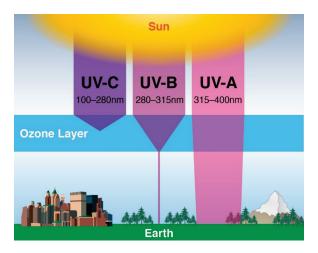
Introduction to Atmospheric Photochemistry

AOSC / CHEM 433 & AOSC 633

Ross Salawitch & Walter Tribett

Class Web Site: http://www.atmos.umd.edu/~rjs/class/spr2019



Lecture 10 12 March 2019

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Admission Ticket Lecture 10

According to Section 2.6 of Chemistry in Context:

- a) what is the upper limit for photodissociation of O₂
- b) what is the classification of this type of radiation?
- c) how damaging is this radiation and would it be a problem, for humans to be exposed to this type of radiation?

According to Section 2.6 of Chemistry in Context:

- a) what is the upper limit for photodissociation of O₃
- b) what is the classification of this type of radiation?
- c) how damaging is this radiation and would it be a problem, for humans to be exposed to this type of radiation?

Figure 2.12 of the Warneck reading shows the photodissocation *frequency* of O_2 and O_3 , termed J_{O2} and J_{O3} , as a function of altitude. An altitude of 15 km corresponds to Earth's upper troposphere. And, the lifetime for loss by dissociation is given by the reciprocal of the photodissociation frequency.

Calculate the lifetimes for loss of $\rm O_2$ and for loss of $\rm O_3$ by photodissociation, for the upper troposphere, in units that you can easily conceptualize. Then, for each molecule, state whether you think the gas will be lost by photodissociation in Earth's upper troposphere.

Admission Ticket Lecture 10

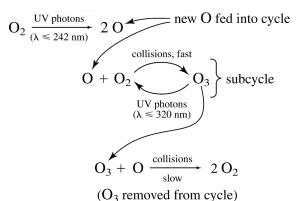


Figure 2.10, Chemistry in Context

$$O_2 \xrightarrow[(\lambda \le 242 \text{ nm})]{\text{UV photons}} 2 O$$

Equation 2.4, Chemistry in Context

$$O_3 \xrightarrow{\text{UV photon}} O_2 + O$$
 Equation 2.5, Chemistry in Context

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According to Section 2.6 of Chemistry in Context:

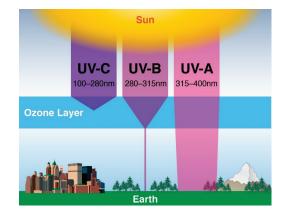
- a) what is the upper limit for photodissociation of O₂
- b) what is the classification of this type of radiation?
- c) how damaging is this radiation and would it be a problem, for humans to be exposed to this type of radiation?

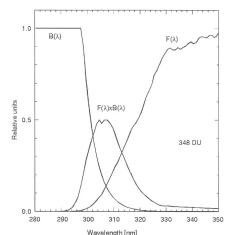
O₂ protects us even more so than O₃

According to Section 2.6 of Chemistry in Context:

- a) what is the upper limit for photodissociation of O_3
- b) what is the classification of this type of radiation?
- c) how damaging is this radiation and would it be a problem, for humans to be exposed to this type of radiation?

When we examine skin cancer and ozone depletion, it is all about UV-B radiation.

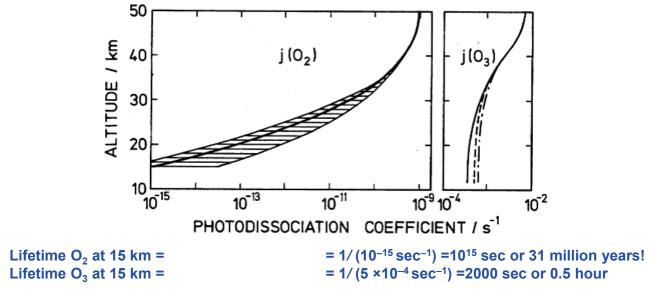




Admission Ticket Lecture 10

Figure 2.12 of the Warneck reading shows the photodissocation frequency of O₂ and O₃ termed J_{02} and J_{03} as a function of altitude. An altitude of 15 km corresponds to Earth's upper troposphere. And, the lifetime for loss by dissociation is given by the reciprocal of the photodissociation frequency.

