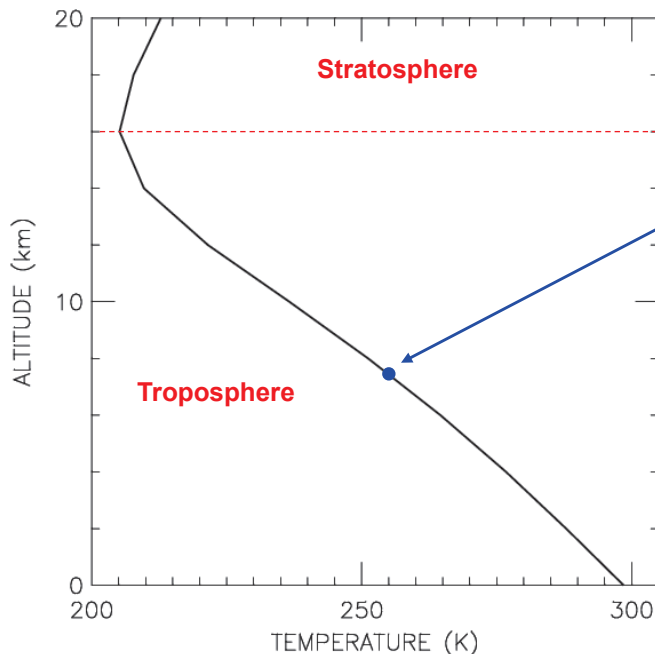
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## Temperature versus Altitude

Let's take a closer look at  $T_{\text{EFF}} = 255 \text{ K}$



Altitude in troposphere where  
 $T = 255 \text{ K}$   
Termed Earth's mean radiating  
altitude to space

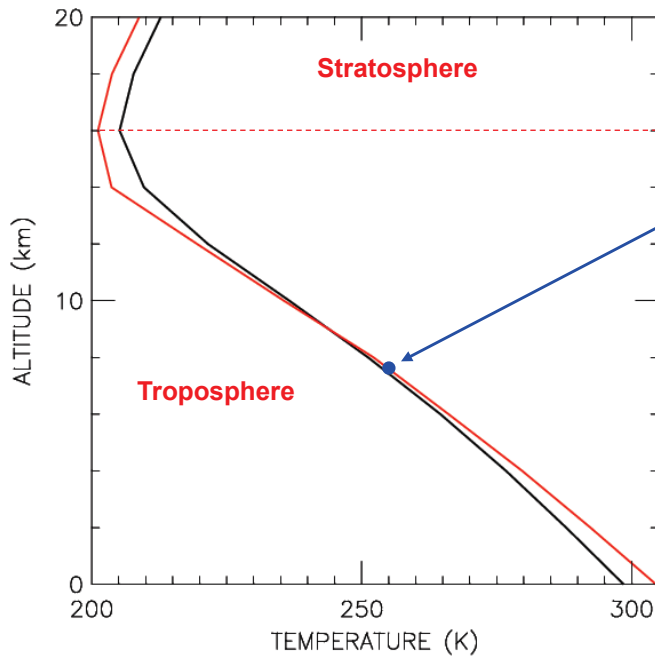
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# Temperature versus Altitude

Let's take a closer look at  $T_{\text{EFF}} = 255 \text{ K}$



Altitude in troposphere where  $T = 255 \text{ K}$

Termed Earth's mean radiating altitude to space

As Earth warms in response to rising GHGs, the lower troposphere will warm, the stratosphere will cool, and the mean radiating level will likely rise slightly higher in altitude

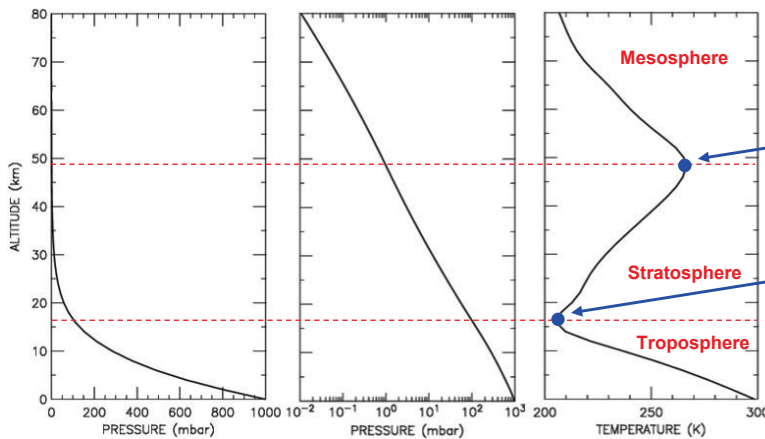
Regardless, the temperature of the mean radiating altitude will not change unless

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# Temperature versus Altitude



Stratopause

Tropopause

- $T$  falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

If the troposphere is dry,  $dT/dz = - \text{grav} / c_p$

where  $c_p$  is specific heat of air at constant pressure =  $1 \times 10^7 \text{ erg gm}^{-1} \text{ K}^{-1}$

Note:  $1 \text{ erg} = 1 \text{ dyne cm} = \text{gm cm}^2 \text{ sec}^{-2}$

$\Rightarrow dT/dz^{\text{DRY}} = - 981 \text{ cm sec}^{-2} / (10^7 \text{ cm}^2 \text{ sec}^{-2} \text{ K}^{-1}) \times 10^5 \text{ cm/km} = 9.8 \text{ K / km}$

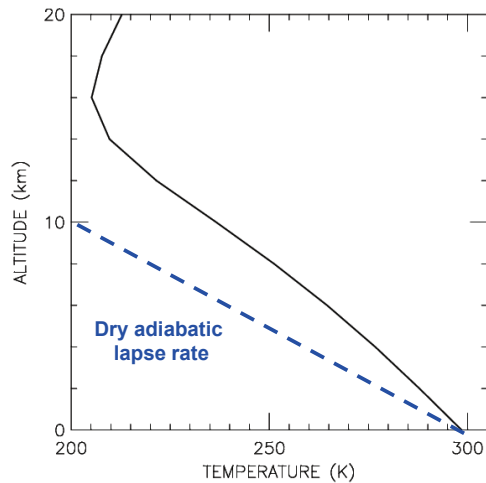
Dry adiabatic lapse rate

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# Temperature versus Altitude



- T falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

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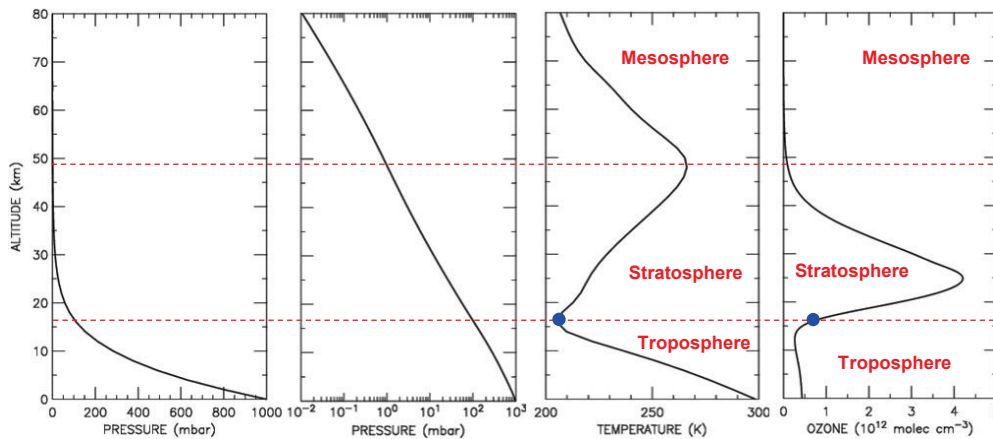
**Dry adiabatic lapse rate**

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# Temperature versus Altitude



- T falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

**Fourth chart expresses abundance of ozone concentration, or ozone density, or  $[O_3]$ , in units of molecules /  $\text{cm}^3$**

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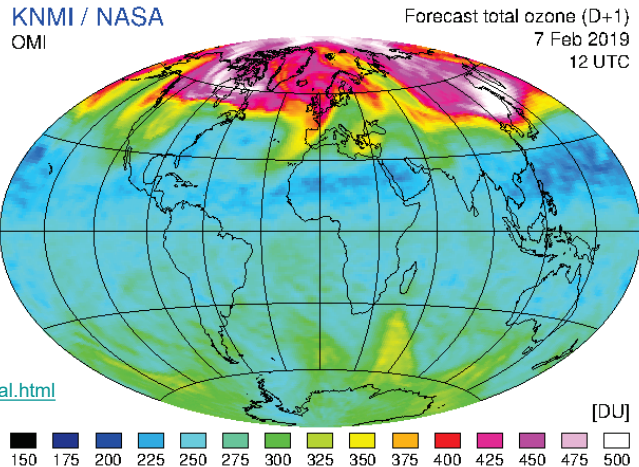
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# Back to the ATs