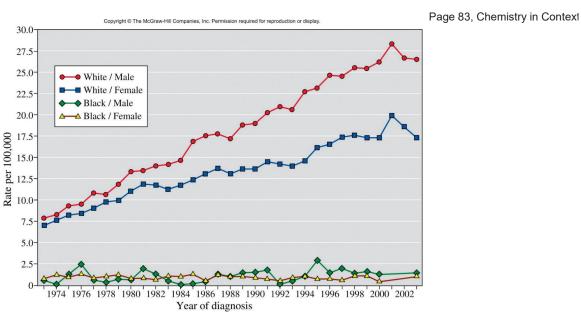
Calculate the lifetimes for loss of O₂ and for loss of O₃ by photodissociation, for the upper troposphere, in units that you can easily conceptualize. Then, for each molecule, state whether you think the gas will be lost by photodissociation in Earth's upper troposphere.



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Skin Cancer versus Time

According to calculations, a given percent decrease in stratospheric ozone is expected to increase the biological damage done by UV radiation by twice that percentage. For example, from a 6% decrease in stratospheric ozone we predict a 12% rise in skin cancer, especially the more easily treated forms such as basal cell and squamous cell carcinomas (non-melanoma)



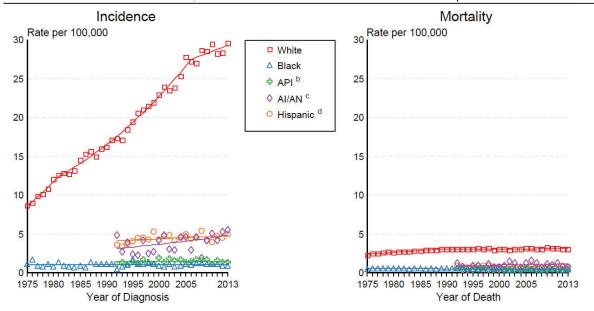
Increase in incidence of melanoma skin cancer in the U.S. Figure 2.12, Chemistry in Context

Skin Cancer versus Time, Update

Melanoma of the Skin, Both Sexes

Joinpoint Analyses for Whites and Blacks from 1975-2013

and for Asian/Pacific Islanders, American Indians/Alaska Natives and Hispanics from 1992-2013



https://seer.cancer.gov/csr/1975 2013/results merged/topic race ethnicity.pdf

SEER: Surveillance, Epidemiology, and End Results Program of the National Cancer Institute. As far as we can tell, census data is used to determine ethnicity.

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Relationship Between UV and Column Ozone

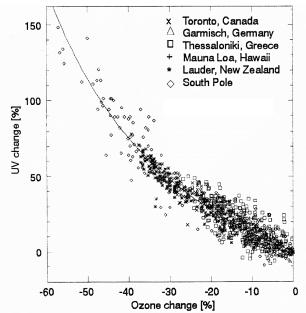
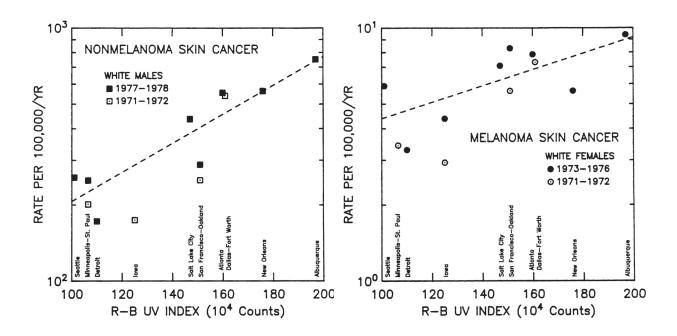


Fig. 2. Dependence of erythemal ultraviolet (UV) radiation at the Earth's surface on atmospheric ozone, measured on cloud-free days at various locations, at fixed solar zenith angles. Legend: South Pole [8]; Mauna Loa, Hawaii [9]; Lauder, New Zealand [10]; Thessaloniki, Greece (updated from Ref. [11]); Garmisch, Germany [12]; and Toronto, Canada (updated from Ref. [13]).

Madronich et al., J. of Photochemistry and Photobiology B, Vol. 46, 5–19, 1998.