## AT3, Q4:

According to *Chemistry in Context*, if all of the ozone in Earth's atmosphere could be isolated and brought to the surface at a pressure of 1 atmosphere and a temperature of 15°C, the resulting gas would have a thickness of how many centimeters?



<http://www.temis.nl/protocols/O3global.html>

**1 Dobson Unit is defined to be a 0.01 millimeter thickness of air, at “standard temperature and pressure”**  
**Today we have about 300 Dobson Units (CC) = 300 × 0.01 millimeter = 3 millimeter or 0.3 cm of ozone, isolated and compressed to STP, between us and outer space.**

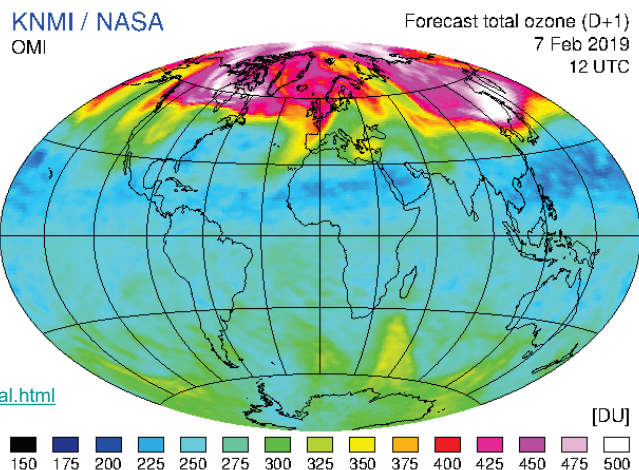
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# Back to the ATs

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<http://www.temis.nl/protocols/O3global.html>

### Mathematically:

$$\text{Ozone Column} = \int_{\text{Ground}}^{\text{Top of Atmosphere}} [\text{O}_3(z)] dz$$

Units :  $[\text{O}_3(z)]$  in  $\frac{\text{molecule}}{\text{cm}^3}$  and  $z$  in cm, leading to Ozone Column in  $\frac{\text{molecule}}{\text{cm}^2}$

It can be shown that  $1 \text{ DU} = 2.687 \times 10^{16} \frac{\text{molecule}}{\text{cm}^2}$

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# Back to the ATs

## AT3, Q4:

According to *Chemistry in Context*, if all of the ozone in Earth's atmosphere could be isolated and brought to the surface at a pressure of 1 atmosphere and a temperature of 15°C, the resulting gas would have a thickness of how many centimeters?

0.5 cm

## AT3, Q5:

if all of the molecules in the entirety of Earth's atmosphere could be isolated and brought to the surface (at a pressure of 1 atmosphere and a temperature of 15°C), the resulting gas would have a thickness of about  $7.4 \times 10^5$  cm

Compute the mean mixing ratio of ozone throughout Earth's atmosphere (this is a simple ratio of two numbers), **express the answer in parts per million.**

Mean mixing ratio of ozone:  $0.5 \text{ cm} / 7.4 \times 10^5 \text{ cm} = 6.8 \times 10^{-7}$  (no units!)

Note: mixing ratio has used throughout this class means number mixing ratio or:  $\frac{\text{moles of a particular gas}}{\text{moles of all gases in the air sample}}$

To convert  $6.8 \times 10^{-7}$  to ppm, need to realize 1 ppm corresponds to:  $\frac{1 \text{ moles of a particular gas}}{10^6 \text{ moles of all gases in the air sample}}$

Hence,  $6.8 \times 10^{-7} \times 1 \text{ ppm} / 10^6 = \mathbf{0.68 \text{ ppm}}$  \*OR\* **ppb**

Compare your answer to the NAAQS for O<sub>3</sub> given in your answer to either Q1 or Q3.

Specifically:

Is the mean mixing ratio of O<sub>3</sub> throughout Earth's atmosphere larger or smaller than the current U.S. NAAQS for O<sub>3</sub>?  
and

What figure in the assigned material from CC provides the "explanation" to the comparison of these two values?

# Back to the ATs

## AT3, Q4:

According to *Chemistry in Context*, if all of the ozone in Earth's atmosphere could be isolated and brought to the surface at a pressure of 1 atmosphere and a temperature of 15°C, the resulting gas would have a thickness of how many centimeters?

0.5 cm

## AT3, Q5:

if all of the molecules in the entirety of Earth's atmosphere could be isolated and brought to the surface (at a pressure of 1 atmosphere and a temperature of 15°C), the resulting gas would have a thickness of about  $7.4 \times 10^5$  cm

Compute the mean mixing ratio of ozone throughout Earth's atmosphere (this is a simple ratio of two numbers), **express the answer in parts per million.**

Mean mixing ratio of ozone:  $0.5 \text{ cm} / 7.4 \times 10^5 \text{ cm} = \underline{\hspace{2cm}}$  (no units!)

Note: mixing ratio has used throughout this class means number mixing ratio or:  $\frac{\text{moles of a particular gas}}{\text{moles of all gases in the air sample}}$

To convert  $6.8 \times 10^{-7}$  to ppm, need to realize 1 ppm corresponds to:  $\frac{1 \text{ moles of a particular gas}}{10^6 \text{ moles of all gases in the air sample}}$

Hence,  $6.8 \times 10^{-7} \times 1 \text{ ppm} / 10^6 = \underline{\hspace{1cm}} \text{ ppm}$  \*OR\*  $\underline{\hspace{1cm}} \text{ ppb}$

Compare your answer to the NAAQS for O<sub>3</sub> given in your answer to either Q1 or Q3.

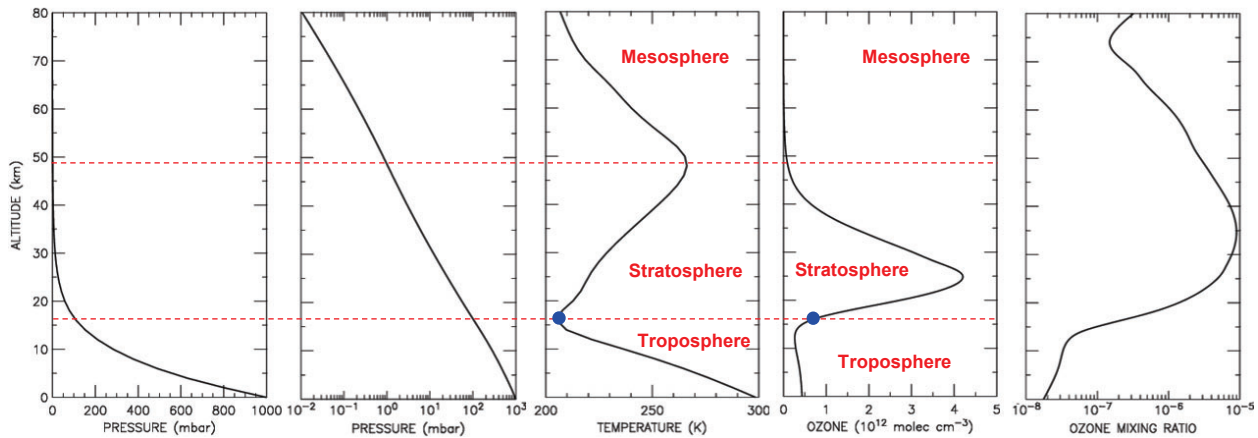
Specifically:

Is the mean mixing ratio of O<sub>3</sub> throughout Earth's atmosphere larger or smaller than the current U.S. NAAQS for O<sub>3</sub>?  
and

What figure in the assigned material from CC provides the "explanation" to the comparison of these two values?