Relationship Between Cancer and UV

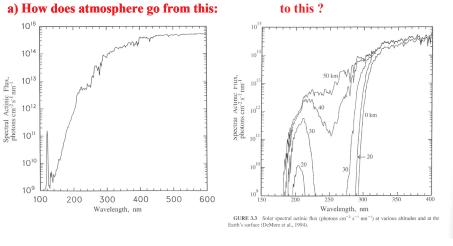


Scotto and Fraumeni, Cancer Epidemiology, W. B. Saunders and Co, Philadelphia, 1982.

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Motivation for Today's Lecture:



From DeMore et al., Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling, Evaluation No. 11, 1994.

From Seinfeld and Pandis, Atmospheric Chemistry and Physics, 1998.

b) Biological Effects of UV Radiation

Energetics of Photolysis

$$O_3 + hv \rightarrow O(^1D) + O_2$$

hy represents a photon with specific energy.

Let's examine enthalpy of this reaction:

33.9 kcal/mole 104.7 kcal/mole 22.5 kcal/mole
$$O_3 + h\nu \rightarrow O(^1D) + O_2(^1\Delta_q)$$
 Δ Enthalpy =

Photon Energy:

$$\varepsilon = \frac{hc}{\lambda} \implies \lambda_{\max} = \frac{hc}{\Delta \text{Enthalpy}}$$

For O_3 photo-dissociating to $O(^1D)$:

$$\lambda_{\text{max}} = \frac{hc}{\Delta \text{Enthalpy}} = \frac{2.85 \times 10^4 \text{ kcal/mole nm}}{\Delta \text{Enthalpy}} = -----=$$

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Energetics of Photodissociation

$$\begin{split} &O_3 + h\nu \rightarrow O(^1D) + O_2(^1\Delta_g) \quad \Delta \text{Enthalpy} = & \text{kcal/mole} \quad \lambda_{\text{max}} = & \text{nm} \\ &O_3 + h\nu \rightarrow O(^3P) + O_2(^3\Sigma_g) \quad \Delta \text{Enthalpy} = \text{25.7 kcal/mole} \quad \lambda_{\text{max}} = & \text{nm} \end{split}$$

Atomic oxygen: (Note: you will not be "responsible" for the material below on any exam (9)

Ground state – two unpaired electrons in the 2p orbitals: $(1s)^2(2s)^2(2p_1)^2(2p_2)^1(2p_3)^1$ Called ³P:

"3" represents 2S+1, where S is spin of all of the unpaired electrons.

There are 2 unpaired electrons, each with spin of $\frac{1}{2}$

Hence, S = 1 and 2S+1 = $3 \leftarrow$ spin angular momentum

P represents orbital angular momentum, found from an electron diagram of filled orbitals:

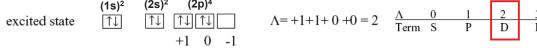
Excited state – one electron moves from $2p_3$ to $2p_2$: $(1s)^2(2s)^2(2p_1)^2(2p_2)^2$ Called ¹D:

"1" represents 2S+1, where S is spin of all of the unpaired electrons.

There are no unpaired electrons!

Hence, S = 0 and $2S+1 = 1 \leftarrow spin$ angular momentum

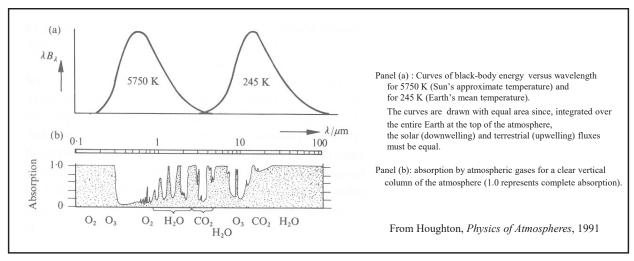
D represents orbital angular momentum, found from an electron diagram of filled orbitals:



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Atmospheric Radiation

• Solar irradiance (downwelling) at top of atmosphere occurs at wavelengths between ~200 and 2000 nm (~5750 K "black body" temperature)