**Mean mixing ratio of O<sub>3</sub> in Earth's Atmosphere >> than U.S. NAAQS**  
**Much higher mixing ratios of O<sub>3</sub> in the stratosphere than ever exists near the surface**

# Temperature versus Altitude



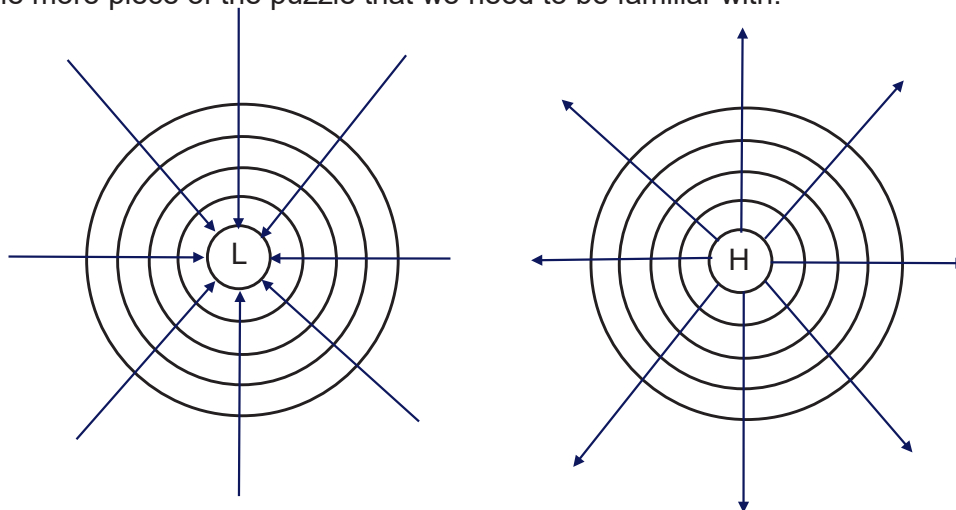
- T falls wrt increasing altit until the tropopause, then rises wrt altit until the stratopause, then falls wrt to rising altitude

Chart on far right expresses ozone mixing ratio,  $O_3$  mr in dimensionless units, where  $O_3$  mr =  $[O_3] / [M]$ , where  $[M]$  is the concentration (or density) of air

# Coriolis Force

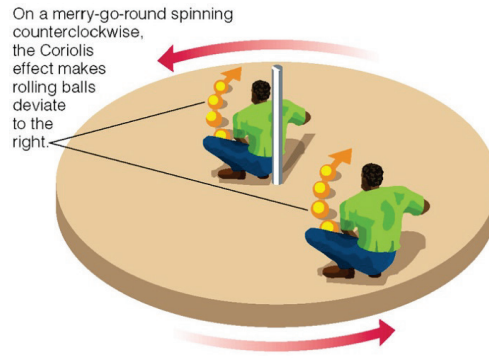
So far, we've reviewed temperature, pressure, and the balance between solar energy input to the atmosphere and terrestrial radiation leaving the atmosphere.

There's one more piece of the puzzle that we need to be familiar with.



In general, air moves from areas of high pressure to areas of low pressure. In the absence of external forces, air will move in a straight line, following pressure gradients

# Coriolis Force



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<http://lasp.colorado.edu/~bagenal/3720/CLASS15/15EVM-Dyn1.html>

Earth's rotation provides an apparent force that deflects air to the right in the Northern Hemisphere, to the left in the Southern Hemisphere.  
Force is proportional to  $\sin(\text{latitude})$ , so vanishes at the equator

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# Geostrophy

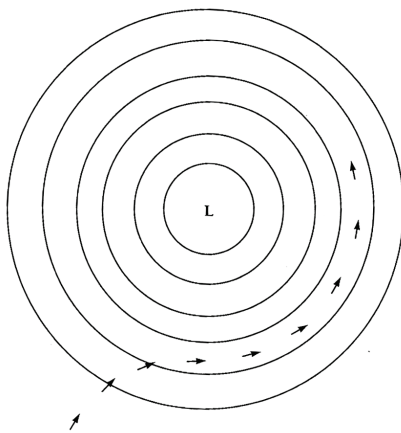


Figure 8.16 Track of an air parcel in the vicinity of a low pressure region in the Northern Hemisphere. The parcel is initially at rest but then adjusts to the pressure gradient force and the Coriolis force to achieve geostrophic balance.

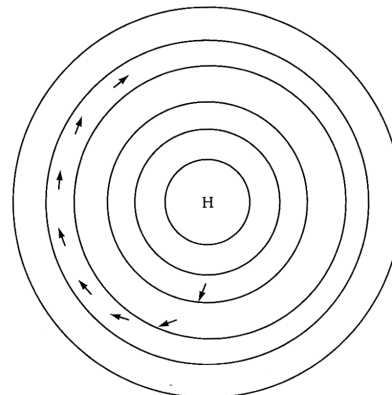


Figure 8.17 Same situation as in 8.16, except that the parcel is in the vicinity of a high pressure region in the Northern Hemisphere.