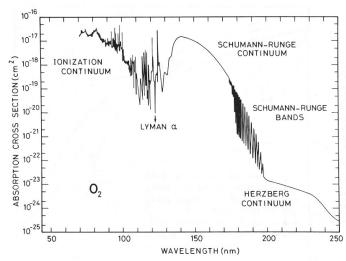
• Absorption and photodissociation in the UV occurs due to changes in the electronic state (orbital configuration) of molecules

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Absorption Cross Section of O₂



From Brasseur & Solomon, Aeronomy of the Middle Atmosphere, 1986

- O₂ can not dissociate longward of ~250 nm
- All absorption shown above is dissociative (e.g., leads to production of two O atoms)
- Structure in the O₂ cross section is related to whether the initial transition involves an unbound electronic state (smooth) or involves a specific vibrational level of an electronic state (banded, due to requirement of specific quanta of energy)

Beer-Lambert Law

$$F(z,\lambda) = F_{TOA}(\lambda) e^{-\tau(z,\lambda)}$$
 (TOA: Top of Atmosphere)

where:

$$\tau(z, \lambda) = m \int_{z}^{\infty} \sigma_{\lambda} [C] dz' \qquad (\tau: optical depth)$$

F : solar irradiance (photons/cm²/sec)

 σ_{λ} : absorption cross section (cm²/molecule)

C: concentration of absorbing gas (molecules/cm³)

m : ratio of slant path to vertical path, equal to $1/\cos(\theta)$ for $\theta < \sim 75^{\circ}$

 θ : solar zenith angle

Governs basics of radiative transfer in the UV and near IR regions

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Photolysis Frequency

For a specific spectral interval, the photolysis frequency (*partial J value*) of a gas is given by the product of its absorption cross section and the solar irradiance:

$$\begin{split} J_{gas}(z,\!\lambda) &= Quantum_Yield(\lambda) \; \sigma_{gas} \; (\lambda,\!T) \; F(z,\!\lambda) \\ &\quad \quad Units: \; s^{-1} \; nm^{-1} \end{split}$$

The total *photolysis frequency* (*J value*) is found by integrating $J_{gas}(z,\lambda)$ over all wavelengths for which the gas photodissociates:

$$J_{gas}(z) = \int_{\lambda_{min}}^{\lambda_{max}} J_{gas}(z, \lambda) d\lambda$$
 Units: s⁻¹

Rate of Reaction =
$$\frac{dO_3}{dt} = -J [O_3]$$
; Units of J are s⁻¹

More precisely, calculations of photolysis frequencies consider the "spectral actinic flux", which represents the amount of available photons integrated over all angles, rather than "solar irradiance". These two quantities differ because of scattering of solar radiation by gases and aerosols, and reflection of radiation by clouds and the surface.

Optical Depth of O₂ Absorption

Recall the Beer-Lambert Law:

$$F(z,\lambda) = F_{TOA}(\lambda) e^{-\tau(z,\lambda)}$$
 (TOA : Top of Atmosphere)

where:

$$\tau(z, \lambda) = m \int_{z}^{\infty} \sigma_{\lambda} [C] dz'$$
 (τ : optical depth)

Also:

$$\int_{0}^{\infty} [O_2] dz' \approx 4 \times 10^{24} \text{ molecules/cm}^2$$

$$\frac{O_{2} \text{ Optical Depth for } \theta = 0^{\circ}, z = 0 \text{ km}}{\sigma_{max}(cm^{2}) \quad \tau \left(0 \text{ km}\right) \quad e^{-\tau \left(0 \text{ km}\right)}}$$

$$\sigma_{\text{max}} (\text{cm}^2) \quad \tau (0 \text{ km}) \qquad e^{-\tau (0 \text{ km})}$$

Schumann-Runge Continuum Schumann-Runge Bands

Herzberg Continuum

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Photolysis Frequency of O₂

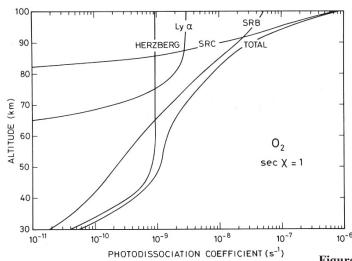


Figure 4.31, Brasseur and Solomon

Where Does Optical Depth = 1.0 for O_2 ?

$$\tau(z, \lambda) = m \int_{z}^{\infty} \sigma_{\lambda} [O_{2}] dz'$$
$$\approx \sigma_{\lambda} m 4 \times 10^{24} e^{-z/H}$$

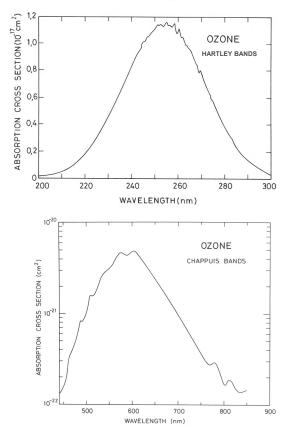
Setting $\tau = 1$ and re-arranging gives:

$$z = H \ln (\sigma_{\lambda} \cdot m \cdot 4 \times 10^{24})$$

Altitude where $\tau = 1$ (for $\theta = 0^{\circ}$)			
	$\sigma_{max}(cm^2)$ z (km)		
Schumann-Runge Continuum	10^{-17}		
Schumann-Runge Bands	10^{-20}		
	3×10^{-23}		
Herzberg Continuum	10^{-23}		

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Absorption Cross Section of O₃



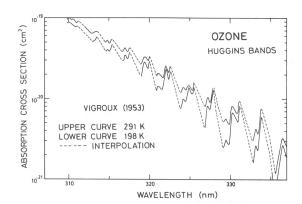


Table 4.6 Theoretical limits corresponding to different photolysis products (nm).

	$O_2(^3\Sigma_g)$	$O_2(^1\Delta_g)$	$O_2(^1\Sigma_g +)$	$O_2(^3\Sigma_u^{}+)$	$O_2(^3\Sigma_u^{\;-})$
$O(^{3}P)$	1180	590	460	230	170
$O(^{1}D)$	410	310	260	167	150
$O(^{1}S)$	234	196	179	129	108

From Brasseur & Solomon, Aeronomy of the Middle Atmosphere, 1986

Optical Depth of O₃ Absorption

A typical mid-latitude column abundance for O₃ is 300 Dobson units (DU):

$$1 \text{ DU} = 2.687 \times 10^{16} \text{ molecules/cm}^2$$
; $300 \text{ DU} = 8 \times 10^{18} \text{ molecules/cm}^2$

Aside:
$$\frac{\text{Column O}_3}{\text{Column Air}} = \frac{8 \times 10^{18}}{2 \times 10^{25}} = 0.4 \text{ parts per million } \Rightarrow \text{Ozone is a trace species!}$$

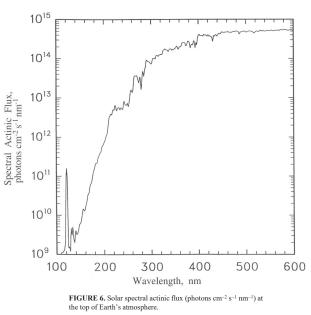
O_3 Optical Depth for $\theta = 0^\circ$, $z = 0$ km						
Hartley (~220 to 280 nm)	σ_{max} (cm ²) τ (0 km)	$e^{-\tau\;(\;0\;km)}$	O_3 Column, $\tau = 1.0$			
Huggins (~310 to 330 nm)						
Chappuis (~500 to 700 nm)						

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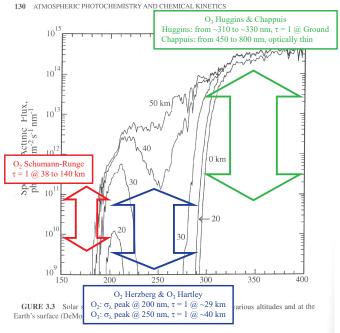
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Solar Spectral Actinic Flux



From DeMore et al., *Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling*, Evaluation No. 11, 1994.



From Seinfeld and Pandis, Atmospheric Chemistry and Physics, 1998.