From "The Atmospheric Environment", M. B. McElroy

**Geostrophic balance: balance between Coriolis Force and pressure gradient**

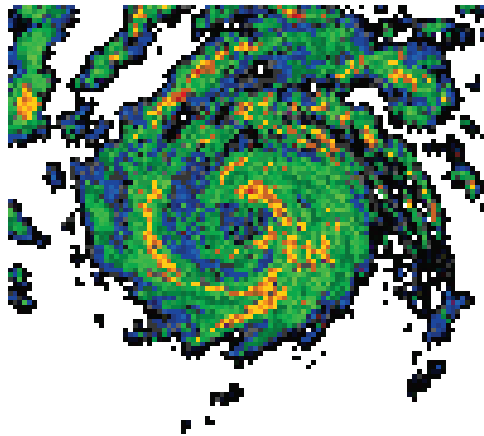
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# Geostrophy

## NH Weather System:

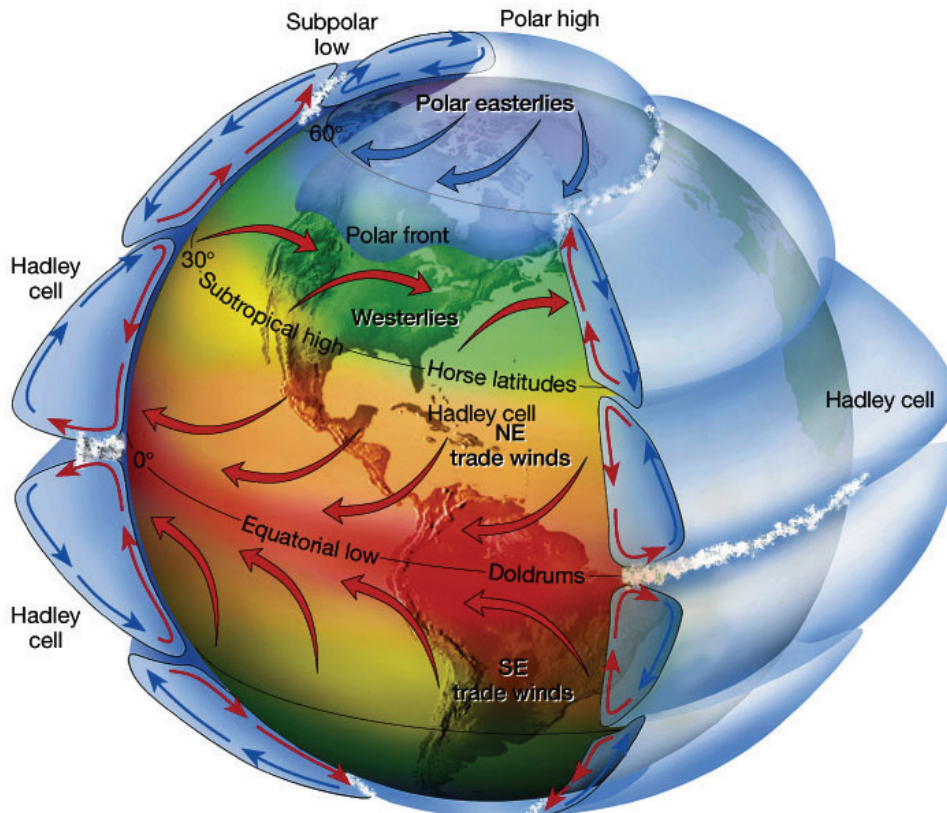


**Cyclonic Flow:** when the wind swirls counter-clockwise in the NH

**Hurricane:** Cyclonic flow occurring in the N Atlantic or NE Pacific Ocean east of the dateline.

**Typhoon:** Cyclonic flow occurring over the NW Pacific Ocean, west of the dateline.

## Ferrel Circulation (Modern View)



# Next Lecture: Climates of the Past

## Next Reading:

*Chemistry in Context*, Secs 2.2, 3.0, 3.1, 3.2 (~14 pgs)

as well as 8 pages from *Global Warming: The Complete Briefing* by Houghton  
7 pages from *Paris Beacon of Hope*

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