Photodissociation Frequencies

Next goal is to understand:

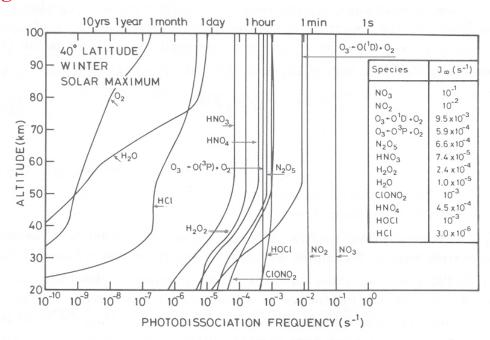


FIGURE 4.58 Photodissociation frequencies for numerous important atmospheric species.

From Brasseur & Solomon, Aeronomy of the Middle Atmosphere, 1986

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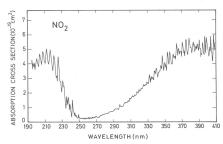
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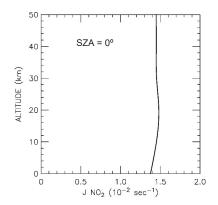
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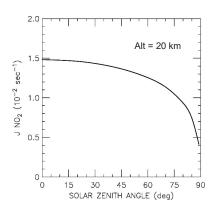
NO₂ Photolysis

The majority of NO_2 photolysis occurs longward of 300 nm, where the atmosphere is optically thin with respect to absorption by O_3 and O_2 :

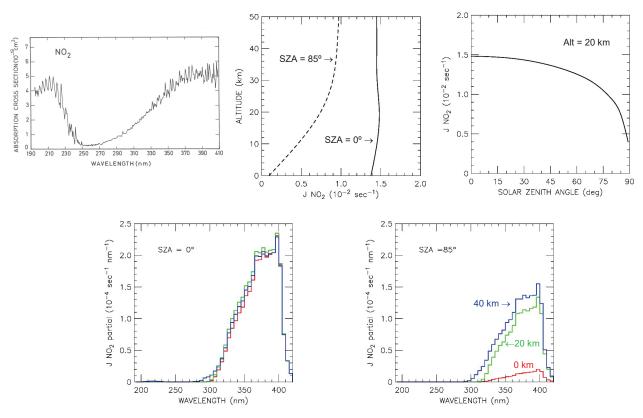


leading to a value for J_{NO2} that is nearly <u>independent</u> of height and SZA:





NO₂ Photolysis

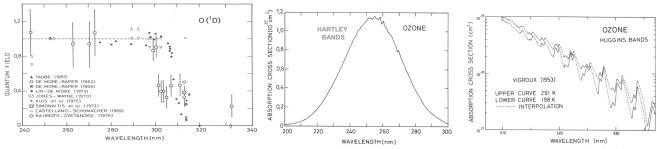


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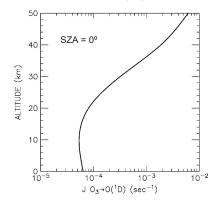
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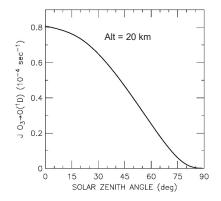
$O_3 \rightarrow O(^1D)$ Photolysis

The production of $O(^1D)$ from photolysis of O_3 occurs shortward of 320 nm, where the atmosphere is basically optically thick with respect to absorption by O_3 :

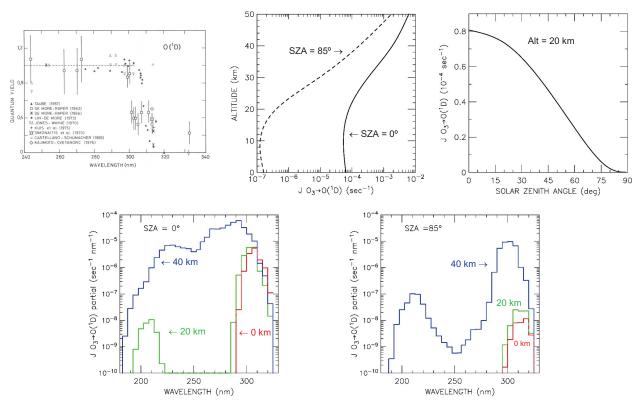


leading to a value for $J_{O3\to O(1D)}$ that is <u>dependent</u> on height and SZA:





$O_3 \rightarrow O(^1D)$ Photolysis

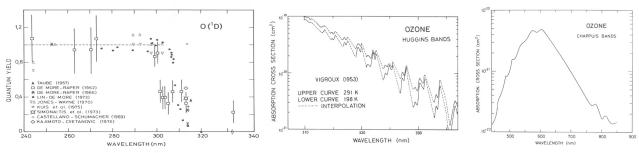


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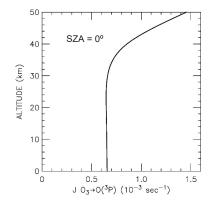
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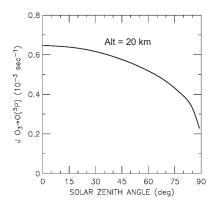
$O_3 \rightarrow O(^3P)$ Photolysis

The production of $O(^3P)$ from photolysis of O_3 occurs mainly longward of 500 nm, where the atmosphere is optically thin with respect to absorption by O_3 :

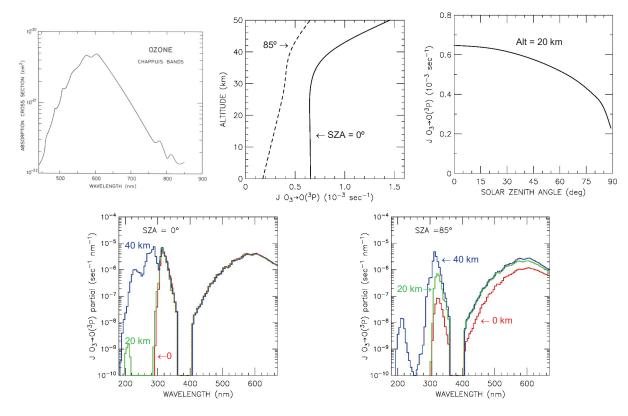


leading to a value for $J_{O3\to O(3P)}$ that is essentially $\underline{independent}$ of height and SZA:





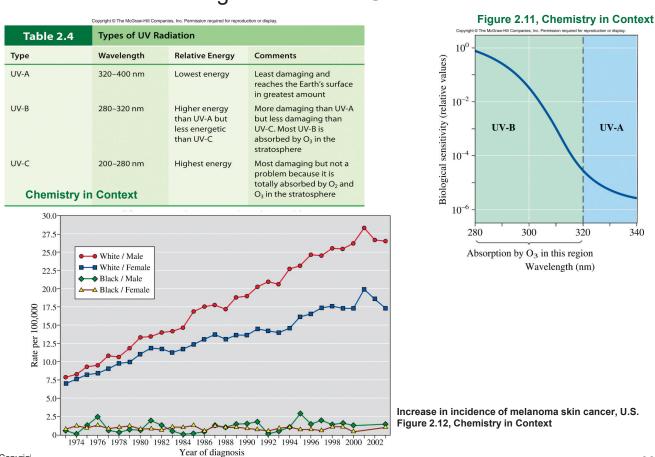
$O_3 \rightarrow O(^3P)$ Photolysis



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Biological Effects of UV Radiation



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Biological Effects of UV Radiation

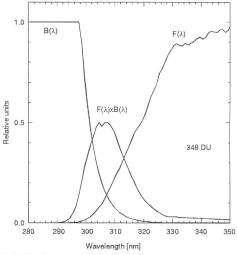


Fig. 1. Biologically active UV radiation. The overlap between the spectral irradiance $F(\lambda)$ and the erythemal action spectrum $B(\lambda)$ given by McKinlay and Diffey [6] shows the spectrum of biologically active radiation, $F(\lambda)B(\lambda)$. The area under the product function $F(\lambda)B(\lambda)$ is the biologically active dose rate. For a total ozone column of 348 DU.

Humans are:

- strongly affected by exposure to UV-C radiation (100 to 280 nm)
- moderately affected by exposure to UV-B radiation (280 to 315 nm)
- weakly affected by exposure to UV-A radiation (315 to 400 nm)

 $\underline{http://www.who.int/uv/uv_and_health/en}$

From Mandronich et al., J. Photochemistry and Photobiology, vol. 46, pg. 5, 1998

The "biologically active dose rate" maximizes in the UV-B region at ~305 nm, where σ_{03} = 3 × 10⁻¹⁹ cm² \Rightarrow τ (0 km) = 2.4 (for O₃ column= 300 DU)

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